

of the quadrangle.

the moderate topography.

## INTRODUCTION

The Cat Mountain 7.5-minute quadrangle is located northeast of the village of Tularosa on the eastern margin of the Tularosa Basin in south-central New Mexico. Main access is via U.S. Highway 70 through the southeastern corner of the area and by a maintained dirt road through Coyote Canyon in the central part of the quadrangle. Southwest-flowing Tularosa Creek and Temporal Creek are the most significant drainages. A mix of private, Bureau of Land Management and State stakeholders own or manage the land. A portion of the western edge of the Mescalero Apache Reservation occupies the northeast corner

Moore et al. (1988) mapped the northwestern part of the Mescalero Apache Reservation, including the northeastern corner of the Cat Mountain quadrangle at a 1:24,000 scale. We did not have permission to work on the reservation, so we used a slightly modified version of the Moore et al. (1988) map on the reservation land. Moore et al. (1988) originally lumped the Triassic Moenkopi Formation and and the Permian Grayburg Formation into a unit described as Triassic Santa Rosa Formation. We modified their "Triassic" contacts using Google Earth images and air photos. Otte (1959) mapped the southern quarter of the quadrangle at a scale of 4 inches to 1 mile (1:15,840). Otte (1959) focused his excellent and thorough work on the Pennsylvanian to Permian Labrocita Formation, which we correlate to the Bursum Formation in this report. Otte (1959) mapped the Labrocita Formation and the lower part of the overlying Abo Formation on a bed-by-bed basis using a plane table, creating a map of unprecedented detail. Kelley (1971) measured a complete section of the Yeso Formation in this area.

The quadrangle is in the transition between the east-dipping Sacramento Mountains and the Sierra Blanca volcanic field. In addition, this area is located on the southwestern margin of the Sierra Blanca basin, a north-northeast trending Laramide basin that contains sedimentary rocks as young as Eocene in age. The volcanic and plutonic rocks of the Sierra Blanca complex and Lincoln County porphyry belt were erupted onto and intruded into the Sierra Blanca basin ~26 to 38 million years ago. Countless dikes and sills associated with Sierra Blanca volcanism criss-cross the Cat Mountain region; the dikes follow a pronounced NE-trend. Rio Grande rift extensional faulting has since disrupted the Sierra Blanca basin and the volcanic centers. Rock exposures are generally quite good due to the sparse Chihuahuan desert vegetation and

	DESCRIPTION OF MAP UNITS QUATERNARY
Qal	<b>Quaternary alluvium (Holocene):</b> Unconsolidated gravel, sand, clay, and silt deposited in major modern drainage and tributary channels and associated floodplains. Locally contains small amounts of alluvial fan, colluvium and othe Quaternary units. up to 5 m thick.
Qc	<b>Undivided colluvium (Holocene)</b> : Hillslope colluvial deposits composed of pebble to boulder size clasts from a variety of lithologies, mainly limestone. Common on north facing slopes. Varies in thickness from ~ 5 – 50 m.
Qt	<b>Terrace deposits (Holocene):</b> Unconsolidated gravel, sand, and silt deposits along major present-day drainages inset into older alluvial deposits _ m above moderr grade. 1-2 m thick.
Qsw	<b>Slopewash deposits (Holocene):</b> Unconsolidated tan silt with lenses of gravel and sand that consists primarily of eolian material reworked by sheetflow. These deposits cover the floors of broad valleys in the central part of the quadrangle. These deposits commonly sit on well-consolidated older alluvium. 1-5 m thick.
Qls	<b>Landslide deposit (middle Pleistocene to Holocene):</b> Undivided landslide deposite composed of locally derived, cohesive blocks of Yeso Formation. Generally associated with a head scarp in their source area. Varies in thickness from ~ 5 – 50 m.
Qtr	<b>Travertine deposit (Pleistocene to Holocene):</b> Gray, porous limestone (tufa cementing alluvium on the south side of the collapse structure at the confluence o Dry Canyon and Rinconada Canyon. Thickness < 2m.
Qao <sub>2</sub>	<b>Old alluvium (Pleistocene):</b> Well-consolidated gravel, sand, and silt deposited by alluvial fans sourced from Black Mountain, located northeast of the quadrangle, and the northern foothills of the Sacramento Mountains. Gravel includes sedimentary (mainly limestone) and volcanic clasts. Red soil with carbonate stringers is well developed at the top of the unit and gypsiferous silt is usually present in the uppermost part of the alluvium. Qao2 is usually covered by a thin veneer of young alluvium or slopewash and is best observed in deeply incised arroyos. Surface developed on this deposit are near modern grade. Younger thin cut and fill deposite along Salinas Draw on the northern edge of the quadrangle are lumped with Qao2. Up to 15 m thick.
Qao1	<b>Older alluvium (Pleistocene):</b> Well-consolidated gravel, sand, and silt deposited by alluvial fans sourced in the northern foothills of the Sacramento Mountains on a high level strath cut on Permian Abo Formation in south-central part of the map area Strath surfaces beneath this deposit are 50 to 70 m above modern grade. The gravels are typically clast-supported, imbricated, subrounded to subangular, poorly sorted and consist of pebbles and cobbles. Gravel includes sedimentary (mainly limestone and volcanic clasts.Up to 25 m thick.
Ti	Undifferentiated dikes and sills mapped from air photos.
Tif	<b>Fine-grained trachyte (Eocene?):</b> Black to dark green, fine-grained dikes with barely discernable needles of plagioclase and pyroxene. Dikes 1 to 2 m wide. <b>Trachyte (Eocene)</b> : Dikes of light gray to dark-gray porphyry with 15 to 25%
Tit	plagioclase laths 10 to 20 mm long set in an equigranular to fine-grained matrix o dark green pyroxene and feldspar. The plagioclase laths are often distinctly aligned parallel to the margins of the dike. Dikes 1 to 4 m wide.
Tita	<b>Porphyritic trachyandesite (Eocene):</b> dikes and sills of light-gray porphyry with 5 to 15% phenocrysts of plagioclase feldspar, dark green pyroxene, and potassium feldspar that are 1-10 mm across. These intrusive bodies may grade into more equigranula textures (Tig). Syenitic and sedimentary xenoliths are locally present.
Titas	<b>Trachyandesite sill (Eocene):</b> Aerially extensive sill north of Highway 70 on the southwestern side of Cat Mountain. Porphyritic with phenocrysts of plagioclase pyroxenes, and biotite. Sill is 10 m thick.
Titd	<b>Porphyritic trachydacite (Eocene)</b> : dikes and sills of light-gray porphyry with 5 to 15% phenocrysts of plagioclase feldspar, biotite $\pm$ hornblende, and dark green pyroxene that are 1-10 mm across.
Tig	<b>Alkali gabbro-syenogabbro (Eocene)</b> : equigranular, fine- to medium-grained, salt and-pepper textured dikes and sills with plagioclase feldspar and pyroxene phenocrysts. These intrusive bodies may grade into more porphyritic textures (Tita) These intrusions often contain xenoliths of syenite (Figure 2). Dikes of this composition cut sills of megacrystic trachyte porphyry south of Coyote Peak and on the Tularosa NI quadrangle to the west. Alkali gabbro may contain amphibole and biotite (Moore e al., 1988). Thick sills of this composition in the northeastern parts of the map area have cumulate pyroxene at the base.
Tbag	<b>Coarse-grained alkali gabbro (Eocene):</b> Black to dark green medium to coarse grained sill with subequal amounts of black pyroxene and plagioclase with sparse apple green pyroxene located on the southwest side of Cat Mountain. The coarse gabbro grades upward into a fine- to medium grained equigranular gabbro.
Tim	<b>Megacrystic trachyte porphyry (Eocene):</b> greenish gray porphyritic sills and diker with megacrysts of embayed tschermakitic hornblende or green pyroxene that are up to 2 to 4 cm across. Often contains hornblende $\pm$ biotite. These intrusives can contain xenoliths of pink coarse-grained syenite with phenocrysts of orthoclase feldspar and hornblende.
Tip	<b>Amphibole-rich trachydacite (Eocene):</b> Black to dark green, medium-grained sills composed of ~30% dark green elongate amphibole needles, and subequal amounts o plagioclase and potassium feldspar. Sills 1 to 5 m thick.
Tis	<b>Syenite (middle Eocene)</b> : pink, medium-grained equigranular sill composed o potassium feldspar, plagioclase feldspar, pyroxene, and biotite.
Tid	<b>Alkali gabbro of Granite Well (Eocene)</b> : Greenish gray, equigranular, fine- to medium-grained, salt-and-pepper textured sill with plagioclase feldspar, pyroxene and biotite phenocrysts exposed south of Granite Well along Temporal Creek near the mountain front. This sill has a pryoxene-rich cumulate phase near the base. Sill is about 30 to 35 m thick.
Tsp	<b>Syenite of Temporal Creek (middle Eocene):</b> Tan to pink, equigranular to porphyritic stock exposed in the northeast corner of the quadrangle (Moore et al., 1988) north of Temporal Creek straddling the western boundary of the Mescalero Apache Indian Reservation. The most common phenocrysts are clinopyroxene, plagioclase feldspare and potassium feldspar. Some phases of this stock have amphibole (Moore et al. 1988), while others have sparse phenocrysts of quartz. Trachyte dikes with large plagioclase phenocrysts cut this stock.
Tg	<b>Gravel (Pliocene?):</b> Unconsolidated gravel resting on San Andres Limestone in a saddle on Coyote Ridge at an elevation of ~5980 feet (1822m). The gravel include sandstone, and pebbles of volcanic rocks from Sierra Blanca, as well as rounded limestone pebbles. < 1 m thick. <b>Mesozoic</b>
Km	<b>Mancos Shale (Upper Cretaceous):</b> Black, finely-laminated, fossiliferous shale with thin beds of limestone and intercalated bentonite beds near the base. Brown to tan cross-bedded to tabular-bedded sandstone and sandstone containing abundan oyster fossils occur higher in the exposed section. These marine deposits are in the lower part of the Rio Salado Tongue of the Mancos Shale (Moore et al., 1988).
Kd	<b>Dakota Sandstone (Upper Cretaceous):</b> Medium to thick-bedded to think laminated, medium- to coarse-grained, cross bedded white sandstone with lenses of dark gray shale. Chert and quartzite pebble conglomerate commonly occurs at o near the base of the unit. Distinctive, cylindrical <i>Ophiomorpha nodosa</i> burrows are abundant at the top of the Dakota Sandstone.
Τκm	<b>Moenkopi Formation (Upper Triassic):</b> Red-brown, fine-grained, thin bedded trough cross-bedded to ripple laminated sandstone. The sand grains are subround to round and include muscovite and lithic fragments. Conglomeratic sandstones with pebbles of chert and quartz are present in a few localities. The upper and lowe contacts are disconformable. The fluvial Moenkopi Formation scoured channels into the underlying Grayburg Formation of the Artesia Group. The unit thickens toward the

east. Thickness up to 10 m.

Pag Grayburg Formation, Artesia Group (Permian): Red siltstone with green reduction spots, red mudstone, and massive gypsum. Siltstone is massive to occasionally crosslaminated. Base with underlying Fourmile Draw Member of the San Andres Limestone is sharp and has a lag gravel. Gypsum occurs in the upper 10 m of the unit. The top is eroded and scoured by the Moenkopi Formation.

San Andres Formation (Permian): Kelley (1971) recognized three members in south central New Mexico. The lower part of the deposit, the Rio Bonito Member, is composed of medium- to thick-bedded dark gray, fossiliferous limestone, golden tan dolomite, and thin beds of yellow siltstone and gypsum. The middle part of the unit, the Bonney Canyon Member, is fossiliferous, gray thin-thick bedded dark gray limestone. The intercalated limestone gypsum unit in the upper part of the San Andres Formation belongs to the Fourmile Draw Member of Kelley (1971). In this area, only the Rio Bonito Member and the Fourmile Draw Member are present.

**Rio Bonito Member:** Dark gray, thin to thick bedded, fossiliferous limestone with thin to medium-bedded gypsum near the top of the unit. Fossils include productid brachiopods, crinoids, echinoderm spines, and bryozoa fragments. The contact with the underlying Yeso Group is generally quite sharp; however thin (<0.5 m) gypsum beds can persist a few meters above the contact. The Rio BonitoMember is 60 to 75 m thick.

Hondo Sandstone: This unit was described by Harbour (1971). The sandstone is present within the Rio Bonito Member approximately 30 m above the base of the limestone. The Hondo Sandstone is a gold-brown to tan, well-sorted, medium to fine-grained quartz arenite with well-rounded sand grains. Bedding is commonly tabular, cross-bedding is rare. The sandstone is discontinuous and ranges from 0 to 10 m thick. The sandstone, where present, is only 0.7 m thick on Cat Mountain (13 S 415916 3670526 NAD27) and is 5 to 10 m thick on the west end of Coyote Ridge.

Fourmile Draw Member: The Fourmile Draw Member is composed of interbedded light gray dolomite, dark gray limestone, and laminated gypsum. The carbonates contain fossils, but fossils are not as common as they are in the Rio Bonito Member. The relative percentage of gypsum to carbonate increases upsection through the member. The thickness of this unit is uncertain because the top is not exposed on the quadrangle and this unit is quite deformed. Estimated thickness is on the order of 300 m.

The contact between the Rio Bonito and Fourmile Draw members is very gradational and is here defined to be at the base of a distinctive limestone bed present across the map area. The marker limestone is dark gray, fossiliferous, and has a gypsum bed that is about 6 m thick below it. The carbonates above the marker limestone become progressively more dolomitic up section. The marker limestone is generally the thickest succession of limestone above cliffs of the Rio Bonito Member, but the marker does vary in thickness (20 to 80 feet; average is 40). The marker limestone forms an easily tracable unit on poorly exposed, gypsum-mantled hills.

The 6 m thick gypsum bed is continuous across the area, but thin, discontinuous gypsum and intercalated limestone beds often occur in the uppermost Rio Bonito Member as much as 25 m below the mappable marker limestone. Similarly, thin sequences of dark gray, fossiliferous limestone that grade up into dolomite and thick beds of gypsum are present above the marker limestone.

**Yeso Formation (Permian):** Red and yellow siltstone and thin beds of gray limestone and laminated to thin bedded gypsum dominate the basal part of the section. A 30 m thick medium-bedded limestone forms a prominent bed in the middle portion of the unit. The upper part of the Yeso on Cat Mountain is gypsum with yellow and red siltstone and thin-bedded limestone. Red siltstone often occurs near the top of the unit. In contrast, the upper Yeso Formation to the north is primarily gypsum and limestone. The limestones are sparsely fossiliferous, containing brachiopods. A solitary nautiloid 15 cm across was found in limestone near the base of the unit. Sea star fossils are at the Yeso/San Andres contact in the northeastern part of the map area. Oncolites are preserved in the limestone at 13 S 416913 3670824 (NAD 27). The Yeso Formation has a gradational contact with the underlying Abo Formation that is particularly well exposed south and southwest of Cat Mountain. The red arkosic sandstone and red and green mudstone of the Abo Formation grade upward into thin beds of black and green shale, fine-grained green and yellow, sandstone, and thin beds of gypsum and limestone. The upper contact is gradational with the San Andres Formation . Kelley (1971) measured a complete section of Yeso Formation on the Cat Mountain quadrangle in the vicinity of Coyote Peak in T13S, R10E, sec. 21 and 22. Maximum thickness 372 m (1220 ft).

**Abo Formation (Permian):** Brick red sandstone, mudstone, siltstone, and conglomerate. The Abo Fomation is dominated by mudstone (50%) and arkosic sandstone and conglomerate (40%). Fossiliferous limestone and pedogenic carbonate, which are primarily found in the basal 150 m, make up the rest of the unit (Otte, 1959). Sandstones are cross-bedded to tabular, coarse to medium-grained, and poorly to moderately sorted. Angular to subround quartz, orthoclase feldspar, and muscovite are common sand grains. Conglomerates contain well-rounded clasts of Proterozoic quartzite, Proterozoic rhyolite porphyry, chert of uncertain age, and sandstone. Thin fossiliferous limestone beds that are 0.2 to 0.7 m thick locally occur near the transitional base of the unit, particularly in the southeast corner of the quadrangle. Thin (4 to 20 cm) chert beds are locally found just above the contact with the Labrocita Formation and above a thin (0.2 m) limestone bed at 13 S 408507 3666281 (NAD 27). Petrified wood is preserved in the Abo Formation at 13 S 408658 3664679 (NAD27). Malachite and hematite mineralization is associated with the Abo Formation. 427 m to 550 m thick (Otte, 1959; Speer, 1983).

Bursum Formation (Late Pennsylvanian to Early Permian): Maroon and green andstone interbedded with fossiliferous limestone and black shale. Biohermal mounds (IPbb) are dominantly composed of algae but have abundant crinoid stems preserved at the top. Bryozoans and brachipods are other common fossils on the mountds. This unit is equivalent to the Labrocita Formation of Otte (1959). We followed Otte (1959) by using a continuous limestone bed (Otte's bed 55) to mark the gradational contact between the Bursum and the Abo formations. Base not exposed. Maximum thickness in the northern Sacramento Mountains is 330 m (Raatz, 2002) and estimated thickness northeast of Tularosa is 305 m (Otte, 1959).

## REFERENCES

Allen, M.S., and Foord, E.E., 1991, Geological, geochemical and isotopic characteristics of the Lincoln County porphyry belt, New Mexico: implications for regional tectonics and mineral deposits: New Mexico Geological Society, Guidebook 42, p. 97-102. Harbour, R. L, 1970, The Hondo Sandstone Member of the San Andres Limestone of southcentral New Mexico: U.S. Geological Survey Professional Paper 700-C, p. 175-182. Kelley, V.C., 1971, Geology of the Pecos country, southeastern New Mexico: New Mexico Bureau of Mines and Mineral Resources Memoir 24, 78 p. Koning, D.J., 1999, Fault segmentation and paleoseismicity of the southern Alamogordo fault, southern Rio Grande rift [M.S. thesis]: Albuquerque, University of New Mexico, 286 p. Koning, D.J., and Pazzalgia, F.J., 2002, Paleoseismicity of the Alamogordo fault zone along the Sacramento Mountains, southern Rio Grande rift, New Mexico: New Mexico Geological Society, 53<sup>rd</sup> Field Conference Guidebook, Geology of White Sands, p. 107-119. Koning, D.J., Pazzaglia, F.J., and Smartt, R., 2002, Alluvial fan stratigraphy along the southern

Sacramento Mountains, N.M., and inferences regarding late Quaternary plaeoclimate, soils, and sedimentation: New Mexico Geological Society, 53rd Field Conference Guidebook, Geology of White Sands, p. 289-302. McManus, C.E.D., and McMillan, N.J., 2002, Subduction or rifting: The Eocene magmas of the Sacramento Mountains, NM: Geological Society of America Abstracts with Programs, v. 34, no. 6, p. 364.

McMillan, N.J., 2004, Magmatic record of Laramide subduction and the transition to Tertiary extension: Upper Cretaceous through Eocene igneous rocks of New Mexico: New Mexico Geological Society Special Publication 11, p. 249-270. Miller A.K., and Youngquist, W., 1949, American Permian Nautiloids: Geological Society of America Memoir 41, 218 p.

Moore, S.L., Foord, E.E., Meyer, G.A., and Smith, G.W., 1988, Geologic map of the northwestern part of the Mescalero Apache Indian Reservation, Otero County, New Mexico: U.S. Geological Survey Map I-1895, scale 1:24,000. Otte, C., 1959, Late Pennsylvanian and early Permian stratigraphy of the northern Sacramento Mountains, Otero County, New Mexico, New Mexico Bureau of Mining and Mineral Resources, Bulletin 50, 111 p.

Pray, L.C., 1961, Geology of the Sacramento Mountains escarpment, Otero County, New Mexico, New Mexico Bureau of Mining and Mineral Resources, Bulletin 35, 144 p. Raatz, W.D., 2002, A stratigraphic history of the Tularosa Basin area, south-central New Mexico: New Mexico Geological Society Guidebook 53, p. 141-157. Speer, S.W., 1983, Abo Formation, north-central Sacramento Mountains: An onlapping fluvial clastic wedge: *in* Guidebook for field trip to the Abo red beds (Permian), central and south-central New Mexico: Roswell Geological Society and the New Mexico Bureau of Mines and Mineral Resources, p. 54-72.