

	MESOZ	OIC STRATA
above, with relatively abundant 1-4 m-thick.	Kccu	Crevasse Canyon Formation, upper unit (Upper Cretaceous, Coniacian North American Stage) A prominent ledge-forming
lolocene) Qfy, as described		sandstone in the upper part of the Crevasse Canyon Formation. Overlies sandstone and shale beds of Kccm . Overlying strata not exposed. Sand is white to light gray, medium- to coarse-grained, subrounded to subangular, well-sorted, and composed of quartz with 1-
topographic high. Surface gravel no to very sparse surface gypsum		5% mafics and cherty lithics. Bedding is laminated to thin, and tangential cross-stratified to planar-horizontal bedded. Sand is strongly cemented. Interpreted to correlate to the basal unit of the Cub Mountain Formation of Arkell (1983), but we follow the strati-
ubordinate gravel. Unconformably		graphic schemes of Weber (1964), Lucas et al. (1989) and Cather (1991) in keeping this sandstone in the upper Crevasse Canyon For- mation. Not observed south of Three Rivers drainage. Up to 110 m-thick in the northeast corner of the quadrangle, but thins to the south.
vy to cross-stratified. No soil de-	Kccm	Crevasse Canyon Formation, middle unit (Upper Cretaceous, Coniacian North American Stage) Fluvial strata consisting of
desert pavement development or is generally less than ~150 years in		interbedded channel-fills of sandstone and shaly floodplain deposits; very minor coal beds. Sandstone is mostly fine- to medium- grained, subrounded to subangular, well-sorted, and composed of quartz, feldspar, and 5-10% lithic grains; high amounts (around 20%)
		of glauconite(?); these were classified as sublithic to subarkosic by Arkell (1982). Colors of the sand range from pale yellow, golden, light olive gray, pale olive, very pale brown, and light brownish gray. Beds are generally tangential to trough cross-stratified
ene) Qfy2, as described above, i0 years in age. 1-2 m-thick.		(laminated or very thin to thin beds), with subordinate to subequal planar-horizontal bedding (laminated to thin). Sandy channel-fill complexes may be up to 6 m-thick, are well-cemented, and form ledges. Floodplain deposits consist of shale and very fine- to fine-
yey sand generally is light brown well consolidated. Base not de-		grained sandstone and silty sandstone, with local, very minor coal beds. Shale is pale yellow to light gray to gray to light olive gray in color and fissile. Carbonaceous shale is dark gray to black in color. Unit correlates with the Barren Member and Coal-Bearing Member of the Mesa Verde Group of Arkell (1983), but coal in this unit is not common in the map area. Base of unit not well-exposed on
generally covered by slopewash on the surface, commonly along		this quadrangle, but noted to be gradational in Arkell (1983). Thickness not well constrained, due to variability in dips, but on the scale of 450 to 600 m.
graphic highs. Unit is recognized	Kccl	Crevasse Canyon Formation, lower unit (Upper Cretaceous, lower Coniacian North American Stage) Interbedded sandstone
		and shale interpreted to reflect marginal marine (shoreface, foreshore, lagoonal pond, and deltaic) and fluvial depositional environ- ments (Arkell, 1983). Unit gradationally overlies Mancos Shale, and its lowest strata are clearly shoreface facies. Base assigned to
- Color is generally pale brown to		lower shoreface sandstone bed that is interbedded in gray, marine shale, below which shale exceeds about 80% of strata. Lower shore- face sandstones are thick, tabular, internally massive to planar-laminated, and interbedded with gray marine shales. Sand is very fine-
fine- to fine-grained sand interbed- and sandy pebbles. Most of the		to fine-grained, calcareous, subrounded to subangular, well-sorted, and composed of quartz with 5-10% lithic grains; glauconite(?) may be present. Colors range from light gray to olive-yellow to light yellowish brown to pale yellow. Upper shoreface sandstones are
erate ped development and stage I sent. Both fining-upward and locally subdivided into Qay2 and		slightly coarser (mostly fine-grained) and in very thin to thick, tabular beds that are internally planar-laminated or cross-stratified (tangential or trough-cross-laminated). Sand is pale yellow to light gray, well-sorted, subrounded to subangular, and a lithic to feld-spathic arenite; glauconite(?) is present in the matrix (20-25% surface area). The upper shoreface sandstone grades upward into a \sim 150
(see description of that unit below). In 1 m), and incorporated into the		m(?) thick interval of interbedded, interpreted upper shoreface sandstones and fluvial sandstones, with subordinate oyster beds and light shale intervals 1-3 m-thick. One coal bed observed. The sand is commonly golden, medium- to thickly bedded, and internally
to upper Holocene in age. Unit I-4 m-thick.		planar-laminated to low angle cross-laminated to very thinly bedded. About 20% glauconite(?) is present in the sand. The sand is fine- to medium-grained, subrounded to subangular, well-sorted, and a lithic to feldspathic arenite. Although the upper shoreface sand
Unconformably overlies Qay1 . soil development. Sand is light		may correlate to the Gallup Sandstone, it is not continuously exposed across the quadrangle and brackish water facies (containing oyster beds), are locally beneath it; therefore, this strata is subsumed into the Crevasse Canyon Formation. Upper contact placed at high-
ngular, moderately to poorly ny up to about 50 cm-tall. No de-		est, golden brown, highly calcerous, very fine- to fine-grained sandstone bed containing marine or brackish water fossils; this bed is generally 30-100 cm-thick. This golden brown, fossiliferous sandstone(s) is relatively continuous throughout the study area and commonly form ledges. 50-150 m-thick.
drangle to the south. This unit is	Km	Mancos Shale (Upper Cretaceous, upper Cenomanian(?) through Turionian North American Stage) Fissile shale that is pla-
ing an erosion surface cut below		nar- to wavy-laminated; colors range from gray to light gray to light olive gray to light greenish gray. Metamorphosed to a black to gray argillite adjacent to laccoliths. Strata hosts laccoliths in the southeast part of the quadrangle; emplacement of these laccoliths has
o light yellowish brown (10YR6/3-		folded and deformed the adjoining Mancos Shale. Unit includes the pale yellow to white, fine-grained, shoreface sands of the Tres Hermanos Formation. Unit gradational overlies the Dakota Sandstone. Cross-section A-A' and well data indicate a thickness of 185-215 m.
- to fine-grained sand interbedded sandy pebbles. In general, the	Kd	Undifferentiated Dakota Sandstone (upper Cretaceous) Combined unit of upper and lower Dakota Sandstone (units Kdu and
erate ped development and stage I sent. Both fining-upward and		Kdl), which are described in detail below. Not differentiated due to insufficient exposure or extensive faulting. 45-55 m-thick.
es most of the undifferentiated Qay eposits. Differentiated from Qs by	Kdu	upper Dakota Sandstone (upper Cretaceous) Intercalated light gray siltstone, very fine-grained sandstone, and light to dark gray shale with subordinate beds of fine- to medium-grained sandstone. Shale is laminated and friable. Sandstone beds are medium to
-thick.		thick, tabular, and internally cross-stratified (very thin beds to laminations that are trough-cross-stratified to tangential-cross-stratified; up to 20 cm-thick); also planar-laminated. Trough cross-stratification commonly indicates a northeast paleoflow direction. Sand is white, well-sorted, subrounded to subangular, and composed of quartz. Sandstone beds are well-cemented. Uppermost sandstone beds
grained sand (minor medium- to		are extensively burrowed, with <i>ophiomorpha</i> burrows common. Slightly purplish brown desert varnish forms on weathered sandstone. Top contact placed at top of upper quartz arenite sandstone with <i>ophiomorpha</i> burrows, which is overlain by dark gray shale of the
and pebbles may be present. subrounded, and poorly sorted;		Mancos Formation. 25-30 m-thick
orted, and contains abundant lithic bed on top of the unit. The calcium thick. Unit is inset into Qao , and	Kdl	lower Dakota Sandstone (upper Cretaceous) Ledge-forming, quartz arenite sandstone. Beds are laminated to very thin and tan- gential- to trough-cross-stratified; also some planar-horizontal-beds (laminated to medium). Sand size is fine-upper to medium-upper. Sand is white subrounded to subangular and mostly well-sorted. Locally beds of aforementioned sand are mixed with minor lithic
thick. Unit is inset into Qao , and tern boundary because it is gener-		Sand is white, subrounded to subangular, and mostly well-sorted. Locally, beds of aforementioned sand are mixed with minor lithic grains (including chert), coarse- to very coarse-grained sand (composed largely of lithic grains), and sparse, very fine to very coarse, subrounded to rounded pebbles of quartz, chert, and quartzite. Lower contact is a scoured unconformity, and deposit locally fills pale-
expose the lower gravel; forms a		ovalleys at its base. This valley-fill is up to 6 m-thick and contains reddish to purplish shale beds and lithic arenite sandstone beds. Slightly purplish brown desert varnish forms on weathered sandstone. Approximately 20 m-thick.
-	Trm	Moenkopi Formation (middle Triassic, Anisian North American Stage) Cross-stratified, fine- to coarse-grained sandstone
Qao3, Qao2, and Qao1 (see de- that are cemented by gypsum. Base defined as top of Santa Fe		(mostly fine- to medium-grained). Colors of the sand range from reddish brown to light reddish brown to pale-weak red to pinkish gray. Varying proportions of weak red to reddish brown shale, claystone, and siltstone; these fine-grained lithologies are subordinate to the south, but may increase in abundance northwards. Sandstone is extensively cross-stratified (commonly trough- to tangential-
and defined as top of Salita Fe		to the south, but may increase in abundance northwards. Sandstone is extensively cross-stratified (commonly trough- to tangential- cross-stratified and laminated to very thinly-bedded; 10-15 cm-thick foresets); minor planar-horizontal bedding and thin to medium, lenticular beds. Sandstone is subrounded to subangular, moderately to well sorted, and a lithic arenite (5-25% lithic grains, including
n unit Qai that are inset into older se to clayey very fine- to fine-		mica, and ~5% estimated potassium feldspar). Locally, very minor (1-5%) beds of pebbly medium- to very coarse-grained sandstone (subrounded and composed of quartz with abundant chert and quartzite grains); pebbles are very fine to coarse and composed of
, highly gypsiferous, and includes Clasts are rounded to subangular		rounded-subrounded chert and quartzite. Lower contact is a planar to scoured unconformity over the Artesia Group (scour relief of 1- 3 m). Strata immediately below the overlying Dakota Sandstone may locally be bleached to a pale green and yellowish gray color.
, is internally massive, and has ower than a pediment surface on tabular deposit up to a few meters		Assigned to Anton Chico Member by Lucas (1991); correlation to this member was verified in a field visit by Lucas during the spring of 2009. No fossils observed. Mostly 30-50 m-thick, but possibly as much as 70 m-thick at a location 0.5 m-thick of lower Three Rivers drainage.
a stage III calcium carbonate hori- tot covered by slopewash. Paleo-	PALEO	ZOIC STRATA
s unit is noteably gravelly and 10- beds.	Pag	Grayburg Formation, Artesia Group (upper Permian, middle Guadalupian North American Stage) Very fine- to fine-grained
luvium subunit, this deposit has a . . Sediment consists of gypsifer-		sandstone and silty to clayey very fine- to fine-grained sandstone; subordinate siltstone and shale. Colors range from orange to red to light red to reddish brown (most to least common). Strata are generally in very thin to thick, tabular (minor irregular) beds. Reduction (bleached) spots 0.5-2 mm in diameter cover 1-15% of rock area, with higher coverage along bedding and fault planes (where they are
d sand and pebbles (scattered or in clay and ped development), with		dm-scale and irregular). Sand is subangular to rounded (mostly subrounded), well-sorted, and composed of quartz, with less than 10% possible feldspar. No fossils observed. Thick gypsum or anhydrite beds are absent to the north but increase to the south (to about 5%-
). Base of deposit overlies Meso- nit Qao1 . A strong gypsum hori-		10% of the unit; more common towards the top). Gypsum and anhydrite beds are generally deformed. Although definitively Artesia Group, we assign this to the Grayburg Formation because of lithologic characteristics, thickness, and the fact that higher formations of
horizon or else a petrogypsic hori- tot covered by slopewash. Several		this group extend progressively shorter distances from the deepest parts of the Delaware Basin to the south (Kelley, 1971 and 1972). We did not recognize the Queen Formation of the Artesia Group because we did not note scattered large, rounded, frosted quartz grains
t to the Alamogordo fault. Well		indicative of the upper part of this formation (Tait et al., 1962). Orangish color, fine texture, and quartz arenite composition serve to differentiate this unit from the overlying Moenkopi Formation. Lower contact with San Andres Formation is a disconformity (see report for more discussion). Thickness is difficult to estimate based on map data because of irregular bedding attitudes where this unit is
ported, sandy gravel filling paleo- bebbles with minor cobbles and		relatively well-exposed to the south. Well data at cross-section A-A' indicate a thickness of 90-110 m.
rtiary intrusives with minor lime- barse-grained, subrounded (minor	Psa	San Andres Formation, undifferentiated (lower to upper Permian, Leonardian to Guadalupian North American Stage) Most of the San Andres Formation on the quadrangle belongs to the Four-mile Draw Member, which is described below. However, inacces-
ining Qoa2 , so it is inferred to be m-thick.		sible strata of the southern Phillips Hills were lumped together into this undifferentiated unit, which includes limestone and dolomite strata below the Four-mile Draw Member (i.e., Rio Bonito Member of Kelley, 1971). 240-250 m-thick, based on subsurface data for wells used in cross-section A-A'.
middle Pleistocene) Generally d (minor medium to very coarse	Psaf	Fourmile Draw Member of the San Andres Formation (upper Permian, Guadalupian North American Stage) Micritic, dark
e. Gravel contains pebbles with and composed of intermediate in-		gray limestone, grayish tan to light gray dolomite, and gypsum. Beds are medium to thick and tabular. Carbonates become more dolo- mitic up-section, and the proportion of gypsum beds increases up-section. Approximately 100-120 m-thick.
wn to very pale brown (10YR 6- enerally several meters above the trath heights, this unit is assigned	Psafd	Dissolution-collapsed Fourmile Draw Member of the San Andres Formation (upper Permian) Chaotic assemblage of limestone
alent petrogypsic horizon devel- o2 becomes increasingly difficult		and fine-grained sandstone that involves the upper Fourmile Draw Member and overlying Grayburg Formations. Interpreted to be due to dissolution of gypsum beds in the Fourmile Draw Member.
oped as Qao1 is actually Qao2 that		RFACE STRATA DEPICTED ONLY IN CROSS-SECTION low the Bursum Formation, depth picks of Paleozoic strata are from King and Harder (1985).
	Ру	Yeso Formation (lower Permian) Yellow and red siltstone, limestone, and gypsum. 630 m-thick.
ined sand, locally with minor scat- to light yellowish brown (10YR	Pa	Abo Formation (lower Permian) Reddish color; consists of overbank deposits of mudstone and clayey fine-grained sandstone that are intercalated with coarse channel-fills of sandstone and pebbly sandstone. Approximately 450 m-thick.
ular-bedded and internally mas- with ~10% estimated potassium	Pb	Bursum Formation (uppermost Pennsylvanian(?) to lowermost Permian) Marine strata (shales and limestone beds) interbedded
calcic horizons (stage II morphol- hin, yellowish red (5YR 5/6), illu- al, very thin to medium lenses of		with fluvial sediment of overbank reddish shale together with channel-fill sandstone and pebbly sandstone. Approximately 500 m-thick.
entioned, relatively fine-grained ravel are subrounded to rounded,	IPh	Holder Formation (upper Pennsylvanian) Limestone, gray and red calcareous shale, sandstone, and conglomerate (from Pray, 1961, fig. 3). 120 m-thick.
subordinate Mesozoic sandstone; ly strongly impregnated by calcium	IPb + IP	g Beeman and Gobbler Formations, undifferentiated (lower to upper Pennsylvanian) Interbedded limestone, shale, and sand-
se of gravelly Qao1 deposits. he footwall of the Alamogordo ea. Greatest exposed thickness is 2		stone (from Pray, 1961, fig. 3). 400-430 m-thick.
bult hanging wall, unit is up to \sim 700	MDS	Undifferentiated strata of Mississippian, Devonian and Silurian age Unit may include Lake Valley, Percha, and Fusselman formations, among others discussed in Pray (1961) and listed in King and Harder, 1985, fig. 21). 64-85 m-thick.
	0	Undifferentiated strata of Ordivician age Unit includes Montoya and El Paso Formations (King and Harder, 1985, fig. 21). 120-150 m-thick.
ocks occur as dikes (marked on	pC	Precambrian rocks Rhyolite or granite, reddish and composed largely of quartz and feldspar (from mud log of the Lewelling #2
ills are common in the Crevasse		well).
ypabyssal rocks occupying sills and ze of groundmass 0.2-0.5 mm, less		Thousands of years (ka) 0 Qam, 20 & Qfy2i Qfy2i
ablende and pyroxene. Phenocrysts and smaller plagioclase (5-20% of		2 — Qamf, & ae a Qamg af Qs& Qse Qc Qec, Qamg Qsgy Qsec Qec, Qec, Qec, Qay2i Qay2
nd sills consist of this rock type,		4 – Qe, & Qe
gray to gray, equigranular to l. Grain size of ground mass is		
ornblende and pyroxene. Exposed		
n color. Minerals include plagio-		40 — Qao3
to degree of porphyritic character		
only small laccoliths. Rock is		$\begin{array}{c c} 80 & \longrightarrow & & & & & & & & & & & & & & & & & $
phenocrysts up to 0.2-10 mm in		Millions of
5% surface coverage by		years (Ma) scale change 0.1 [2] 2, 2, 2, 2, 2, 2, 2, 2, 2, 2, 2, 2, 2,
at extends 5 km in a ENE direction ldspar (0.5-10 mm), pyroxene (0.2-		
ong) of aligned feldspar in a		
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Δ <		5.0 — Tsf
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3000		45 — break 85 —
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-0		100 — C Rivers 7.5-minute
-1000		break and scale change quadrangle, N.M.
-2000		