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Comments to Map Users

A geologic map displays information on the distribution, nature, orientation, and age relationships of rock and deposits and the occurrence of structural features. Geologic and fault contacts are irregular surfaces that form boundaries between different types or ages of units. Data depicted on this geologic quadrangle map may be based on any of the following: reconnaissance field geologic mapping, compilation of published and unpublished work, and photogeologic interpretation. Locations of contacts are not surveyed, but are plotted by interpretation of the position of a given contact onto a topographic base map; therefore, the accuracy of contact locations depends on the scale of mapping and the interpretation of the geologist(s). Any enlargement of this map could cause misunderstanding in the detail of mapping and may result in erroneous interpretations. Site-specific conditions should be verified by detailed surface mapping or subsurface exploration. Topographic and cultural changes may not be shown due to recent

Cross sections are constructed based upon the interpretations of the author made from geologic mapping and available geophysical and subsurface (drillhole) data. Cross sections should be used as an aid to understanding the general geologic framework of the map area, and not be the sole source of information for use in locating or designing wells, buildings, roads, or other man-made structures.

development.

expedite dissemination of these geologic maps and map data to the public as rapidly as possible while allowing for map revision as geologists continued to work in map areas. Each map sheet carries the original date of publication below the map as well as the latest revision date in the upper right corner. In most cases, the original date of publication coincides with the date of the map product delivered to the National Cooperative Geologic Mapping Program (NCGMP) as part of New Mexico's STATEMAP agreement. While maps are produced, maintained, and updated in an ArcGIS geodatabase, at the time of the STATEMAP deliverable, each map goes through cartographic production and internal review prior to uploading to the Internet. Even if additional updates are carried out on the ArcGIS map data files, citations to these maps should reflect this original publication date and the original authors listed. The views and

Digital layout and cartography by the NMBGMR Map Production Group: Phil L. Miller, Amy L. Dunn, Ann D. Knight, and Justine L. Nicolette

Correlation of Map Units (contained in the San Luis Basin)

	01.01.01 Contact—Identity and existence are certain. Location is accurate.	II	31.10 Cross section line and label
	01.01.03 Contact—Identity and existence are certain. Location is approximate.	UD	02.11.02 Fault showing local normal offset (2nd option) $-U$, upthrown block; D, downthrown block.
	01.01.07 Contact—Identity and existence are certain. Location is concealed.	DU	02.11.03 Fault showing local reverse offset—Showing dip value and direction. U, upthrown block; D, downthrown block.
	01.01.09 Internal contact—Identity and existence are certain. Location is accurate.	Δ	02.14.04 Fault-breccia zone or zone of broken rock around fault
	01.01.11 Internal contact—Identity and existence are certain. Location is approximate.		02.15.01 Small, minor inclined fault—Showing strike and dip.
	02.01.01 Fault (generic; vertical, subvertical, or high-angle; or unknown or unspecified orientation or sense of slip)—Identity and existence are certain. Location is accurate.	-	04.03.02 Small, minor inclined joint (1st option)—Showing strike and dip.
	02.01.03 Fault (generic; vertical, subvertical, or high-angle; or unknown or unspecified orientation or sense of slip)—Identity and existence are certain. Location is approximate.		06.02 Inclined bedding—Showing strike and dip.
	02.01.04 Fault (generic; vertical, subvertical, or high-angle; or unknown or unspecified		08.03.02 Inclined metamorphic or tectonic foliation—Showing strike and dip.
?	orientation or sense of slip)—Identity or existence are questionable. Location is approximate.	ţ	09.017 Inclined slickenline, groove, or striation on fault surface (1st option)—Showing bearing and plunge.
	02.01.07 Fault (generic; vertical, subvertical, or high-angle; or unknown or unspecified orientation or sense of slip)—Identity and existence are certain. Location is concealed.	1	12.05 Fluvial transport direction
?	02.01.08 Fault (generic; vertical, subvertical, or high-angle; or unknown or unspecified orientation or sense of slip)—Identity or existence are questionable. Location is concealed.	ξ	30.03.12 Spring, as shwon on topographic maps or on general-purpose or smaller scale maps
•	02.02.01 Normal fault—Identity and existence are certain. Location is accurate. Ball and bar on downthrown block.		02.11.17 Thrust fault or reverse fault (in cross section)—Arrows show relative motion.
. — —	02.02.03 Normal fault—Identity and existence are certain. Location is approximate. Ball and bar on downthrown block.	G	03.01.03 Boundary located by gravity survey
•	02.02.07 Normal fault—Identity and existence are certain. Location is concealed. Ball and bar on downthrown block.		31.02.25 Well location (in cross section)—The location and depth of a well used to establish stratigraphy and geologic unit depth.
·?	02.02.08 Normal fault—Identity or existence are questionable. Location is concealed. Ball and bar on downthrown block.		01.02.01 Key bed (Tsbo)—Identity and existence are certain. Location is accurate.
7 0 4 0 4 0, 5 0 4 6 4 0,	02.14.04 Fault-breccia zone or zone of broken rock around fault		02.11.17 Thrust fault or reverse fault (in cross section)—Arrows show relative motion.

<u>_____</u> 03.02.10 Faults located by geophysical methods.

→ 19.03.07 Trench (generalized trace)

The New Mexico Bureau of Geology and Mineral Resources created the Open-file Geologic Map Series to conclusions contained in these map documents are those of the authors and should not be interpreted as



02-00-00-00-heading01-Sedimentary deposits of the Rio Grande Rift & Adjacent Highlands-Sedimentary deposits of the Rio Grande Rift & Adjacent Highlands (Cenozoic)-Sedimentary deposits of the Rio Grande Rift & Adjacent Highlands 02-01-00-00-map unit-tailings-Tailings ponds (modern-historic)-Tailings ponds (modern-historic) – Areas of artificially deposited fill and debris; delineated where aerially extensive; consists predominantly of mining-related mill tailings and tailings dams west of Questa; the geologic map shows the pre-tailings geology based on an

-00-00-00-map unit-water-water (Cenozoic)-Water - surface water of the Rio

nterpretation of aerial photos.

Mountains east of Questa.

several meters.

02-02-00-00-00-map unit-af-Artificial fill and disturbed land (Cenozoic (modernhistoric))—Artificial fill and disturbed land (modern-historic)—Excavations and areas human-deposited fill and debris; shown only where aerially extensive. 02-03-00-00–map unit–ds–Mine waste rock and related features (Cenozoic odern-historic))—Mine waste rock and related features (modern-historic)—Angular and finer deposits, mainly from Tertiary plutonic rocks; principally located in adjacent to the open pit molybdenum mine located in the Sangre de Cristo

02-04-00-00-map unit-Qal-Alluvium (latest Pleistocene and Holocene)—Alluvium (latest Pleistocene and Holocene)—Generally brownish and/or reddish, poorly to moderately sorted, angular to rounded, thinly to thickly bedded, loose silt and silty sand with subordinate coarse lenses and thin to medium beds of mostly locally derived clasts; mapped in active channels, floodplains, low (young) alluvial terraces, tributary-mouth fans, and some valley-slope colluvial deposits; weak to no soil development; clasts along the Rio Hondo are principally granitic rock types, quartzite, and basalt; clasts along tributaries draining the western side of the Rio Grande are principally volcanic rock types; clasts along the Rio Lucero are principally granitic rock types with some quartzite; clasts along the Rio Pueblo de Taos are principally granitic rock types, quartzite, and sandstones; drainages south and east of the Rio Pueblo de Taos are dominated by sandstone and other sedimentary rock types; clasts in the Taos and Los Cordovas quadrangles are principally granitic, metamorphic, volcanic, and sandstone rock types; up to 7 m estimated thickness.

02-05-00-00-map unit-Qm-Marsh deposits (Holocene)-Marsh deposits (Holocene)—Silt, sand, and clay in low relief, saturated flatlands; high organic content; hosts a variety of streams, springs, and bogs; located primarily between US-64 and Taos Pueblo, bordering the Rio Lucero.

02-06-00-00-map unit-Qc-Colluvium (middle Pleistocene to Holocene)—Colluvium (middle Pleistocene to Holocene)—Mostly locally derived, light- to dark-brown, orange, and rarely reddish, poorly to moderately sorted, angular to well-rounded, silty to sandy conglomerate/breccia with clasts locally to >1m; mapped on hill slopes and valley margins only where it obscures underlying relations; mantles slopes in Red River gorge and northeastern side of Red River fault zone in the eastern part of the Guadalupe Mountain quadrangle; widespread along the bases of mountain-front facets; dominated by quartzite and granitic rock types north of the Rio Pueblo de Taos; dominated by sandstone and pebble conglomerate with minor limestone clasts south of the Rio Pueblo de Taos; in the northwestern part of the Arroyo Hondo quadrangle, west of the Rio Grande, the deposits consist of thin mantles overlying volcanic bedrock; estimated at generally less than 5 m thick.

02-07-00-00-map unit -Qs -Talus and scree (Pleistocene to Holocene) -Talus and scree (Pleistocene to Holocene) – Angular rock fragments as much as 1 m in diameter forming talus cones, talus aprons, and scree slopes; locally well sorted; grades into colluvium as sand and silt content increases; shown only in the Sangre de Cristo Mountains by Lipman and Read (1989).

02-08-00-00-map unit-Qad-Alluvium in closed depressions (latest Pleistocene to Holocene)-Alluvium in closed depressions (latest Pleistocene to Holocene)-Light to dark-brown, very thin- to medium-bedded, loose, massive, shady to silty beds with thin, discontinuous layers of pebbles and rare cobbles (to ~15 cm) found on toreva (rotational) blocks associated with landslide complexes; local thickness of at least

02-09-00-00-map unit-Qsw-Sheetwash alluvium (late Pleistocene to olocene?)—Sheetwash alluvium (late Pleistocene to Holocene?)—Alluvial apron composed mostly of pebbly to silty sand that accumulated on gentle slopes, such as ose on Servilleta Basalt (Tsb); some of the silt- to fine sand-size fraction in these deposits may be of eolian origin (Shroba and Thompson, 1998); deposits of unit Qsw exist along the shores of intermittent ponds or small lakes on the Servilleta Basalt (Tsb); low-lying areas of unit Qsw are susceptible to sheet flooding due to unconfined overland flow, and locally to stream flooding and gullying; recently disturbed surface of unit Qsw may be susceptible to minor wind erosion; estimated thickness is 1 to 5 m, put possibly as much as 10 m (Thompson et al., 2014).

02-10-00-00-map unit-Qty-Young stream terrace deposits (latest Pleistocene to Holocene) – Young alluvial-fan (Qfy) and stream terrace deposits (Qty) (latest Pleistocene to Holocene)—Poorly sorted deposits of silt, sand, pebbles, cobbles and boulders; deposits are typically clast-supported and poorly bedded; pebble and cobble clasts are typically imbricated; terrace deposits unconformably overlie the local bedrock; clasts are primarily sedimentary rocks, quartzite, slate, schist, metavolcanic, granitic rocks, and Tertiary granitic and volcanic rocks; uppermost sediments are typically silty sand probably deposited from overbank flow; weak to moderate pedogenic development, including A, Bw, Bwk and Bk soil horizons and stage I to II calcium carbonate development; map unit Qty is typically on valley floors of large to medium drainages, whereas Qfy exists as young mountain-front fans and valley fills in small tributaries; includes units Qf6 along the mountain front (Arroyo Seco quadrangle), Qt6 (Los Cordovas quadrangle) and Qt8 (Arroyo Hondo and Arroyo Seco

quadrangles) of Kelson (1986); thickness up to 5 m. 02-11-00-00-map unit-Qfy-Young alluvial-fan deposits (latest Pleistocene to Holocene)—Young alluvial-fan (Qfy) and stream terrace deposits (Qty) (latest Pleistocene to Holocene)—Poorly sorted deposits of silt, sand, pebbles, cobbles and boulders; deposits are typically clast-supported and poorly bedded; pebble and cobble clasts are typically imbricated; terrace deposits unconformably overlie the local pedrock; clasts are primarily sedimentary rocks, quartzite, slate, schist, metavolcanic, granitic rocks, and Tertiary granitic and volcanic rocks; uppermost sediments are typically silty sand probably deposited from overbank flow; weak to moderate pedogenic development, including A, Bw, Bwk and Bk soil horizons and stage I to II calcium carbonate development; map unit Qty is typically on valley floors of large to medium drainages, whereas Qfy exists as young mountain-front fans and valley fills in

quadrangle), Qt6 (Los Cordovas quadrangle) and Qt8 (Arroyo Hondo and Arroyo Seco quadrangles) of Kelson (1986); thickness up to 5 m. 02-12-00-00–map unit–Qfyv–Young alluvial-fan deposits from volcanic terrane (latest Pleistocene to Holocene) - Young alluvial-fan deposits from volcanic terrane (latest Pleistocene to Holocene)-Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of volcanic rock types; associated soils have stage I calcium

small tributaries; includes units Qf6 along the mountain front (Arroyo Seco

the Rio Grande and drainages on Guadalupe Mountain.

direction was from the southwest; up to several meters thick.

Grande and Red River gorges.

carbonate development; source areas are primarily the volcanic terrane on west side of

02-13-00-00-map unit-Qe-Eolian deposits (late Pleistocene to Holocene)-Eolian deposits (late Pleistocene to Holocene)-Light-colored, well-sorted, fine to medium sand and silt deposits that are recognized as laterally extensive, low-relief, sparsely vegetated, mostly inactive, sand dunes and sand sheets that overlie Servilleta Basalt on the Taos Plateau; rare gravel lag; weak to moderate soil development; northeasttrending longitudinal dune-crest orientations indicate that the predominant wind

02-14-00-00-map unit-Qls-Landslides in the Rio Grande gorge and tributaries (late Pleistocene to Holocene)-Landslides in the Rio Grande gorge and tributaries (late Pleistocene to Holocene)—Poorly sorted rock debris and sand to boulder debris transported downslope; occurs on slopes marked by hummocky topography and downslope-facing scarps; includes small earth flow, block-slump, and block-slide deposits; includes large rotational Toreva slide blocks within the Rio Grande and Rio Pueblo de Taos gorges, which include large, rotated and detached blocks of intact Servilleta Basalt; may also include areas underlain by Holocene colluvium in the Rio

02-15-00-00–map unit–Qlsm–Landslides in the Sangre de Cristo Mountains Pleistocene to Holocene) – Landslides in the Sangre de Cristo Mountains (Pleistocene to Holocene)—Lobate accumulations of poorly sorted soil and rock debris on slopes marked by hummocky topography and downslope-facing scarps; derived from bedrock and glacial deposits, and includes small earth flow, block-slump, and block-

slide deposits (from Lipman and Read, 1989). 02-16-00-00-map unit-Qfo-Alluvial fan deposits, undivided (middle to late Pleistocene) – Alluvial fan deposits, undivided (middle to late Pleistocene) – Poorly sorted silt, sand, pebbles, and cobbles; in the Guadalupe Mountain quadrangle, Qfo is composed primarily of intermediate and basaltic volcanic clasts; moderate pedogenic development, including A, Bt, Btk and Bk soil horizons and stage III and IV calcium carbonate development; upper soil horizons are commonly affected by surface erosion; probably overlaps with units Qf2 through Qf4, and with alluvial units Qt2 through Qt6, but not assigned to other fan units because of lack of well-defined age control, clear

stratigraphic position, and distinct lithologic characteristics; thickness up to 3 m. 02-17-00-00–map unit–Qfu–Alluvial fan deposits, undivided (middle to late Pleistocene) – Alluvial fan deposits, undivided (middle to late Pleistocene) – Poorly sorted silt, sand, pebbles, and cobbles; mapped along majority of Sangre de Cristo range front, but not correlated to other fan units because of lack of well-defined age control, clear stratigraphic position, or distinct lithologic characteristics; probably correlative with alluvial fan deposits Qf1 through Qf6.

02-18-00-00-map unit-Qmt-Moraine and till (Pleistocene)-Moraine and till (Pleistocene) – Terminal and lateral moraines, and thick valley-bottom till; poorly sorted and generally unstratified clay, silt, and sand containing erratic boulders; characterized by hummocky or ridged topography; some till is mapped with colluvium by Lipman and Read (1989), although some Qmt deposits were remapped in this study. 02-19-00-00–map unit—Qtu—Stream terrace deposits, undivided (middle to late Pleistocene)—Stream terrace deposits, undivided (middle to late Pleistocene)—Poorly sorted silt, sand, pebbles, and boulders; clasts primarily of quartzite, schist, granite, and sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, volcanic rock types; associated soils have stage II to III calcium carbonate development; granite, and volcanic rock types; associated soils have stage I to II calcium carbonate typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock or unit QTl; probably correlative with Qt1 through Qt4. 02-20-00-00-map unit-Qf1-Alluvial-fan deposits (middle Pleistocene)-Alluvial-

fan deposits (middle Pleistocene)—Poorly sorted silt, sand, pebbles, and boulders; stage III to IV calcium carbonate development, although soil horizons are commonly affected by surface erosion; in the Los Cordovas quadrangle, the clasts are principally granitic, intermediate volcanic, basalt, and metamorphic rock types; granitic clasts are also present east of Arroyo del Alamo; in the Taos quadrangle, clasts are primarily granitic and metamorphic rock types; the deposit is finer grained to the north and away from the Picuris Mountains range front; Qf1 is differentiated from QTI by larger clast size (Kelson, 1986), less oxidation, poor sorting, absence of abundant manganese oxide staining, and clasts that are less weathered; slope of Qf1 surface on the Taos and Los Cordovas quadrangles is southwesterly, and is dissected by numerous southwesterly trending arroyos; on the Taos quadrangle, Qf1 is correlative with Unit Q1p of Kelson (1986); a tephra within Qf1 deposits on the Taos SW quadrangle was dated at 1.27 ± 0.02 Ma (40Ar-39Ar method, W. McIntosh, personal commun., 1996); the deposit is more than 12 m thick in the northeastern part of Los Cordovas quadrangle, and is thinner from northeast to southwest; directly southwest of Taos Municipal Airport, Qf1 is less than about 1 m thick and unconformably overlies Servilleta Basalt (Tsb); elsewhere, Qf1 appears to overlie unit QTl or Tsb; Qf1 is more than 12 m thick in the northwestern part of the Taos and Arroyo Hondo quadrangles, and is thinner from northeast to southwest; on Blueberry Hill, Qf is about 3 m thick at US-64, and about 2 m thick to the southwest, where it unconformably overlies the unit QTI; it is more than 5 m thick in the northeastern part of the Arroyo Seco quadrangle, and thins from northeast to southwest; it is more than 5 m thick in the northeastern part of the Questa quadrangle, and thins from northeast to southwest.

02-21-00-00-heading02-Terrace Deposits of the Rio Grande-Terrace Deposits of the Rio Grande (Cenozoic)—Terrace Deposits of the Rio Grande

2-21-01-00-00-map unit-Oao3-Older alluvium (middle? Pleistocene)-Older lluvium (middle? Pleistocene) – Poorly sorted silt, sand, and pebbles; clasts primarily of granitic, metamorphic, basaltic, and intermediate volcanic rocks; distinctly smaller clast sizes than units Qt2rr, Qt1rr, and Qt0rr; upper soil horizons are locally affected by surface erosion; may be mantled locally by unit Qe; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock near the rim of the Rio Grande gorge; located only upstream of the Red River fault zone; correlative with unit Qao3 of Ruleman et al. (2007) in the Sunshine quadrangle.

02-21-02-00-00-map unit-Qt2rg-Stream terrace deposits of the Rio Grande (middle) Pleistocene)—Stream terrace deposits of the Rio Grande (middle? Pleistocene)—Poorly sorted silt, sand, pebbles, and boulders; clasts primarily of granitic, metamorphic, intermediate volcanic, basalt, and sedimentary rocks; locally contains clasts of Amalia Tuff; associated soils have stage III to IV calcium carbonate development, thick argillic Bt soil horizons, and 7.5YR to 10YR hues in soil Bt horizons; upper soil horizons may be affected by surface erosion; may be mantled locally by unit Qe; possibly faulted by the Dunn fault in the Arroyo Hondo quadrangle; modified from Kelson (1986); estimated thickness 1 to 10 m.

02-21-03-00-00 – map unit – Qt1rg – Stream terrace deposits of the Rio Grande (early to middle? Pleistocene)—Stream terrace deposits of the Rio Grande (early to middle? Pleistocene)—Poorly sorted sand, pebbles, and cobbles; clasts of basalt, quartzite, slate, schist, and other metamorphic rock types, volcanic rock types, and (rarely) sandstone and limestone; locally contains clasts of 25 Ma Amalia Tuff; where preserved, associated relict soils have stage III to IV calcium carbonate development, thick argillic Bt soil horizons, and 7.5YR hues in soil Bt horizons; upper soil horizons commonly

affected by surface erosion; may be mantled locally by unit Qe; estimated thickness 1 to

02-21-04-00-00—map unit—Qt0rg—Stream gravel deposited by the ancestral Rio Grande (early? to middle? Pleistocene)—Stream gravel deposited by the ancestral Rio Grande (early? to middle? Pleistocene)—Poorly sorted sand, pebbles, and cobbles; clasts of basalt, quartzite, slate, schist, other metamorphic rock types, and volcanic rock types; very rare Amalia Tuff clasts; associated with the broad, highest terrace west of he Rio Grande; upper soil horizons commonly affected by surface erosion; locall

02-22-00-00–heading02–Terrace Deposits of the Rio Pueblo de Taos–Terrace Deposits of the Rio Pueblo de Taos (Cenozoic)-Terrace Deposits of the Rio Pueblo de

mantled by eolian san

02-22-01-00-00-map unit-Qt7rp-Stream terrace deposits of the Rio Pueblo de Taos early to middle Holocene)-Stream terrace deposits of the Rio Pueblo de Taos (early to middle Holocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts

primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage I calcium carbonate development; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock or unit QTI. 2-22-02-00-00-map unit-Qt6rp-Stream terrace deposits of the Rio Pueblo de Tao (latest Pleistocene)-Stream terrace deposits of the Rio Pueblo de Taos (latest

Pleistocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage I to II alcium carbonate development; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock or unit QTI; associated with the Q6 surface of Kelson (1986).

02-22-03-00-00-map unit-Qf4-Alluvial-fan deposits of the Rio Pueblo de Taos and tributaries (middle? to late Pleistocene) - Alluvial-fan and stream terrace deposits of the Rio Pueblo de Taos and tributaries (middle? to late Pleistocene)-Poorly sorted silt, sand, pebbles, and boulders; associated soils have stage III calcium carbonate development, argillic Bt soil horizons and 10YR to 7.5YR hues in Bt horizons; clast primarily of granitic and metamorphic rocks north of the Rio Pueblo de Taos, and granitic, metamorphic, and sedimentary rock types south of the Rio Pueblo de Taos; clasts also include basaltic rock types along Arroyo Seco and along the Rio Pueblo de

Taos downstream of Los Cordovas; modified from Kelson (1986). 2-22-04-00-00-map unit-Qt4rp-Stream terrace deposits of the Rio Pueblo de Taos nd tributaries (middle? to late Pleistocene)—Alluvial-fan and stream terrace deposits f the Rio Pueblo de Taos and tributaries (middle? to late Pleistocene)—Poorly sorted silt, sand, pebbles, and boulders; associated soils have stage III calcium carbonate development, argillic Bt soil horizons and 10YR to 7.5YR hues in Bt horizons; clasts primarily of granitic and metamorphic rocks north of the Rio Pueblo de Taos, and ranitic, metamorphic, and sedimentary rock types south of the Rio Pueblo de Taos lasts also include basaltic rock types along Arroyo Seco and along the Rio Pueblo de

aos downstream of Los Cordovas; modified from Kelson (1986).

02-22-05-00-00-map unit-Qt3rp-Alluvial-fan and stream terrace deposits of the Rio Pueblo de Taos and tributaries (middle? to late Pleistocene)—Alluvial-fan and stream terrace deposits of the Rio Pueblo de Taos and tributaries (middle? to late Pleistocene)—Poorly sorted silt, sand, pebbles, and boulders; associated soils have stag

II to III calcium carbonate development; clasts primarily of quartzite, slate, and schist; ranitic clasts also exist east of Arroyo del Alamo; possible Qt3rp remnant inset into errace Qt2rp on the western side of Taos may be an artificial terrace related to esidential development; modified from Kelson (1986). 02-22-06-00-00-map unit-Qf2-Alluvial fan deposits of the Rio Pueblo de Taos and

tributaries (middle? Pleistocene)-Alluvial fan and stream terrace deposits of the Rio Pueblo de Taos and tributaries (middle? Pleistocene)-Poorly sorted silt, sand, pebbles, and boulders; clasts primarily of granitic and metamorphic rocks north of Rio Pueblo de Taos, and granitic, metamorphic and sedimentary rocks south of Rio Pueblo de Taos; associated soils have stage III to IV calcium carbonate development, thick argillic Bt soil horizons, and 7.5YR to 10YR hues in soil Bt horizons; upper soil horizons may be ffected by surface erosion; modified from Kelson (1986).

and tributaries (middle? Pleistocene) – Alluvial fan and stream terrace deposits of the Rio Pueblo de Taos and tributaries (middle? Pleistocene)-Poorly sorted silt, sand, pebbles, and boulders; clasts primarily of granitic and metamorphic rocks north of Rid Pueblo de Taos, and granitic, metamorphic and sedimentary rocks south of Rio Pueblo de Taos; associated soils have stage III to IV calcium carbonate development, thick argillic Bt soil horizons, and 7.5YR to 10YR hues in soil Bt horizons; upper soil horizons may be affected by surface erosion; modified from Kelson (1986).

02-22-08-00-00-map unit-Qt1rp-Stream terrace deposits of the Rio Pueblo de Tao (middle Pleistocene)-Stream terrace deposits of the Rio Pueblo de Taos (middle Pleistocene)—Poorly sorted silt, sand, pebbles, and boulders; clasts primarily of quartzite, slate, and schist; granitic clasts also exist east of Arroyo del Alamo; finer grained to the north, away from the Picuris Mountains range front; stage III to IV alcium carbonate development; a tephra within Qf1 deposits in the Taos SW quadrangle was dated at 1.27 ± 0.02 Ma (40Ar-39Ar method, W. McIntosh, personal commun., 1996); present along the west rim of the Rio Grande del Rancho valley. 2-23-00-00–heading02–Terrace Deposits of the Rio Hondo–Terrace Deposits of

the Rio Hondo (Cenozoic)-Terrace Deposits of the Rio Hondo 02-23-01-00-00-map unit-Qt8rh-Stream terrace deposits of the Rio Hondo (middle to late Holocene)-Stream terrace deposits of the Rio Hondo (middle to late Holocene)-Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of

quartzite, schist, granite, and volcanic rock types; deposits have negligible soil development; typically present as thin (< 5 m) alluvial deposit beneath high-stage floodplain or adjacent to active alluvial channels. 02-23-02-00-00-map unit-Qt7rh-Stream terrace deposits of the Rio Hondo (early to middle Holocene)-Stream terrace deposits of the Rio Hondo (early to middle

02-23-03-00-00-map unit-Qt6rh-Stream terrace deposits of the Rio Hondo (latest Pleistocene)—Stream terrace deposits of the Rio Hondo (latest Pleistocene)—Poorly development; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock or unit QTl; associated with the Q6 surface of Kelson (1986). 02-23-04-00-00-map unit-Qt5rh-Stream terrace deposits of the Rio Hondo (late Pleistocene)-Stream terrace deposits of the Rio Hondo (late Pleistocene)-Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage II to III calcium carbonate

volcanic bedrock or unit QTl; associated with the Q5 surface of Kelson (1986). 02-23-05-00-map unit-Qt4rh-Stream terrace deposits of the Rio Hondo (middle? to late Pleistocene) – Stream terrace deposits of the Rio Hondo (middle? to late Pleistocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage III calcium carbonate development, argillic Bt soil horizons and 10YR to 7.5YR hues in Bt horizons; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on

development; typically present as thin (<5 m) alluvial deposit on strath surfaces cut on

volcanic bedrock or unit QTI; associated with the Q4 surface of Kelson (1986). 02-23-06-00-00-map unit-Qt3rh-Stream terrace deposits of the Rio Hondo (middle? to late Pleistocene)—Stream terrace deposits of the Rio Hondo (middle? to late Pleistocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage III calcium carbonate development, argillic Bt soil horizons and 10YR to 7.5YR hues in Bt horizons.

02-23-07-00-00-map unit-Qt2rh-Stream terrace deposits of the Rio Hondo (middle? Pleistocene)—Stream terrace deposits of the Rio Hondo (middle? Pleistocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage III to IV calcium carbonate development, thick argillic Bt soil horizons, and 7.5YR to 10YR hues in soil Bt horizons; upper soil horizons are locally affected by surface erosion. 02-24-00-00–heading02–Terrace Depostis of the Red River–Terrace Depostis of the

Red River (Cenozoic)—Terrace Depostis of the Red River 02-24-01-00-00-map unit-Qt8rr-Stream terrace deposits of the Red River (middle to late Holocene)—Stream terrace deposits of the Red River (middle to late Holocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; deposits have negligible soil development; typically present as thin (< 5 m) alluvial deposit beneath high-stage floodplain or adjacent to active alluvial channels; equivalent to Qt8 of Kelson (1986) and

Pazzaglia (1989).

02-24-02-00-map unit-Qt7rr-Stream terrace deposits of the Red River (early to middle Holocene)—Stream terrace deposits of the Red River (early to middle Holocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage I calcium carbonate development; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock; equivalent to Qt7 of Kelson (1986) and Pazzaglia

02-24-03-00-00-map unit-Qt6rr-Stream terrace deposits of the Red River (latest Pleistocene)—Stream terrace deposits of the Red River (latest Pleistocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage I to II calcium carbonate development; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock or unit QTI; associated with the Q6 surface of Kelson (1986).

02-24-04-00-00-map unit-Qt5rr-Stream terrace deposits of the Red River (late Pleistocene)—Stream terrace deposits of the Red River (late Pleistocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage II to III calcium carbonate development; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock or unit QTl; associated with the Q5 surface of Kelson (1986).

02-24-05-00-00-map unit-Qt4rr-Stream terrace deposits of the Red River (middle? to late Pleistocene)—Stream terrace deposits of the Red River (middle? to late Pleistocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage III calcium carbonate development, argillic Bt soil horizons and 10YR to 7.5YR hues in Bt horizons; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock or unit QTl; associated with the Q4 surface of Kelson (1986).

02-24-06-00-00-map unit-Qt3rr-Stream terrace deposits of the Red River (middle? to late Pleistocene)—Stream terrace deposits of the Red River (middle? to late Pleistocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage III calcium carbonate development; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock or unit QTI; equivalent to Qt3 of Kelson (1986) and Pazzaglia (1989).

02-24-07-00-00 – map unit – Qt2rr – Stream terrace deposits of the Red River (middle? Pleistocene)—Stream terrace deposits of the Red River (middle? Pleistocene)—Poorly sorted silt, sand, pebbles, cobbles, and boulders; clasts primarily of quartzite, schist, granite, and volcanic rock types; associated soils have stage III to IV calcium carbonate development; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock or unit QTI; includes correlative terrace deposit flanking the southwestern side of Lama Canyon on the Questa and Guadalupe Mountain quadrangles; equivalent to Qt2 of Kelson (1986) and Pazzaglia (1989).

02-24-08-00-00-map unit-Qt1rr-Stream terrace deposits of the Red River (middle Pleistocene)-Stream terrace deposits of the Red River (middle Pleistocene)-Poorly sorted silt, sand, pebbles, and boulders; clasts of basalt, quartzite, metamorphic rock types, and volcanic rock types; soil development not documented but upper soil horizons are probably affected by surface erosion; present only locally along the rim of 03-03-00-00-map unit – Tbo – Older Basalt (Miocene) – Older Basalt (Miocene) – Ir the Red River gorge, where Qt1rr is inset into Qt0rr gravel deposits and Tertiary volcanic rocks.

02-24-09-00-00-map unit-Qt0rr-Old stream terrace deposits flanking the Red River and tributaries (early? to middle? Pleistocene)—Old stream terrace deposits flanking the Red River and tributaries (early? to middle? Pleistocene)—Poorly sorted sand, pebbles, and cobbles; clasts of basalt, quartzite, and many volcanic and metamorphic rock types; the upper part is commonly affected by surface erosion; the unit is present upstream and downstream of the Red River Fish Hatchery and in the confluence area between the Rio Grande and the Red River; Qt0rr merges with unit Qt0rg in the southernmost part of the Guadalupe Mountain and Arroyo Hondo quadrangles. 02-25-00-00-00-heading02-Santa Fe Group-Santa Fe Group (Cenozoic)-Santa Fe

Group 02-25-01-00-00—map unit—QTl—Lama formation (Pliocene to early? Pleistocene)—Lama formation (Pliocene to early? Pleistocene)—Poorly sorted sand, pebbles, and cobbles; clasts of basalt, quartzite, other metamorphic rock types, and other volcanic rock types; locally high percentage of angular to subangular quartzite pebbles and cobbles; commonly cross-bedded, and stained with black manganese oxide and yellowish-orange iron oxide coatings; oxidized; clasts are typically weathered or grussified; contains distinct discontinuous sandy interbeds; commonly crudely

imbricated; imbrication suggests westerly flow direction in area north of Taos Municipal Airport, and southerly flow direction in areas north and west of Rio Pueblo de Taos, with northwesterly flow direction in area southeast of Rio Pueblo de Taos; wel drillers records in the Questa area show clay layers in the shallow subsurface that are interpreted as lacustrine deposits (Bauer et al., 2015); the unit is present between the Sangre de Cristo Mountains range front and the Rio Grande gorge over most of the map area; correlative with Lambert's (1966) two informal facies of the "Servilleta Formation" (the "sandy gravel facies" found south of the Rio Hondo, and the "gravelly silt facies" found between the Rio Hondo and the Red River); correlative with Kelson's (1986) informal "Basin Fill deposit;" correlative with the unit previously informally called "Blueberry Hill formation" in the Taos area; also correlative with Pazzaglia's

(1989) late Neogene-Quaternary rift fill sequence (unit Q1) which he informally named the Lama formation; herein, for this study area, the Lama formation is defined as the uppermost, pre-incision, sedimentary rift fill, and where extant represents the uppermost member of the Santa Fe Group; the unit therefore includes all of the basin fill between the oldest Servilleta Basalt (40Ar/39Ar age of 5.55 ± 0.37 Ma near Cerro Azul, D. Koning, personal commun., 2015) and the oldest Rio Grande (and tributary) terrace gravels (e.g., Qt0rg, Qt0rr); the Lama formation and the underlying Chamita Formation are texturally and compositionally similar and may be indistinguishable in

boreholes, although Koning et al. (2015) noted a coarsening of sediment (south of this map area) that roughly coincides with the Chamita/Lama contact in this map area; the top of the Lama formation is typically marked by a sharp unconformity and color/textural contrasts with overlying gravels; the unit contains several laterally variable components of sedimentary fill that are associated with various provenance reas related to east- or west-flowing tributary watersheds that have been fair persistent in the late Cenozoic; locally contains tephra layers; reworked tephra in a road

cut near the Red River Fish Hatchery (elevation ca. 7160 ft) was probably derived from nearby ca. 5 Ma volcanic units (R. Thompson, personal comm., 2015); a tephra in the uppermost Lama formation yielded a date of ~1.6 Ma based on a chemical correlation with the 1.61 Ma Guaje Pumice eruption in the Jemez Mountains (elevation ca. 7660 ft, M. Machette, personal comm., 2008); thickness ranges from zero to an exposed thickness of about 25 m at the southwestern end of Blueberry Hill, but may be considerably thicker in other parts of the map area.

02-25-01-01-00—subunit—Tcc—Clay layers in the Lama formation of the Santa Fe Group (Pliocene)-Clay layers in the Lama formation of the Santa Fe Group (Pliocene)—In cross section only; boreholes east of Guadalupe Mountain have quartzite, schist, granite, and volcanic rock types; associated soils have stage I calcium loped behind lava dams along the ancestral Red River drainage; lateral continuity and extent are unknown; clays appear to influence the characteristics of groundwater flow systems by perching or mounding water above the regional aquifer, resulting in a locally elevated water table; layers are locally up to at least 100 feet thick.

Description of Map Units (contained in San Luis

02-25-02-00-00-map unit-QTsf-Santa Fe Group, undivided (Miocene to early Pleistocene)—Santa Fe Group, undivided (Miocene to early Pleistocene)—In cross

02-25-03-00-00—map unit—Tc—Chamita Formation, undivided, Santa Fe Group (Miocene? and Pliocene)-Chamita Formation, undivided, Santa Fe Group (Miocene? and Pliocene)—In cross sections only. Sedimentary deposits between the lowest Servilleta Basalt and the Tesuque Formation; typically rounded to subrounded pebbleto cobble-size clasts in a sand to silt matrix; thick sections to the south reflect Proterozoic clast provenance and are dominated by schist, quartzite, and amphibolite with lesser volcanic clasts derived from the Latir volcanic field; locally, thin interbeds are typically dominated by pebble-size clasts in a fine sand to silt matrix and commonly top of Tc is herein defined as the sediments below the oldest Servilleta Basalt flows.

02-25-04-00-00—map unit—Tt—Tesuque Formation, Santa Fe Group (Miocene)—Tesuque Formation, Santa Fe Group (Miocene)—In cross section only. Basin-fill deposits of clay, silt, sand, pebbles, cobbles, and boulders of the Rio Grande

02-26-01-00-00-map unit-Tp-Picuris formation, undivided (Oligocene to Miocene)-Picuris formation, undivided (Oligocene to Miocene)-In cross section only. In the Picuris Mountains area (Bauer, et al., 2017) this unit consists of an upper member of tuffaceous and pumiceous silty sandstones and volcaniclastic sandstone and conglomerate; a member of buff to white and/or pinkish, silty sandstone to fine cobble conglomerate and nonfriable to strong, very fine lower to very coarse upper, very poorly to moderately sorted, rounded to subangular, thinly to thickly bedded, silicacemented silty to pebbly sandstone which locally contains a basal portion of poorly sorted pebbly/gravelly sandstone and/or cobble/boulder conglomerate composed exclusively of Proterozoic clasts; a member of light buff, yellowish, and locally white, ash-rich, quartzose, silty, fine sand to pebbly, pumiceous sandstone; a lower member of

subrounded, pebbly/silty sandstone and mudstone containing very thick(?) to thin beds and/or lenses and/or isolated clasts of subangular to rounded Proterozoic quartzite (up to 3 m across) and massive quartzite conglomerate; paleoflow measurements near the Picuris Mountains indicate source to the north (Rehder, 1986; Aby et al., 2004); age range is from at least 35.6 Ma to less than 25 Ma; thickness unknown, but at least 450 m in the Picuris Mountains area. 03-00-00-00-heading01-Rocks of the Taos Plateau Volcanic Field-Rocks of the

Taos Plateau Volcanic Field (Cenozoic)-Rocks of the Taos Plateau Volcanic Field 03-01-00-00-map unit—Tsb—Servilleta Basalt (Pliocene)—Servilleta Basalt (Pliocene)-Flows of dark-gray tholeiitic basalt characterized by small olivine and tabular plagioclase phenocrysts, diktytaxitic texture, and local vesicle pipes and segregation veins; forms thin, fluid, widespread pahoehoe basalt flows of the Taos Plateau volcanic field erupted principally from large shield volcanoes in the central part of the Taos Plateau (Lipman and Mehnert, 1979) but also from several small shields and vents to the northwest of the map area near the Colorado border (Thompson and Machette, 1989; K. Turner, personal comm., 2014); additional buried vents west of the

with a maximum thickness of approximately 50 m in the Rio Grande gorge 16 km northwest of Taos; Tsb can locally be subdivided into the lower Servilleta Basalt (Tsbl the middle Servilleta Basalt (Tsbm), and the upper Servilleta Basalt (Tsbu) that are separated by sedimentary intervals as much as 70 m thick in the southern part of the map area (Leininger, 1982); 40Ar/39Ar ages from basalts exposed in the Rio Grande gorge (Cosca et al., 2014) range in age from 4.78 ± 0.03 Ma for the lowest basalt near the Gorge Bridge, to 3.59 ± 0.08 Ma for the highest basalt flow at the Gorge Bridge, broadly

consistent with previous results by Appelt (1998); the base of the upper Servilleta Basalt lava flow section at La Junta Point yielded an 40Ar/39Ar age of 3.78 ± 0.08 Ma (sample 10RG05 - M. Cosca, personal comm., 2014), whereas a lava flow at the base of the section south of Cerro Chiflo yielded an 40Ar/39Ar age of 3.78 ± 0.08 Ma (sample RT08GM02 - M. Cosca, personal comm., 2014). 03-01-01-00-00-subunit-Tsbu-Servilleta Basalt, upper (Pliocene)-Servilleta Basalt,

the middle Servilleta Basalt (Tsbm), and the upper Servilleta Basalt (Tsbu) that are separated by sedimentary intervals as much as 70 m thick in the southern part of the map area (Leininger, 1982); 40Ar/39Ar ages from basalts exposed in the Rio Grande gorge (Cosca et al., 2014) range in age from 4.78 ± 0.03 Ma for the lowest basalt near the Gorge Bridge, to 3.59 ± 0.08 Ma for the highest basalt flow at the Gorge Bridge, broadly consistent with previous results by Appelt (1998); the base of the upper Servilleta Basalt lava flow section at La Junta Point yielded an 40Ar/39Ar age of 3.78 ± 0.08 Ma (sample 10RG05 - M. Cosca, personal comm., 2014), whereas a lava flow at the base of the

section south of Cerro Chiflo yielded an 40Ar/39Ar age of 3.78 ± 0.08 Ma (sample RT08GM02 - M. Cosca, personal comm., 2014). 03-01-02-00-00-subunit-Tsbl-Servilleta Basalt, lower (Pliocene)-Servilleta Basalt, lower (Pliocene)—Tsb can locally be subdivided into the lower Servilleta Basalt (Tsbl) the middle Servilleta Basalt (Tsbm), and the upper Servilleta Basalt (Tsbu) that are separated by sedimentary intervals as much as 70 m thick in the southern part of the

map area (Leininger, 1982); 40Ar/39Ar ages from basalts exposed in the Rio Grande gorge (Cosca et al., 2014) range in age from 4.78 ± 0.03 Ma for the lowest basalt near the Gorge Bridge, to 3.59 ± 0.08 Ma for the highest basalt flow at the Gorge Bridge, broadly consistent with previous results by Appelt (1998); the base of the upper Servilleta Basalt lava flow section at La Junta Point yielded an 40Ar/39Ar age of 3.78 ± 0.08 Ma (sample 10RG05 - M. Cosca, personal comm., 2014), whereas a lava flow at the base of the section south of Cerro Chiflo yielded an 40Ar/39Ar age of 3.78 ± 0.08 Ma (sample RT08GM02 - M. Cosca, personal comm., 2014).

03-02-00-00–map unit—Tsbo—Older Servilleta Basalt (Pliocene)—Older Servilleta Basalt (Pliocene)—In cross section only. Identified in borehole BOR-6 and aeromagnetic cross section only. Identified in borehole BOR-5 and aeromagnetic data.

03-04-00-00-00-map unit-Tdmc-Dacite of unnamed cerrito east of Montoso (UCI near-vent deposits (Pliocene)—Dacite of unnamed cerrito east of Montoso (UCEM) near-vent deposits (Pliocene)—Near-vent deposits associated with lava flows of map unit Tdm; predominantly cinder, spatter agglutinate and local volcanic bombs.

03-05-00-00-map unit-Tdm-Dacite of unnamed cerrito east of Montoso (UCEM) (Pliocene)—Dacite of unnamed cerrito east of Montoso (UCEM) (Pliocene)—Dark gray sparsely phyric, low-silica, calc-alkaline dacite (64 wt% Si02, 6 wt% Na20+K20) lava flows erupted from two vent areas east of Cerro Montoso; contains rare skeletal pyroxene phenocrysts and resorbed, subhedral olivine and quartz xenocrysts in microcrystalline to glassy groundmass; locally includes small volume, aerially restricted andesite flows (McMillan and Dungan, 1986); 40Ar/39Ar age determination of 4.08 ± 0.04 Ma (sample 11RG42) and 4.6 ± 0.02 Ma (sample 11RG27) from north and south UCEM areas respectively (M. Cosca, personal comm., 2014); Appelt (1998) reported a similar 40Ar/39Ar age determination of 4.11 ± 0.13 Ma from a northern UCEM exposure; UCEM locally caps the west rim of the Rio Grande gorge, forming a

and local interbedded sedimentary deposits (Leininger, 1982; Peterson, 1981); these relations are not shown at the scale of this map due to the extensive landslide deposit (unit Qls); scoria and spatter agglutinate are common near the poorly defined vent 03-06-00-00-map unit-Tvr-Volcanic deposits of the Red River volcano (Pliocene)-Volcanic deposits of the Red River volcano (Pliocene)-Dacite lava flows and near-vent pyroclastic deposits of moderate relief on the south side of Guadalupe Mountain and in canyon exposures in the middle and upper reaches of the Red River where dacite lava flows cap the gorge sequence on both sides of the drainage; the lava flows exposed on both sides of the Red River were fed locally by dikes exposed on both

sides of the canyon; McMillan and Dungan (1986) reported chemical compositions for the underlying basaltic andesite (unit Tvhc) to dacite suite ranging from 52 to 61 wt% SiO2 and from 4.2 to 7.4 wt% Na2O+K2O; medium-grey dacite lavas are porphyritic, containing 5-15% phenocrysts of augite and bronzite with common olivine xenocryst in a fine-grained to glassy groundmass of plagioclase, glass, pyroxenes, and tantanomagnetite (McMillan and Dungan, 1986); dacite lavas are typically thick, up to tens of meters locally, and are characteristically discontinuous and aerially restricted; deposits of the Red River volcano overlie andesitic lava flows of the Hatchery volcano (unit Tvh) and locally deposits of south Guadalupe Mountain (unit Tdgs); 40Ar/39A

2014) was obtained from a sample collected near the northeastern limit of exposed deposits. 03-07-00-00-map unit-Tv-Basalt (Pliocene?)-Basalt (Pliocene?)-Basa

03-08-00-00-map unit-Tvh-Volcanic deposits of Hatchery volcano 'Pliocene)—Volcanic deposits of Hatchery volcano (Pliocene)—Includes a sequence o lava flow, intercalated volcanic breccia, and near vent pyroclastic deposits in canyon exposures in the middle and upper reaches of the Red River drainage and as low relief hills adjacent to the Red River; lava flows include a series of predominantly basaltic andesite and andesite lava flows; McMillan and Dungan (1986) reported chemical compositions for the basaltic andesite to overlying dacite (unit Tvr) ranging from 52 to 61 wt% SiO2 and from 4.2 to 7.4 wt% Na2O+K2O; dark gray basaltic andesite and andesite lava flows typically contain 5-10% phenocrysts of olivine and plagioclase; olivine phenocrysts can be large (up to 6mm) exhibiting well-developed skeletal overgrowths (McMillan and Dungan, 1986); andesite lava flows with aa flow tops and well exposed basal flow breccias tend to be thin, a few meters to 10 m thick, and are

laterally continuous based on exposures in the Red River canyon; deposits of the Hatchery volcano overlie dacite lava flows of Guadalupe Mountain, and locally overly two lava flows of Servilleta Basalt at the base of the Red River gorge near the New Mexico State Fish Hatchery (not differentiated at the map scale); 40Ar/39Ar age determination of 4.82 ± 0.07 Ma (sample 11RG42 - M. Cosca, personal comm., 2014) was obtained from a sample at the base of the section approximately 0.6 km southwest of the New Mexico State Fish Hatchery.

Geologic Cross Section A–A' digitized from AirBOR6 cross section by Paul W. Bauer, V. J. S. Grauch, Keith I. Kelson, Ren A. Thompson, and Benjamin J. Dreuth, Dec 2016

Los Cordovas Quadrangle | Taos Quadrangle

Taos Regional Airport BOR-6 Airport well

carbonate development; typically present as thin (< 5 m) alluvial deposit on strath surfaces cut on volcanic bedrock or unit QTI.

02-22-07-00-00-map unit-Qt2rp-Stream terrace deposits of the Rio Pueblo de Taos

red, greenish, and yellowish, moderately to very poorly sorted, subangular to

Rio Grande likely exist; flows typically form columnar-jointed cliffs where exposed,

thin veneer that is typically a single flow thickness over the Servilleta Basalt (unit Tsbu

age determination of 4.67 ± 0.06 Ma (sample RT08GM12 - M. Cosca, personal comm.

03-09-00-00—map unit—Tvhc—Volcanic deposits of the Hatchery volcano, near vent (Pliocene)—Volcanic deposits of the Hatchery volcano, near vent (Pliocene)—Near-vent section only. Basin-fill clay, silt, sand, pebbles, cobbles, and boulders of the Rio Grande deposits associated with lava flows of map unit Tvh; predominantly cinder, spatter, and utinate exposed in the Red River drainage approximately 1.25 km northwest of the New Mexico State Fish Hatchery; near-vent spatter, agglutinate, and volcanic bombs are common near hill 7590' on the south side of the Red River.

03-10-00-00-map unit—Tao—Andesite of Cerro de la Olla (Pliocene)—Andesite of Cerro de la Olla (Pliocene)—Dark gray to black, porphyritic, olivine andesite (58.5 wt% SiO2, 6.9 wt% Na20+K20) lava flows that erupted from vents near the summit of Cerro de la Olla, one of the largest, petrologically uniform, shield volcanoes of the Taos Plateau volcanic field (Lipman and Mehnert, 1979); contains 2-3% phenocrysts of includes the rock types above in addition to subangular and subrounded volcanic clasts olivine in a microcrystalline groundmass of plagioclase, olivine, augite, Fe-Ti oxides, derived locally from adjacent volcanic highlands of the Taos Plateau volcanic field; the the lower slopes of Cerro de la Olla in the northwestern part of the map area are commonly mantled in colluvium and rarely preserve well-developed flow morphology; instead outcrops typically exhibit blocky flow tops and remnants of numerous discontinuous and aerially restricted flow lobes; Appelt (1998) reported an 40Ar/39Ar age of 4.97 ± 0.06 Ma for a groundmass separate from the west side of Cerro de la Olla.

> 03-11-00-00–map unit–Tdgn–Dacite of Guadalupe Mountain, north (Pliocene)-Dacite of Guadalupe Mountain, north (Pliocene)-Predominantly trachydacite lava flows (62 wt% SiO2, 6.3 wt% Na20+K20) and associated near-vent pyroclastic deposits; contains sparse, small phenocrysts of plagioclase, hypersthene, and augite in a pilotaxitic glassy groundmass; proximal lava flows, lava dome remnants, and near-vent pyroclastic deposits consisting mostly of spatter and agglutinate of the geographic north peaks of Guadalupe Mountain; spatter and cinder leposits are found locally in association with flank lavas and may represent remobilized central vent deposits or mark the location of satellite vents on the flanks of north Guadalupe Mountain; distinguished from lava flows of south Guadalupe Mountain on the basis of reversed magnetic polarity based on paleomagnetic and aeromagnetic determinations (M. Hudson and V.J.S. Grauch respectively, personal comm., 2014; Grauch et al., 2015; Bauer et al., 2015); 40Ar/39Ar age determination of 5.04 ± 0.04 Ma (sample 10RG06 - M. Cosca, personal comm., 2014).

03-12-00-00–map unit–Tdgs–Dacite of Guadalupe Mountain, south (Pliocene)—Dacite of Guadalupe Mountain, south (Pliocene)—Predominantly trachydacite lava flows (62 wt% SiO2, 6.3 wt% Na20+K20) and associated near-vent pyroclastic deposits; contains sparse, small phenocrysts of plagioclase, hypersthene, and augite in a pilotaxitic glassy groundmass; proximal lava flows, lava dome remnants, and near-vent pyroclastic deposits consisting mostly of spatter and agglutinate of the geographic south peaks of Guadalupe Mountain; distinguished from lava flows of north Guadalupe Mountain on the basis of reversed magnetic polarity based on paleomagnetic and aeromagnetic determinations (M. Hudson and V.J.S. Grauch respectively, personal comm., 2014; Grauch et al., 2015; Bauer et al., 2015); 40Ar/39Ar age determination of 5.00 ± 0.04 Ma (sample 10RG07 - M. Cosca, personal comm., 2014); stratigraphic position relative to unit Tdgn is based on geophysical modeling of aeromagnetic data (B. Drenth, V.J.S. Grauch, personal comm., 2014) and age constraints relative to geomagnetic time scale; Appelt (1998) reported 40Ar/39Ar ages of 5.11 ± 0.08 Ma and 5.34 ± 0.06 Ma for groundmass separates from the south side of Guadalupe Mountain.

03-13-00-00-00—map unit—Tdg—Dacite of Guadalupe Mountain, undivided (Pliocene)—Dacite of Guadalupe Mountain, undivided (Pliocene)—Predominantly trachydacite lava flows (62 wt% SiO2, 6.3 wt% Na20+K20); contains sparse, small phenocrysts of plagioclase, hypersthene, and augite in a pilotaxitic glassy groundmass; distal lava flows exposed in the Rio Grande gorge and the Red River gorge are highly elongate and individual flows are laterally restricted, typically forming overlapping finger-like lobes characterized by radial cooling fractures and concentric brecciated rapaces where exposed in cross section; flows exposed in the Rio Grande gorge range considerably in thickness from a few meters to several tens of meters; lava flow directions exposed in the Rio Grande gorge appear to be predominantly from east to west, suggesting a primary source area at Guadalupe Mountain; dacite lava flows overlie both Cerro Chiflo dome deposits and lower Servilleta Basalt lava flows in the Rio Grande gorge; 40Ar/39Ar age determination of 5.27 ± 0.05 Ma (sample 11RG08 - M Cosca, personal comm., 2014).

upper (Pliocene)—Tsb can locally be subdivided into the lower Servilleta Basalt (Tsbl), ___________03-14-00-00-map unit—Tdcn—Dacite of Cerro Negro (Miocene)—Dacite of Cerro Negro (Miocene) — Dark gray to black, extensively fractured, two-pyroxene dacite; yielded an 40Ar/39Ar age of approximately 5.7 Ma (McIntosh et al., 2004).

> 03-15-00-00-map unit—Tam—Andesite of Cerro Montoso (Miocene)—Andesite of Cerro Montoso (Miocene) – Dark gray to black, porphyritic, olivine andesite (57.6 wt% SiO2, 8 wt% Na20+K20) lava flows erupted from vents on Cerro Montoso, one of the largest, petrologically uniform, shield volcanoes of the Taos Plateau volcanic field (Lipman and Mehnert, 1979); contains 2-3% phenocrysts of olivine in a microcrystalline groundmass of plagioclase, olivine, augite, and Fe-Ti oxides; the lower slopes of Cerro Montoso in the western map area are commonly mantled in colluvium, and rarely preserve well-developed flow morphology; instead outcrops typically exhibit blocky flow tops and remnants of many discontinuous and aerially restricted flow lobes; Appelt (1998) reported an 40Ar/39Ar age of 5.88 ± 0.18 Ma for a groundmass separate from the west side of Cerro Montoso.

03-16-00-00-map unit-Tvc-Trachyandesite of Cerro Chiflo (Miocene) – Trachyandesite of Cerro Chiflo (Miocene) – Eroded remnants of large lava dome of porphyritic trachyandesite (63 wt% SiO2, 7.7 wt% Na2O+K2O; rock designation based on IUGS classification (Le Bas et al., 1986); formerly described by Lipman and Mehnert (1979) as quartz latite; forms prominent cliff outcrops along the Rio Grande gorge in the northern part of the map area; light brown to gray, weakly to strongly flow laminated, with phenocrysts of plagioclase, hornblende, and sparse otite in a devitrified groundmass; xenoliths of Proterozoic schist, gneiss, and granite common; flow breccias preserved around the margins of dome and ramp structures are common throughout the exposed interior; Appelt (1998) reported 40Ar/39Ar ages of 5.31 ± 0.31 Ma and 5.32 ± 0.08 Ma for groundmass separates from the west and east sides of the dome, respectively; more recently, preliminary 40Ar/39Ar total fusion ag determinations of 10.65 Ma and 9.86 Ma on biotite and hornblende separates, respectively (sample 11RG43 - M. Cosca, personal comm., 2014), are more consistent with previously determined Miocene potassium-argon ages reported by Lipman and Mehnert (1979).

04-00-00-00-heading01-Early-Rift Volcanic & Volcaniclastic Rocks-Early-Rift Volcanic & Volcaniclastic Rocks (Cenozoic) – Early-Rift Volcanic & Volcaniclastic Rock

04-01-00-00-map unit-Tvb-Volcanic deposits of Brushy Mountain (Oligocene)-Volcanic deposits of Brushy Mountain (Oligocene)-Volcanic rocks and deposits consisting primarily of andesite to dacite lava flows and flow breccias and rhyolite block-and-ash flows and ash-flow tuff with volumetrically minor air-fall deposits (Thompson et al., 1986; Thompson and Schilling, 1988). Light tan, poorly welded, lithic-rich, rhyolite ash-flow tuff forms the base of the section near the low saddle of Brushy Mountain; the lower rhyolite contains phenocrysts of plagioclase and altered biotite, light brown altered pumice and angular to subangular vitrophyric inclusions (<0.5 cm to several cm) containing plagioclase phenocrysts and reddishbrown dacite inclusions (2 cm to several cm); locally overlain by thin outflow deposits of Amalia Tuff (unit Tat); post-Amalia Tuff deposits include light-grey to white rhyoli dome deposits including locally, block-and-ash flows, ash-flow tuffs, and air-fall deposits; all deposits are mineralogically similar, containing sanidine, quartz, and minor biotite phenocrysts in a devitrified glass matrix; exposed in the quarry on the south side of Brushy Mountain and the north side of Cerro Montoso; thin (< 2-3 m) andesite lava flows locally overlie rhyolite dome deposits and consist of medium- to dark-brown, porphyritic flows and flow remnants containing olivine, clinopyroxene, and plagioclase phenocrysts, plagioclase glomerocrysts, and minor orthopyroxene microphenocrysts in a fine- to medium-grained trachytic groundmass composed predominantly of plagioclase, clinopyroxene, and Fe-Ti oxides; the upper part of the section consists of light- to dark-gray, aphyric to porphyritic, dacite lava flows and flo breccias containing variable amounts of hornblende, plagioclase, clinopyroxene, Fe-Ti oxides, and minor orthopyroxene, sanidine, sphene, and zircon in a fine-grained to microcrystalline groundmass; lava flows are locally variable in thickness, discontinuous and commonly delineated on the basis of blocky rubble deposits and float; in the Rio Grande gorge, deposits are dominantly andesite to dacite breccias and reworked pyroclastic deposits overlying biotite- and hornblende-bearing dacite lava flows; Zimmerer and McIntosh (2012) reported 40Ar/39Ar age determination of 25.17 ± 0.04 Ma on sanidine from a basal rhyolite of the Brushy Mountain section, and 22.69 ± 0.08 Ma from groundmass concentrates from an andesite lava flow in the upper part of the section.

04-02-00-00-map unit—Tvt—Volcanic deposits of Timber Mountain (Oligocene)-Volcanic deposits of Timber Mountain (Oligocene)-Volcanic rocks and deposits consisting primarily of andesite to dacite lava flows and flow breccias and lesser rhyolite flows and ash-flow tuff (Thompson et al., 1986; Thompson and Schilling 1988); light-brown, lithic-poor, densely welded, rhyolite ash-flow tuff forms the base of the section and contains moderately to highly flattened pumices, phenocrysts of plagioclase, sanidine, quartz, and biotite with subordinate amounts of Fe-Ti oxides, clinopyroxene, and orthopyroxene in a glassy to partially devitrified matrix; rhyolite pyroclastic deposits containing variable amounts of plagioclase, clinopyroxene, Fe-Ti oxides, hornblende, plus or minus biotite, and sanidine xenocrysts; locally contains abundant micropillows of basaltic lava and dacite xenoliths as much as 10 cm in diameter; lower dacite sequence is separated from an upper dacite sequence locally by medium- to dark-brown, porphyritic lava flow remnants containing olivine, clinopyroxene, and plagioclase phenocrysts, plagioclase glomerocryts, and minor orthopyroxene microphenocrysts in a fine- to medium-grained trachytic groundmass composed of plagioclase, clinopyroxene, and Fe-Ti oxides; upper dacite sequence contains medium- to light-gray porphyritic, glassy lava flows and lava dome remnants containing phenocrysts of hornblende, biotite, plagioclase, clinopyroxene, and Fe-Ti

oxides in variable proportion; a whole rock 40Ar/39Ar age of 24.22 ± 0.12 Ma (M. Cosca, personal comm., 2014) was obtained from a basal dacite vitrophyre southwest of the map area.

05-01-00-00-heading02-Volcanic-Volcanic (Cenozoic)-Volcanic

05-01-01-00-00—map unit—Tqi—Latite and quartz latite (Miocene and Oligocene)-Latite and quartz latite (Miocene and Oligocene)-Light tan to gray latite and quartz latite, often stained rust brown, with 15-30% phenocrysts of sanidine, roxene and/or hornblende, sparse quartz, and altered cubes of pyrite; plagioclase phenocrysts to several centimeters in length are present; occurs as dikes up to 20 m wide and elongate intrusive masses north of the D.H. Lawrence Ranch (Lipman and Read, 1989).

5-01-02-00-00 — map unit — Tri — Aphanitic rhyolite (Miocene and cene)—Aphanitic rhyolite (Miocene and Oligocene)—Aphanitic to sparsely orphyritic rhyolite, otherwise similar to Trp (Lipman and Read, 1989).

05-01-03-00-00—map unit—Trpp—Peralkaline rhyolite (Miocene and Oligocene)—Peralkaline rhyolite (Miocene and Oligocene)—Dikes and irregular intrusions of alkali rhyolite and granite porphyry (76-77% SiO2) chemically similar to the Amalia Tuff (Tat) and associated caldera-related rhyolitic lava flows; contains 1-25% henocrysts of quartz and sodic alkali feldspar; locally contains small phenocrysts of redsonite and acmite, especially in the caldera-margin ring dike along Jaracito Canyon and in the Virgin Canyon-Virsylvia Peak area, where the peralkaline rhyolite forms marginal facies of metaluminous biotite-bearing intrusions of granite porphyry within the caldera; dated by K-Ar and F-T methods at about 26 Ma (Lipman and Read,

05-01-04-00-00—map unit—Trp—Porphyritic rhyolite (Miocene and Oligocene)—Porphyritic rhyolite (Miocene and Oligocene)—White to light tan to light gray porphyritic rhyolite typically containing 5-20% phenocrysts of quartz, sanidine, nd sparse plagioclase and biotite; occurs as dikes 1-10 m wide and local irregular and shallow intrusions (Lipman and Read, 1989); generally only observed as float. 05-01-05-00-00—map unit—Tqk—Potassium feldspar quartz latite (Miocene and Oligocene)—Potassium feldspar quartz latite (Miocene and Oligocene)—Coarsely

porphyritic, light-gray quartz latite containing potassium feldspar phenocrysts as long as 5 cm (Lipman and Read, 1989). 05-01-06-00-00—map unit—Tapi—Porphyritic andesite and dacite (Miocene and Oligocene)—Porphyritic andesite and dacite (Miocene and Oligocene)—Fine-grained, dark gray, aphanitic and porphyritic andesite and minor basalt; where present, phenocrysts include hornblende, plagioclase, biotite, and little or no quartz or sanidine

(Lipman and Read, 1989).

05-01-07-00-00-map unit-Tu-Tertiary rocks along the Sangre de Cristo fault, undivided (Miocene and Oligocene) – Tertiary rocks along the Sangre de Cristo fault, undivided (Miocene and Oligocene)—Rocks related to the Questa magmatic system at are undivided along the Sangre de Cristo fault zone; these units were mapped as Quaternary deposits by Lipman and Read (1989), but are herein mapped as unknown bedrock that is generally covered by surficial deposits that are too thin to delineate as map units.

05-01-08-00-00—map unit—Tat—Amalia Tuff (Oligocene)—Amalia Tuff (Oligocene)—Light gray to light brown moderately welded porphyritic, peralkaline, rhyolite ash-flow tuff erupted from the Questa caldera just east of the Village of Questa (Lipman and Reed, 1989); consists primarily of quartz and sanidine phenocrysts in a devitrified matrix; Fe-Ti oxides, titanite, and alkali amphibole phenocrysts are minor, lithic fragments are common; forms low erosional hills of outflow near Brushy Mountain in the Guadalupe Mountain quadrangle; Zimmerer and McIntosh (2012) reported a mean age of 25.39 ± 0.04 Ma based on 13 40 Ar/39Ar laser fusion analyses of Amalia Tuff.

05-01-09-00–map unit–Tatl–Lithic-rich lower facies of Amalia Tuff (Oligocene)-Lithic-rich lower facies of Amalia Tuff (Oligocene)-Nonwelded to partly welded tuff up to 30 m thick, containing as much as 5% fragments of andesitic volcanic s; sparse fragments of Proterozoic rocks present locally; generally grades upward into main unit; locally difficult to distinguish from older tuff of Tetilla Peak (Ttp) (Lipman and Reed, 1989).

05-01-10-00-00-map unit-Tq-Lava flows and domes (Oligocene)-Lava flows and domes (Oligocene) – Massive quartz latite, locally flow layered; commonly gray to altered flows are light red-brown or light gray; intrusive quartz latite (Tqi) locally ifficult to distinguish from flow rocks; maximum thickness is at Latir Mesa, where sections through seemingly single flows or domes exceed 500 m (Lipman and Read,

05-01-11-00-00-map unit-Ta-Andesitic lava flows (Oligocene)-Andesitic lava flow Oligocene)—Purplish-gray to gray, aphanitic to porphyritic andesite lava flows and preccias, with minor interbedded volcaniclastic sediments; phenocrysts include plagioclase and hornblende (Lipman and Read, 1989).

05-01-12-00-00-map unit-Trc-Rhyolite of Cordova Creek (Oligocene)-Rhyolite of Cordova Creek (Oligocene)—Light-tan to light-gray rhyolitic lava flows and domes (74-7% SiO2) containing about 5% phenocrysts of quartz, alkali feldspar, plagioclase, and biotite; commonly massive and devitrified; locally flow laminated; large domes centered at Cordova Creek, Van Diest Peak, and Italian Creek also appear to be sources for main accumulations of tuff of Tetilla Peak; as thick as 250 m at head of Cordova Creek (Lipman and Read, 1989).

05-01-13-00-00 – map unit – Ttp – Tuff of Tetilla Peak (Oligocene) – Tuff of Tetilla Peak (Oligocene) – Quartz-rich, light-colored, weakly welded, rhyolitic ash-flow tuff containing abundant small volcanic fragments; contains 10-30% phenocrysts of quartz, nidine, plagioclase and sparse chloritized biotite; lithic fragments mostly andesite and quartz-bearing rhyolite (Lipman and Read, 1989). 05-02-00-00-heading02-Plutonic-Plutonic (Cenozoic)-Plutonic

05-02-01-00-00 – map unit – Tgy – Lucero Peak pluton (Miocene) – Lucero Peak pluton (Miocene) - White to pale pink, medium to coarse grained equigranular granite toquartz monzonite (Lipman and Read, 1989).

05-02-02-00-map unit—Tgb—Biotite granite (Oligocene)—Biotite granite (Oligocene)—Granitic roof phase of the Rio Hondo pluton emplaced in the Questa caldera at about 26 Ma, during volcanism and caldera formation; medium-grained and equigranular, with sparse aplite and no hornblende (Lipman and Read, 1989).

05-02-03-00-00-map unit-Tgp-Granite porphyry (Oligocene)-Granite porphyry (Oligocene) – Fine-grained porphyritic biotite granite and aplite, texturally transitiona between mapped bodies of granite (Tgb) and intrusive porphyritic rhyolite (Trp) or rhyolite (Tri), especially in the Rito del Medio and Canada Pinabete areas (Lipman and Read, 1989).

05-02-04-00-00 — map unit — Tgd — Rio Hondo pluton (Oligocene) — Rio Hondo pluton groundmass consists of a microcrystalline mosaic of quartz, plagioclase, K-feldspar, (Oligocene) – White to pale, grayish-orange, medium- to fine- grained, massive to locally foliated granodiorite; white to pale orange, aphanitic-porphyritic border facies has quartz phenocrysts and local breccia; has potassium feldspar phenocrysts up to 4 cm in size; generally forms rounded outcrops with abundant grus (Lipman and Read,

05-03-00-00-heading02-Sedimentary-Sedimentary

05-03-01-00-00-map unit-Tvs-Volcanic sedimentary rocks (Oligocene)-Volcanic sedimentary rocks (Oligocene)-Relatively well-bedded and well-sorted volcanic sedimentary rocks of andesitic to rhyolitic composition at many levels in the volcanic sequence; dominantly fluviatile and deltaic deposits; the volcanic sedimentary rocks are locally tuffaceous and interfinger and intergrade in places with the tuff of Tetilla Peak (Ttp); also included are local air-fall and reworked silicic tuff underlying the Amalia Tuff (Trt); exposed thickness nowhere more than about 50 m (Lipman and Read, 1989).

05-03-02-00–map unit–Tps–Prevolcanic sedimentary rocks (Lower Oligocene or Eocene)—Prevolcanic sedimentary rocks (Lower Oligocene or Eocene)—Discontinuous lenses of weakly indurated shale, sandstone, and conglomerate derived from Proterozoic sources; commonly expressed mainly by reddish-brown silty soil; cobble of green quartzite are locally distinctive; outcrops rare, except where baked near granitic intrusion along the Red River; indurated Tertiary sedimentary rocks, which have been correlated with Permian and Pennsylvanian Sangre de Cristo Formation (McKinlay, 1957; Clark and Read, 1972), occur only within areas of Tertiary thermal metamorphism and lack limestone interbeds characteristic of the Sangre de Cristo in

adjacent areas; probably correlative with the Vallejo Formation of Upson (1941) in the Sangre de Cristo Mountains in southern Colorado, and with the Blanco Basin Formation and Telluride Conglomerate in the San Juan Mountains; thickness 0-100 m (Lipman and Read, 1989).

06-00-00-00-00-heading01—Sedimentary Deposits of the Taos Trough—Sedimentary osits of the Taos Trough (Pennsylvanian)—Sedimentary Deposits of the Taos Trough

06-01-00-00-map unit-IPu-Sedimentary rocks of the Taos Trough, undivided (Pennsylvanian)—Sedimentary rocks of the Taos Trough, undivided (Pennsylvanian) – Poorly exposed; greenish, reddish, yellowish, buff, tan, black, and brown; very friable-to-firm; sandy to clayey; thinly to thickly bedded; poorly to moderately well-cemented(?), sandy to clayey siltstone, mudstone, and shale interbedded with mostly greenish and brownish, firm to very strong, poorly to moderately well-sorted, poorly to moderately well rounded, thin- to very thickly bedded, moderately to very well-cemented, quartzose, feldspathic, and arkosic, silty to pebbly sandstone and sandy conglomerate and less common thin- to thick-bedded, gravish and blackish limestone of the Alamitos and Flechado Formations; contains a h assortment of fossils; sandstones commonly contain plant fragments that have een altered to limonite(?); contacts between beds are generally sharp, rarely with minor scour (less than ~20 cm); the lower contact is not exposed in the map area, although to the south, the lower contact is mapped at the top of the Del Padre Sandstone or highest Mississippian carbonate, or at the base of the lowest sedimentary bed where Mississippian rocks are absent; conglomeratic layers in the lower part of the unit locally contain rare, sometimes banded, chert pebbles; equivalent to the Sandia, Madera, and La Posado Formations to the south; fusilinids collected in the Taos quadrangle are Desmoinesian in age (Bruce Allen, personal commun., 2000); may

locally contain minor amounts of Mississippian strata; Miller et al. (1963) measured an incomplete section of 1756 m along the Rio Pueblo near the Comales Campground, and an aggregate thickness of Pennsylvanian strata of >1830 m. 06-02-00-00–map unit—IPcq—Conglomerate (Pennsylvanian)—Conglomerate

Pennsylvanian)—Lenses of quartzite pebble/boulder, clast-supported conglomerate sandy matrix; clasts average about 6 inches in diameter, although some are greater than 2 feet; rounded to well-rounded; very poorly sorted; thickness ranges up to several meters.

06-03-00-00-map unit-IPlst-Limestone (Pennsylvanian)-Limestone (Pennsylvanian)—Light gray limestone in scattered, discontinuous layers; fossiliferous non-fossiliferous; fossils include phylloid algae, crinoids, brachiopods, and other nell fragments; well bedded to poorly bedded; outcrops are typically highly weathered, and locally are highly fractured; limestone represents a very small percentage (<1%) of the volume of Pennsylvanian rock in the map area.

06-04-00-00-00-map unit—IPss—Sandstone (Pennsylvanian)—Sandstone (Pennsylvanian)—Medium-grained to very coarse-grained sandstone; tan to gray; poorly sorted to well sorted; typically non-calcareous; locally contains lenses of conglomerate composed mainly of quartzite pebbles; locally contains wood chips and plant fossils (possibly ferns); thickness ranges from less than a meter to tens of meters.

07-00-00-00-heading01-Metamorphic Rocks of the Sangre de Cristo Mountains—Metamorphic Rocks of the Sangre de Cristo Mountains (Proterozoic) – Metamorphic Rocks of the Sangre de Cristo Mountains

7-01-00-00–map unit—Tu & XYu—Bedrock units of Sange de Cristo Mountains (Miocene to Proterozoic)—Bedrock units of Sange de Cristo Mountains (Miocene to Proterozoic)-In cross section only. Bedrock units of Sange de Cristo Mountains have been combined to form a single unit for the purposes of the cross section. This represents undivided Tertiary rocks along the Sangre de Cristo fault (Tu) and Proterozoic rocks (XYu).

07-02-00-00–map unit–XYu–Proterozoic rocks, undivided (Paleoproterozoic and Mesoproterozoic) — Proterozoic rocks, undivided (Paleoproterozoic and fesoproterozoic)—Supracrustal metamorphic rocks and plutonic and metaplutonic

rocks; shown in cross sections and on Taos Pueblo lands that remain unmapped. 07-03-00-00-00-heading02-Metasedimentary-Metasedimentary (Proterozoic)-Metasedimentary

07-03-01-00-00—map unit—Xq—Quartzite (Paleoproterozoic)—Quartzite (Paleoproterozoic) – White to gray, massive, vitreous quartzite with crossbeds defined by heavy mineral concentrations; pervasively fractured into decimeter-scale, angular zenges by joints, irregular fractures, and bedding (Lipman and Read, 1989); the map cludes the unpublished, detailed mapping of J. Grambling (Univ. of New Mexico, personal communication, 1991) in the San Cristobal Canyon and Cerrito Colorado areas, which supersedes the work of Lipman and Read (1989).

-02-00-00—map unit—Xms—Biotite muscovite schist and gneis (Paleoproterozoic)—Biotite muscovite schist and gneiss (Paleoproterozoic)—Medium to coarse-grained, thinly layered to massive, lustrous, quartz-mica schist and gneiss, ommonly contains sillimanite; locally contains garnet, andalusite, and cordierite (Lipman and Read, 1989).

7-03-03-00-00-map unit-Xlg-Layered gneiss (Paleoproterozoic)-Layered gneiss (Paleoproterozoic)-Conspicuously layered and well-foliated, fine- to medium-grained, biotite gneiss, biotite-hornblende gneiss, hornblende gneiss, and amphibolite; rocks consist of various proportions of quartz, oligoclase-andesine, blue-green hornblende, brown biotite, epidote, and magnetite; layers range in thickness from a few centimeters everal meters and commonly display rootless isoclinal fold noses and variations in kness due to ductile deformation; thin lenses and layers of ferruginous quartzite, magnetite ironstone, and quartz-epidote-calcite marble are commonly interleaved; compositions suggest that many of the layers could have been derived from intermediate volcanic or volcaniclastic rocks; local graded bedding suggests derivation from graywackes, perhaps with a significant volcanic component (Lipman and Read,

07-04-00-00-heading02-Metavolcanic-Metavolcanic (Proterozoic)-Metavolcani

07-04-01-00-00-map unit-Xfg-Felsic gneiss (Paleoproterozoic)-Felsic gneiss (Paleoproterozoic)—Pale gray to orange-brown, micaceous, weakly to moderately iated, quartzofeldspathic gneiss locally grading to micaceous quartzite; commonly terlayered with amphibolite and amphibole gneiss (Lipman and Read, 1989).

07-04-02-00-00-map unit-Xa-Amphibolite (Paleoproterozoic)-Amphibolite Paleoproterozoic) — Thinly layered to massive, fine- to coarse-grained, medium green to dark green to black amphibolite and amphibole gneiss; locally contains calc-silicate iss, biotite-hornblende gneiss, felsic gneiss, and muscovite biotite schist (Lipman and Read, 1989).

07-04-03-00-00-map unit-Xcg-Metaconglomerate Paleoproterozoic)—Metaconglomerate (Paleoproterozoic)—Composed of closely cked 0.5-4-cm angular to subrounded white, blue-gray, and red-brown quartz ebbles in a fine-grained arkosic matrix; interlayered with muscovitic felsic gneiss south of Lama Canyon (Lipman and Read, 1989).

7-04-04-00-00 – map unit – Xvf – Felsic metavolcanic rocks (Paleoproterozoic) – Felsic netavolcanic rocks (Paleoproterozoic)—Fine-grained, light gray, greenish-gray, or pink, massive to strongly foliated, felsic, blastoporphyritic gneiss containing conspicuous 2-5mm ovoid grains of bluish-gray quartz and 1- to 5-mm laths of white feldspar; pidote, and scattered flakes of biotite; feldspar porphyroblasts include both agioclase (oligoclase) and grid-twinned microcline with irregular blotches of albite; composition is similar to rhyolite or rhyodacite; widespread layering and local graded bedding show that a large part of the unit is derived from tuffs or volcaniclastic rocks; zircon from volcaniclastic rock northeast of Gold Hill yielded an upper-intercept concordia age of 1,765 Ma (Lipman and Read, 1989).

07-05-01-00-00-map unit-Zd-Diabase dikes (Neoproterozoic? or early Paleozoic?)—Diabase dikes (Neoproterozoic? or early Paleozoic?)—Nonfoliated, dark gray-green, medium- to fine-grained rocks with well-preserved ophitic texture; 10-20 cm thick with chilled margins (Lipman and Read, 1989).

07-05-00-00-heading02-Metaplutonic-Metaplutonic (Proterozoic)-Metaplutonic

7-05-02-00–00–map unit—Xqc—Quartz monzonite of Columbine Creek (Paleoproterozoic) – Quartz monzonite of Columbine Creek (Paleoproterozoic) – White to gray to pale tan, moderately to strongly foliated quartz monzonite; recrystallized to sugary textured, non-foliated rock near Tertiary plutons; age is 1730 Ma (Lipman and Reed, 1989).

07-05-03-00–00—map unit—Xmi—Mafic and ultramafic rocks (Paleoproterozoic) – Mafic and ultramafic rocks (Paleoproterozoic) – Medium- to coarse-grained dark-green to greenish-gray weakly foliated gabbro and serpentinized ultramafic rocks; gabbro consists of equant clots of hornblende in a matrix of calcic plagioclase, epidote, and sparse quartz; in smaller bodies the gabbro is medium to fine grained, distinctly foliated, and displays chilled margins; original ophitic or tergranular textures are locally preserved and a few bodies display relict cumulus rering; ultramafic rocks are similar to gabbro, except that quartz is absent and plagioclase sparse; mapped only where intrusive into supracrustal rocks; similar rocks

are widespread as inclusions in plutonic rocks where they are mapped as amphibolite (Xa); age of most bodies undetermined, but zircon from gabbro sill west of Gold Hill yielded an upper-intercept concordia date of 1741 Ma, interpreted as the emplacement age (Lipman and Read, 1989). 07-05-04-00-00-map unit-Xtr-Tonalite of Red River (Paleoproterozoic)-Tonalite of Red River (Paleoproterozoic)—Gray to green, medium- to coarse-grained, strongly

foliated, biotite-hornblende tonalite (quartz diorite); consists of a mosaic of recrystallized quartz, plagioclase, and biotite studded with 0.5- to 1-cm ovoid yroclasts of faintly zoned andesine; hornblende forms large irregular sievetured grains and scattered small grains in the mosaic; locally grades into gabbro; slabs and blocks of layered gneiss are locally abundant as inclusions; zircon from tonalite along road to Middle Fork Lake gives an upper-intercept concordia date of 1,750 Ma, interpreted as the emplacement age (Lipman and Read, 1989).

