

Correlation of Units Silver Creek 7.5' quadrangle

DESCRIPTION OF MAP UNITS

Surficial Deposits

- Qsg Sand and gravel (Holocene)—Sand, gravel, and minor mud in modern active river and arroyo channels. 0–20(?) m thick.
Qal Alluvium (Holocene)—Sand, gravel, and mud adjacent to modern arroyo and river channels. Alluvium is typically at or near grade of modern channels, except in local areas. 0–20(?) m thick.
Qae Eolian and alluvial deposits (upper Pleistocene–Holocene)—Eolian sands and loessic silts locally reworked by alluvial processes. Deposits are stabilized by vegetation in most areas. Includes thin, discontinuous eolian veneers on stable upland surfaces that are intimately intertongued with alluvium. 0–2 m thick.
Qt Terrace deposits (upper Pleistocene–Holocene)—Fluvial gravels and sands that form small terraces <15 m above modern stream grades. 1–5 m thick.
Qxy Younger valley-fill deposits (upper Pleistocene)—Gravel, sand, and mud that form terraces less than about 30 m above modern stream grade. Alluvium is representative of deposition in a variety of piedmont environments, including alluvial fans, paleovalleys and arroyo fills, small terraces, fill terraces, and pediments. 0–20(?) m thick.
Qvo Older valley-fill deposits (middle Pleistocene)—Gravel, sand, and mud that form terraces more than about 30 m above modern stream grade. Range of depositional environments is similar to Qxy. At least two aggradational episodes are represented by Qvo. 0–30(?) m thick.
Qpo Older piedmont deposits (middle to lower(?) Pleistocene)—Weakly cemented alluvial gravels and sands that occupy the highest post-Santa Fe Group landscape positions in the northern part of the quadrangle. 1–10 m thick.

SANTA FE GROUP

Deposits of the Santa Fe Group were subdivided based on textural criteria and were then further subdivided on the basis of provenance (Cather, 1997). Deposits containing predominantly Paleozoic and Precambrian detritus are denoted by (s). Those consisting dominantly of volcaniclastic rocks are indicated by (v). Deposits of mixed siliciclastic, calciclastic, and volcaniclastic provenance are denoted by (m).

Sierra Ladrones Formation (Pliocene to lower Pleistocene).

Q1sbcs Conglomerate-sandstone piedmont facies—Characterized by varying, subequal proportions of conglomerate and sandstone. Conglomerate is mostly clast-supported and poorly sorted. Sandstone is typically medium to very coarse-grained, commonly pebbly, and exhibits crossbedding or horizontal stratification. Mudstone is minor unit. Consists of upper and lower tongues. The lower tongue occurs mostly east of the Cerro Colorado fault and exhibits mostly southerly paleoflow. The upper tongue buries the Cerro Colorado fault and shows mostly southerly to easterly paleoflow. Includes pediment veneer on southwest flank of Cerro Colorado that consists of poorly sorted conglomerate composed primarily of Paleozoic limestone and sandstone clasts with subordinate granite. Lower contact has considerable relief in many places and dips more steeply to the east than the constructional top which projects westward to a dissected low relief pediment cut into Proterozoic basement.

Q1sbcs Sandstone-dominated basin-floor facies—Characterized by conglomerate/sandstone ratio of less than 1/2. Conglomerate is clast-supported and occurs in tabular or lenticular units <2 m thick. Sandstone is very fine to very coarse grained and exhibits mostly of horizontal stratification. Mudstone is common and occurs as tabular units which locally compose as much as 20% of the unit. Unit was deposited by low-gradient bedload streams and associated floodplains. Paleoflow was non-systematic.

Q1sbms Basin-floor deposits—Consists of subequal sandstone and mudstone with subordinate conglomerate. Sandstone and pebbly sandstone is horizontally stratified or trough crossbedded and forms broad, lenticular bodies 0.5–3 m thick. Mudstone is indistinctly bedded, forming tabular units 0.5–2 m thick and commonly contains weakly developed calcareous paleosols. Conglomerate is mostly clast supported and occurs in tabular units, commonly incised units as much as 2 m thick. Paleoflow was highly variable in direction.

Q1sbcs Basin-floor travertine and related calcareous deposits—Mesa-forming carbonate deposits interbedded with Q1sbms. To the north and east, unit consists of tabular bedded travertine deposits as much as 5–7 m thick. Outcrops present in thin sections attest to the former presence of ponds in this area. To the south and west, unit thins and becomes more clastic-rich. Near La Jencia Creek the unit grades into carbonate-cemented sandstone and pinches out westward against fluvial deposits of Q1sbms near the Cerro Colorado fault.

Popotona Formation (lower Miocene–upper Miocene)
T1pc Conglomerate piedmont facies—Characterized by conglomerate/sandstone ratio of greater than 2/1. In this and other facies within the Popotona Formation, detritus was derived dominantly from mid-Tertiary volcanic rocks. Siliciclastic and calciclastic deposits in the Popotona Formation occur locally in the stratigraphically highest part of the unit east of the Silver Creek fault near the Rio Salado. These are probably equivalent to the fanglomerate of Ladrón Peak of Machete (1978). Conglomerate is mostly clast-supported, crudely imbricated, and poorly sorted. Matrix-supported debris-flow deposits are common in this facies; in areas where debris-flow deposits are voluminous, they are mapped separately as facies T1pc. Sandstone in T1pc is medium to very coarse grained and commonly exhibits crossbedding or horizontal laminations. Mudstone is rare, occurring mostly as thin discontinuous drapes. Paleoflow was eastward.

T1pc Debris-flow-dominated facies—Characterized by a dominance of very poorly sorted conglomerate that is matrix-supported and typically very well indurated. Clasts are virtually all derived from Tertiary volcanic rocks. Fills paleovalleys in the basal part of the Popotona Formation in the Lemitar Mountains area (Chamberlin, 1982; Cather et al., 1994a).

T1ps Conglomerate-sandstone piedmont facies—Characterized by conglomerate/sandstone ratio greater than 2/1. Conglomerate is mostly clast-supported and poorly sorted. Sandstone is fine to very coarse grained and commonly is horizontally stratified or trough crossbedded. Mudstone is minor. Paleoflow was eastward.

T1ps Sandstone-dominated piedmont facies—Characterized by conglomerate/sandstone ratio less than 1/2 and sandstone/mudstone ratio greater than 2/1. Sandstone is dominantly horizontally stratified with subordinate trough crossbedding. Conglomerate is mostly clast-supported and occurs as drapes and as tabular units less than a meter thick. Paleoflow was generally eastward.

T1ps Transitional playa-margin facies—Characterized by sandstone/mudstone ratio of between 2/1 and 1/2. Sandstone is horizontally laminated and forms thin tabular beds (<0.5 m). Sandstone is intimately interbedded with tabular mudstones that are structureless and dominantly red-brown in color. Conglomerate is rare. The playa-margin facies represents interfingering of distal piedmont and sand-flat deposits with playa mudstone. Near Silver Creek, paleoflow was generally northward due to deflection of regional easterly paleoflow by the Silver Creek volcano.

T1pm Playa facies—Characterized by a dominance of mudstone (sandstone/mudstone ratio is less than 1/2). Mudstone is mostly red-brown with uncommon greenish-gray zones that are parallel to bedding. Bedding is generally indistinct, although horizontal lamination was occasionally noted. Sandstone is medium to very fine grained and occurs as thin tabular beds (<0.5 m). Conglomerate is virtually absent. Gypsum is common. Deposited in large playa system east of the Silver Creek fault and by ponding adjacent to the Silver Creek volcano.

Tpa Silver Creek andesite—Pyroxene- and plagioclase-bearing basaltic andesite flows and tephra. Erupted by the Silver Creek volcano that was subsequently cut by the Silver Creek fault (Weber, 1971). Volcano consists of a shield volcano that buried an earlier tuff cone that consisted of phreatomagmatically erupted greenish basaltic glass and recycled detritus (Tpo) from the underlying Popotona Formation. Radial dikes (Td) are exposed in the core of the volcano. Volcano has been K-Ar dated at 16.2 ± 1.5 Ma by Weber (1971). ⁴⁰Ar/³⁹Ar dates of 15.33 ± 0.07 and 15.49 ± 0.20 have recently been obtained from a lava and a feeder dike, respectively, from the Silver Creek volcano (R. M. Chamberlin and W. C. McIntosh, 2003, written commun.).

UPPER EOCENE–OLIGOCENE VOLCANIC, INTRUSIVE AND VOLCANICLASTIC ROCKS

Tlp La Jara Peak Basaltic Andesite (upper Oligocene)—Mafic flows with associated tephra and minor eolian sandstone. Flows exhibit trachytic textures locally and phenocrysts are clinopyroxene and plagioclase ± amphibole. Individual flows are 1–5 m thick. Not dated in quadrangle; age range is ~27–24 Ma regionally.

Ti Basaltic andesite dikes (upper Oligocene)—Mafic dikes 0.5–1.5 m thick. Undated on quadrangle, but regionally are associated with La Jara Peak Basaltic Andesite. En echelon dikes exposed within fault slices of the Cerro Colorado fault suggest a dextral component to this moderately dipping normal fault.

Tvp Vicks Peak Tuff (upper Oligocene)—Distal outflow of crystal-poor, high-silica rhyolitic ignimbrite erupted from Nogal Canyon caldera in the southern San Mateo Mountains. Age is 28.6 Ma (McIntosh et al., 1991).

Tm Hells Mesa Tuff (lower Oligocene)—Crystal-rich ignimbrite containing crystals of biotite, quartz, plagioclase, and sandine. Erupted from the Socorro caldera at 32.1 Ma (McIntosh et al., 1991).

Tds Spears Group (upper Eocene–lower Oligocene)—Grayish-purple volcaniclastic sandstone, mudstone, and conglomerate that interfinger regionally with extrusive volcanic rocks in the Mogollon-Datil (Cather et al., 1994a). Only exposed in a small fault block along Cerro Colorado fault on north side of the Rio Salado.

PALEOZOIC ROCKS

Py Yeso Formation (Permian)—Gray to orange sandstone, siltstone, limestone and gypsum. Exposed in fault blocks along Cerro Colorado and Silver Creek faults.

IPb Bursum Formation (Pennsylvanian–Permian)—Interbedded limestone and red sandstone–siltstone exposed in fault slices along the Cerro Colorado fault.

IPm Madera Limestone (Pennsylvanian)—Gray to brown fossiliferous limestone and shale, and sandstone. About 600 m thick.

Mk Kansas Limestone (lower Mississippian)—Gray to light brown limestone (commonly crinoidal), sandstone, and shale. Paleokarst features are common. These rocks are strongly altered (oxidized iron) in places, particularly where the Cerro Colorado and Ladrón faults intersect. About 30 m thick; thins to north.

PROTEROZOIC ROCKS

Ys Mesoproterozoic(?) granite—Medium-grained unfoliated granite and associated granitic pegmatites. This granite appears to intrude Xgn and is only seen northeast of the Ladrón Fault.

Xgn Paleoproterozoic medium-grained granite gneiss—Generally strongly foliated/lineated. Two-feldspar granite gneiss is pink due to higher proportion of K-spar to plagioclase. Biotite is often altered to chlorite. Appears to intrude Xmg and older rocks. This gneiss does not appear to be a higher-strain variety of Xmg because the primary foliation is so different as is the proportion of biotite (much more in Xmg) and the finite strain appears to be comparable between Xgn and Xmg.

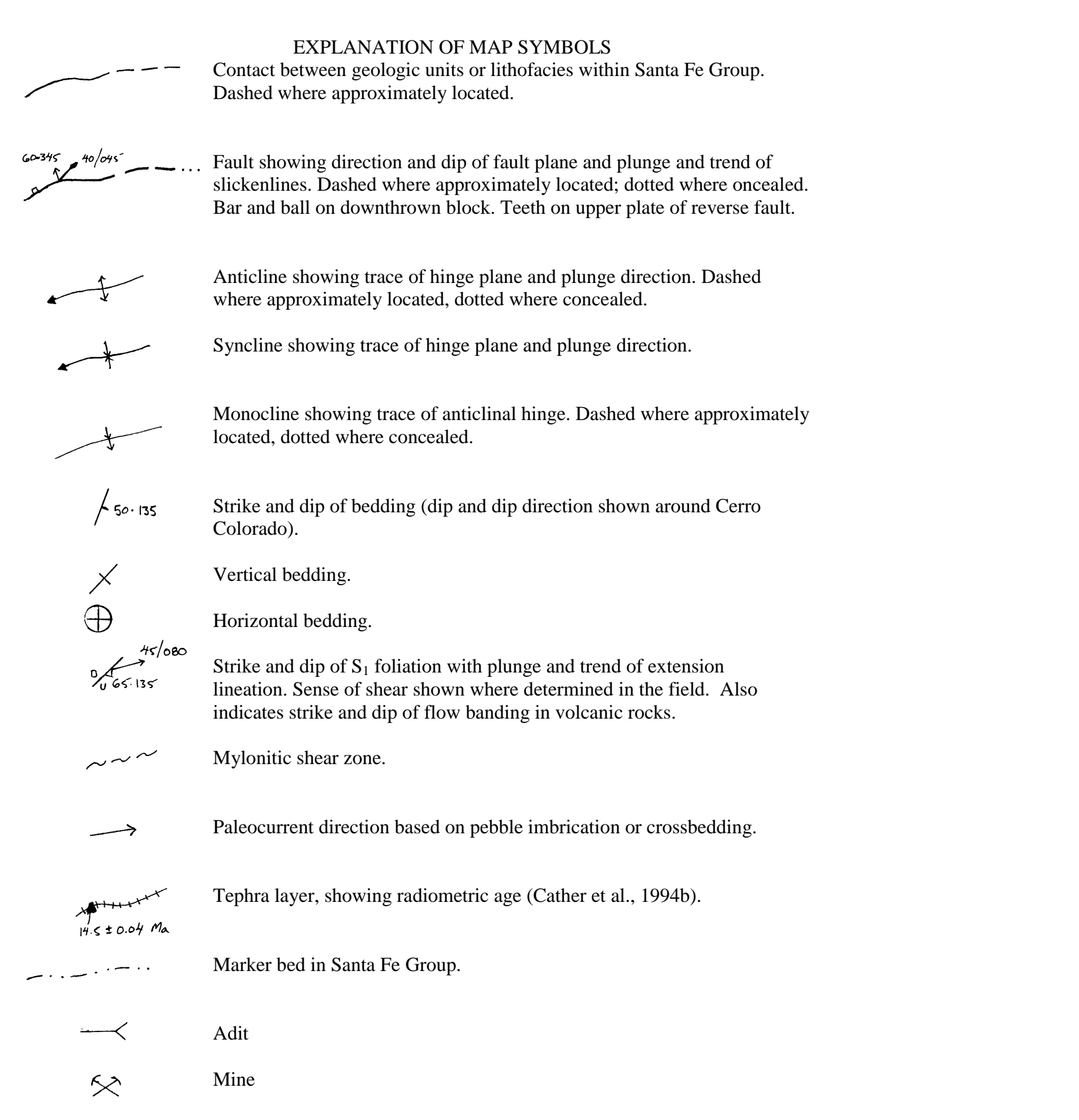
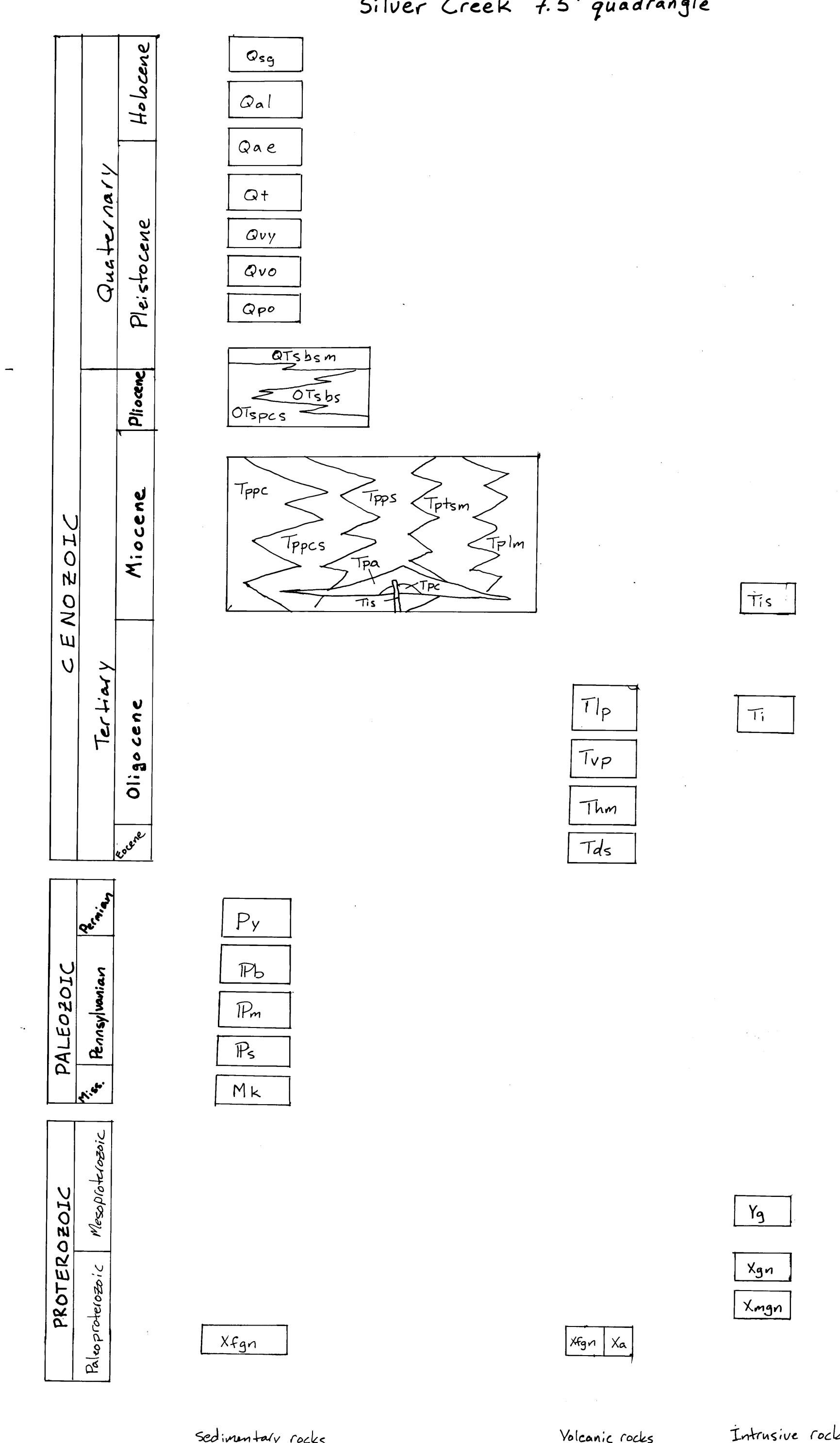
Xmgn Paleoproterozoic megacrystic grano-diorite gneiss—Two-feldspar gneiss is very similar in composition to Xgn except that the proportion of biotite is much higher. Composition ranges from granitic to grano-diorite probably reflecting the local composition of assimilated country rock (Xa and Xgn). The strongly foliated/lineated gneiss is generally pink due to large K-spar crystals that are plastically deformed into ribbons 10–20 cm long. Shear sense is difficult to determine in outcrop, probably reflecting a large component of pure shear. However, K-spar porphyroclasts viewed parallel to lineation and perpendicular to foliation occasionally indicate top-west thrust sense in present orientation/shear.

Xgn Paleoproterozoic fine-grained gneiss—Felsic to intermediate gneiss of uncertain protolith. Gneiss is generally felsic, composed mostly of K-spar and quartz and is quite fine-grained. In places, mylonite zones are present within this gneiss, however it appears that the fine grain size is primary. It is generally strongly foliated/lineated to a similar degree as Xgn, Xa, and Xmgn with respect to the finer grain size. This rock is intimately interbedded with amphibolite with sometimes indistinct contacts leading to more intermediate compositions. The association with amphibolite suggests that the protolith was a felsic volcanic rock or sedimentary rock derived from rhyolites in a bi-modal volcanic field. However, no primary volcanic or sedimentary features were observed in the field.

Xa Paleoproterozoic amphibolite—Amphibole-plagioclase mafic schist. Generally strongly foliated similar to the Xgn and is interbedded with. Also appears as sometimes large screens within younger metaplatic rocks. In places, what appear to be relict amphibolites are filled with plagioclase suggesting that these rocks represent basal flows rather than intrusive rocks.

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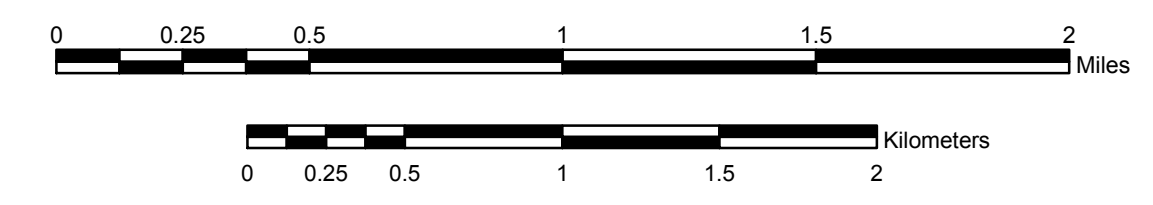


Geologic Map of the Silver Creek 7.5-min quadrangle

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May, 2003

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COMMENTS TO MAP USERS

A geologic map displays information on the distribution, nature, orientation, and age relationships of rock and deposits and the occurrence of structural features. Geologic and fault contacts are irregular surfaces that form boundaries between different types or ages of units. Data depicted on this geologic quadrangle map may be based on any of the following: reconnaissance field geologic mapping, compilation of published and unpublished work, and photogeologic interpretation. Locations of contacts are not surveyed, but are plotted by interpretation of the position of a given contact onto a topographic base map; therefore, the accuracy of contact locations depends on the scale of mapping and the interpretation of the geologist(s). Any enlargement of this map could cause misunderstanding in the detail of mapping and may result in erroneous interpretations. Site-specific conditions should be verified by detailed surface mapping or subsurface exploration. Topographic and cultural changes associated with recent development may not be shown.

This work was performed under the STATEMAP component of the USGS National Cooperative Geologic Mapping Program. Funding was provided by the U.S. Geological Survey and the New Mexico Bureau of Geology and Mineral Resources, a division of New Mexico Tech.
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