

Ocp Camp Rice Formation, piedmont-slope facies—Deposits associated with

piedmont slopes that grade into or interfinger with ancestral Rio Grande fluvial

deposits (Qcf); weakly to moderately cemented boulder to cobble fan deposits

and erosion-surface veneers near mountain fronts grading to gravelly silt, loam,

gravel and grades downward into gravel and sand lenses almost everywhere glomerate facies (QTcc), undivided veak saline soil is generally present in upper horizons; approximately 20-25 Camp Rice Formation, basin-floor deposits, undifferentiated—Includes Oct Obfs Basin-floor, fluvial-fan deposits of Mimbres River, sand and gravel fastrata as well as somewhat older basin-floor sediments exposed in Camel Mountain escarpment; the older strata consist of yellow, tan, and pale-orange, noderately cemented sandstone and siltstone and red to gray mudstone, plains and fans and that stand 2-3 ft (0.6-1 m) above the interfluve lowlands:

largely of eolian and alluvial-flat origin, respectively; a thickness of approxisheet of intertonguing Qbff and Qbfs that is 20-25 ft (6-8 m) thick mately 100 ft (30 m) is present in Camel Mountain escarpment but is mostly QI Playa deposits—Loam, silt, and clay of small, nonalkaline playa lakes and Mimbres Formation, younger piedmont-slope deposits—Boulder to cobble sions; as much as 15 ft (4.5 m) thick fan deposits and erosion-surface veneers near mountain fronts grading to Qs | Eolian deposits—Dunes and irregular hummocks of quartz sand, especially

gravelly silt, loam, or clay on piedmont toeslopes; weakly to moderately co mented; surficial gravelly layers generally contain pedogenic calcrete as much as 5 ft (1.5 m) thick, whereas finer-grained facies may have thick red clay Bt Mountains. The sand is derived largely from Qbfs and Qbff and from basin horizons; Qm is correlative with Qcp of the Camp Rice Formation, the Qm floor facies of the Camp Rice Formation (QTcb); the symbol Qs/Qb or Qs/Qcp symbol used west of ancestral Rio Grande drainage systems; generally less etc. shows the rock unit that the sand has buried; as much as 10 ft (3 m Mimbres Formation, piedmont-slope deposits, undifferentiated—Unit in-

Qpu Older piedmont-slope deposits (Qpo) and correlative Camp Rice Forcludes Qm as well as older well-cemented conglomerate and fanglomerate mation, piedmont-slope facies (Qcp) and Mimbres Formation, younger that may be as old as late Pliocene; unit is correlative with QTcu of Camp piedmont-slope deposits (Qm), undifferentiated Rice Formation; 50 ft (15 m) or more exposed, probably considerably thicker Opo Older piedmont-slope deposits—Fan and terrace deposits and erosion-sur face veneers on piedmont slopes graded to Pleistocene closed-basin floors: Basalt of West Potrillo Mountains-Alkali-olivine basalt flows (Qb) and cinder mostly weakly consolidated gravel and sandy gravel grading downslope to cones (Qbc); K-Ar radiometric age of 1.2 Ma from flow on northern edge of gravelly loam; thin surficial and buried horizons of soil-carbonate and clay accumulation; gravelly carbonate horizons are commonly indurated and form

Basalt flows, cones, and dikes—Black to gray alkali-olivine basalt flows (Tb) volcanic field; radiometric ages of 5.2 ± 0.1 , 3.9 ± 0.2 , and 3.0 ± 0.07 Ma (Seager Qpa Younger (Qpy) and older (Qpo) piedmont-slope deposits Older axial river deposits—Mostly gravel and sand interbedded with sheets Fanglomerate—Brown to reddish-brown boulder conglomerate and breccia of pink and pale-orange basin-floor loam, silt, and clay; moderate to strong and tan, gray, or purplish-gray to maroon-gray conglomeratic sandstone, sand soil carbonate in gravelly zones, less strong in fine-grained facies; as much stone, and siltstone derived from a wide variety of local volcanic rocks and,

near the East Potrillo Mountains, derived from Paleozoic or Cretaceous sedi-Obfo Older basin-floor deposits—Pink, tan, cream, and pale-orange silt, loam, and mentary rocks and Precambrian granite; well cemented in lower parts and fine sand on basin floors that are graded either to Qpo (upper Pleistocene) or moderately cemented in upper part; unit intertongues downward with Tba tilted; broken by range-boundary faults; locally silicified along faults; correlative Qm (middle Pleistocene) piedmont slopes; soil carbonate is moderate to strong with bulk of the Gila Conglomerate and Santa Fe Group; at least 1,000 ft (300 and locally calcrete, a few feet thick, may have developed; as much as 200 Obfg Older gypsiferous basin-floor deposits—Light-gray to cream gypsiferous Tsrc Santa Fe Group, Rincon Valley Formation—Reddish-brown to tan fanglomerate, conglomeratic sandstone, and sandstone of Sierra de las Uvas-Sleeping

entysix Draw drainage; as much as 20 ft (6 m) exposed Lady Hills area; derived from a variety of local volcanic rocks; unconformably overlies Tu; tilted; faulted; at least 500 ft (152 m) thick (Clemons, 1976) Ot Maar-rim tuff and breccia—Crossbedded juvenile basaltic lapilli, accretionary Uvas Basaltic Andesite—Brown, gray, and black basaltic-andesite, hornapilli, and disaggregated sand and silt from the Camp Rice Formation on the ende-latite, and andesite flows; vesicular, amygdaloidal to dense with platy rims of Potrillo, Mt. Riley, and Malpais maars; upper mantle and lower crusta pinting; includes thin tongues of brown sandstone and conglomeratic sandxenoliths as well as fragments of upper crust conglomerate, limestone, and stone; intrusive masses thought to be correlative with Uvas Basaltic Andesite lows are labeled Tui; 200–500 ft (61–152 m) thick; radiometric ages of 25.9 ± 1.5 , as 50 ft (15 m) thick; K-Ar radiometric age from associated Potrillo maar basal: 26.1 ± 1.4, and 27.4 ± 1.2 Ma (Clemons and Seager, 1973; Clemons, 1976, of 0.18 ± 0.03 Ma (Seager et al., 1984) contrasts with age of 24 ka for the similar Kilbourne Hole maar determined by comparative soil studies (Gile, Basaltic andesite—Brown, gray, red, and black tholeiite*, basaltic-andesite,

ments; commonly reworked by wind; surficial-carbonate accumulation (K) and

locally reddish-brown horizons of clay accumulation (Bt); carbonate horizons

are commonly indurated, forming pedogenic calcrete zones as much as 5 ft

(1.5 m) thick; unit overlain by veneer of Qs or locally by upper Quaternary

TD 4011'

no vertical exaggeration

playa or alluvial-flat deposits (QI, Qbf); as much as 50 ft (15 m) exposed

latite, and andesite flows; vesicular, scoriaceous, amygdaloidal to dense with platy jointing; some glassy flows; local interbeds of brown sandstone, con-Qba Aden Basalt—Alkali-olivine basalt flows associated with Aden volcano (crater), flows overlie La Mesa surface; K–Ar radiometric age of 0.53 ± 0.04 Ma (Seager et al., 1984) contrasts with 3 He surface-exposure ages from 15.7 \pm 2 to 18.2 \pm 3 Lower fanglomerate—Gray to tan boulder conglomerate, tan tuffaceou Oct | Camp Rice Formation, sediments associated with La Mesa surface—Basinsandstone and siltstone, and brown conglomeratic sandstone, consisting of a variety of local volcanic detritus; generally well cemented; probably correlfloor sediments, the constructional top of which is La Mesa geomorphic surative with "early Gila" strata: intertonques upward with Tba in the Cedar face (shown on geologic map with diagonal line pattern). Consists generally of sand, silt, loam, or clay deposited in fluvial, playa, and alluvial-flat environ

Mountains region; approximately 400 ft (122 m) thick Basaltic-andesite conglomerate—Brown to black basalt boulder conglom erate derived largely from Tba; poorly exposed in the Greasewood Hills where unit is associated with Tba and Tlg; 100 ft (30 m) or more thick Intermediate- to silicic-composition lavas of West Potrillo Mountains-

banding; unit is probably correlative with latite and rhyolite of Greasewood

Mountains. The rhyolite probably varies in age from place to place, but most is probably middle to late Oligocene Tuff of Johnson Mountain—Pinkish-gray, dacitic* crystal ash-flow tuff conaining 0-2% quartz, 1-6% sanidine, 2-5% biotite, and as much as 48% microperthite, the latter probably a product of potassium metas trix is welded devitrified pumice and glass shards; at least 200 ft (61 m) thick Purple, maroon, and light- to medium-gray, porphyritic latite to dacite* flows;

Tsb Younger tuffs, flows, and sedimentary rocks of Carrizalillo Hills—Brown, yellow, tan, and red, rhyolitic volcaniclastic rocks, breccia, and minor tuff, all intensely altered (K-metasomatism) and complexly faulted and veined by quartz and carbonates; some intermediate-composition flows (TII) are inter as is a single bed of fresh-water limestone approximately 20 ft (6 m) thick:

Andesite-latite flows—Greenish-gray to brown andesite to latite flows; fine

Limestone conglomerate—Reddish-brown to gray limestone boulder congrained, aphyric to slightly porphyritic; only two outcrops at southwest end omerate derived primarily from Paleozoic and Lower Cretaceous(?) rocks and to a lesser extent from Precambrian granite and intermediate-composition volcanic rocks; locally sandstone, siltstone, and mudstone dominated by vol-

be a tongue within Tli; approximately 100 ft (30 m) or more thick rmediate-composition volcanic rocks—Light- to dark-gray, reddish-brown, Silicic plutonic intrusions—Coarse-grained granite to quartz monzonite of purplish-gray, and maroon flows ranging in composition from andesite* to dacite* to latite; aphyric to moderately porphyritic; intensely altered (K-metasomatism) in the Carrizalillo Hills but relatively unaltered in the Cedar Mountains where andesine, pyroxene, and hornblende phenocrysts are in a trachytic

matrix; at least 500 ft (152 m) thick Fuff of Carrizalillo Hills—Reddish-brown, rhyolite* crystal ash-flow tuff; 3-% quartz, 8-17% sanidine, minor biotite, and as much as 11% xenocrystic K-feldspar in a matrix of pumice fragments and glass shards; extensive Kmetasomatism: densely to moderately welded; unit has two members sep arated by a few feet of air-fall tuff or epiclastic rocks; the upper member, less than 100 ft (30 m) thick, is a simple cooling unit; the lower member, as much as 300 ft (91 m) thick, is a compound cooling unit

with interbedded epiclastic sandstone; 0-300 ft (0-91 m) thick; 3) ash-flow

Intermediate-composition volcanic rocks-Gray, purplish- to maroon-gray

laterally with Rubio Peak Formation: maximum thicknesses approximately 3.50

plagioclase, sanidine, quartz, and biotite phenocrysts; abundant white, flat-

tened pumice; crops out in Little Florida Mountains; 200 ft (61 m) thick

ft (1,067 m) (Clemons, 1979, 1976, 1977); K-Ar radiometric age of 37-43 Ma

Tuff of Lonesome Canyon Draw—Cream, light-gray, and pale-brown ashflow tuff containing small volcanic lithic fragments and quartz and sanidine atite, dacite*, or andesite* flows. Andesine and sanidine phenocryst laths are crystals; extremely welded and dense; aspect ratio of pumice is 12:1; found in a trachytic groundmass of plagioclase and devitrified glass; K-alteration in only in west-central Tres Hermanas Mountains; at least 300 ft (91 m) thick rizalillo Hills; similar, possibly correlative rocks in Burdick Hills are labeled T/2: 30-70 ft (9-21 m) thick Tuff of Big Tank Draw-Brownish-gray ash-flow tuff with abundant lithic frag nents as much as 2 inches in diameter; welded; many cavities from weathwer crystal tuff of Carrizalillo Hills—Pale purplish-gray, medium-grained, ered pumice and lithic fragments; platy foliation; basal conglomerate (Tc) consists welded crystal ash-flow tuff, dacite* in composition; 5% quartz, <1% sani-dine, 2–3% bronze biotite, and as much as 37% altered xenocrystic K-feldspar of vollcanic boulder clasts in mud-sand matrix; found only in west-central Tres

Hermianas Mountains; Ttf and Tc total thickness at least 500 ft (152 m) in a matrix of devitrified glass shards; K-alteration in Carrizalillo Hills; at least Rhyolite of southern Tres Hermanas Mountains—Yellow, cream, and gray, nassive to delicately flow banded rhyolite to quartz-latite dome-flow(?) com-Didest exposed tuffs and sedimentary rocks of Carrizalillo Hills—Cream, plex with basal purplish-gray tuff and breccia (Trtb); similar rocks in northin, pale-purple, and yellowish-tan air-fall tuff and breccia and epiclastic strata eastern Tres Hermanas Mountains may be correlative with Trt; at least 1,000 rhyolitic in composition; variably silicified and altered; at least 100-200 ft (30-

afu Undifferentiated ash-flow tuff—Gray to red, lithic-vitric ash-flow tuff in the West Potrillo Mountains and reddish-brown crystal ash-flow tuff west of Black Mountain; 50-100-ft-thick (15-30-m-thick) sections exposed in each area

Latite-rhyolite of Greasewood Hills-Flow-banded to nearly massive, dense

o vuggy rhyolite* to latite flows and intrusions(?); from medium- to dark-gray

to purplish- or reddish-gray; slightly porphyritic; rocks of Midway Butte and

Jog Tank Hill are probably correlative as is unit Tlis of the West Potrillo Moun-

tains; age relationship of Tlg with Tba and Tbc in the Greasewood Hills is not

of Adlen Hills; possibly correlative with Tvi; at least 300 ft (91 m) thick

Andesite of Little Florida Mountains-Medium-gray to brownish-gray flows

and intrusions of andesite and dacite; aphanitic; aphyric; at least a few hundred

res Hermanas stock; biotite-latite porphyry to monzonite of Providence cone

leucomonzonitic facies of the diorite-monzonite intrusion west of Camel

Intermediate-composition intrusions—Gray diorite to hornblende-monzo-

Eagle Nest hills, digrite-andesite of northern Tres Hermanas Mountains, and

porphyritic andesite of Mt. Riley-Mt. Cox; medium- to coarse-grained except

are fine to medium grained; coarse-grained intrusions are small stocks or thick

Hills intrusion (at the Calumet mine) a subvolcanic plug

dikes; Mt. Riley-Mt. Cox intrusion is probably a volcanic dome and Carrizalillo

or Tres Hermanas, Carrizalillo Hills, and Mt. Riley-Mt. Cox intrusions, which

nite porphyry of Camel Mountain area, hornblende-monzonite porphyry of

Bell Top Formation, undifferentiated Rhyolite intrusions—Tan, white, and cream, massive rhyolite generally lackng or with only weakly developed flow banding; contains abundant euhedral dimentary rocks and tuffs—White, tan, and gray tuffaceous sandstone, juartz phenocrysts in Bisbee Hills; unit is present as eroded dome-plugs in nudstone, and conglomerate and rhyolitic air-fall tuff above and below Tbte Bisbee Hills but elsewhere as dikes, small plugs, and irregular, small intrusive 0-300 ft (0-91 m) thick in southern Good Sight Mountains and Sleeping masses; age probably varies from place to place but probably is middle to late Oligocene; dikes in Florida Mountains are 29.1 ± 1.3 Ma (Clemons, 1985) Ash-flow tuff 6—Grayish-red, densely welded crystal-vitric ash-flow tuff; Dacitie flows of Burdick Hills-From medium- to light-gray to brown dacite* -7% quartz, 2-13% sanidine, 0-15% plagioclase, 0-3% biotite plus horn-

flows; aphyric to porphyritic; saussuritized plagioclase phenocrysts common blende; simple cooling unit; 0-100 ft (0-30 m) thick; radiometric age of in some flows; at least a few hundred feet thick 33.5 ± 0.12 Ma (McIntosh et al., 1991) Ttv Vitric tuff of Burdick Hills—Pale grayish-red, fine-grained, densely welded, Lower tuffs and epiclastic rocks—In descending order the unit includes ritric ash-flow tuff with scattered pumice; interbedded in Ttu; approximately) ash-flow tuff 4; porous to densely welded, pale grayish-red vitric-crystal ash-flow tuff; 0–180 ft (0–55 m) thick; radiometric age of 34.96 ± 0.04 Ma (McIntosh et al., 1991); 2) white to yellowish-gray air-fall tuff and breccia

an, rhyolitic air-fall tuff and tuffaceous sandstone, siltstone, and shale; forms tuff 3; moderately to poorly welded, pinkish-gray, vitric ash-flow tuff; 0–350 ft (0–107 m) thick; radiometric age of 35.7 ± 0.02 Ma (McIntosh et al., 1991); two members separated by Tty: upper member contains abundant agate. which has been quarried extensively in the Burdick Hills; lower member is and 4) white to brownish-gray and pale yellowish-brown air-fall tuff and mostly covered but may be as much as 500 ft (152 m) thick; upper (agatebreccia with interbedded conglomerate and sandstone, as much as 100 ft (30 ft) thick. In the southern Good Sight Mountains only unit 1 is present; bearing) member is approximately 100 ft (30 m) thick all units are well developed in the Sleeping Lady Hills (Clemons, 1976) Crystal tuff of Burdick Hills-Pink to brownish-gray, rhyolitic to dacitic, crystalric welded ash-flow tuff; phenocrysts include 2-11% quartz, 5-11% sani-Basaltic andesite within Tbl-Gray, vesicular basaltic-andesite dikes and dine, 13-32% plagioclase, and 1-5% biotite in a matrix of pumice and devit-

Upper tuffs and sedimentary rocks of Burdick Hills-White, light-gray, and

ified glass; simple(?) cooling unit; at least 250 ft (76 m) thick

Lower tuffs and sedimentary rocks of Burdick-Bisbee Hills-Pink, cream, latite, dacite, and andesite flows, intrusions(?), and laharic(?) breccia; aphyric nd gray, rhyolitic crystal ash-flow tuff and associated epiclastic, tuffaceous to moderately porphyritic; generally fine grained; locally includes epiclastic sandstone, shale, and conglomeratic sandstone; thin beds of travertine and rocks derived from Tvi; exposures are limited to Mt. Riley and Camel Mountain fresh-water limestone appear to be interbedded with the unit but may be region, but correlative rocks may be units Ta and Tal; at least several hundred younger units related to major faults; approximately 300-400 ft (91-122 m) Palm Park Formation—Gray, red, and purple, tuffaceous epiclastic rocks inlegabreccia in rhyolite—Boulders to house-size blocks of Tj, Trf, and other cluding lahar breccia and minor flows of intermediate composition; correlative ks in a matrix of greenish-yellow, devitrified rhyolite and rhyolitic glass and

in air-fall tuff and breccia; small domes, plugs, and dikes of Trf locally intrude the unit or are graditional into the megabreccia; much of the unit probably ents a proximal tuff-cone facies; exposed thickness approximately 500 Trs Sugarlump Formation—Grayish-pink vitric-crystal ash-flow tuff with 7–30% Flow-banded rhyolite—Pink, cream, brown, tan, yellow, or gray flow-banded rhyolite*, locally spherulitic, forming small plugs, moderate to large domes (Red Mountain), dikes, or irregular intrusive masses in the Carrizalillo Hills,

es Hermanas Mountains, Florida Mountains, Sierra de las Uvas, and Potrillo

Rubio Peak Formation—Gray, purple, and maroon lava flows, lahar breccia, and epiclastic rocks of intermediate composition; generally propylitized; 1,600 ft (488 m) thick in Florida Mountains (Clemons and Brown, 1983), at least 2,000 ft (610 m) thick in the Good Sight Mountains (Clemons, 1979), and as much as 6.500 ft (1.980 m) thick in the subsurface near Deming (R. E. Clemons pers. comm. 1984); K-Ar radiometric ages of 44.7 ± 1.9 , 38.0 ± 1.5 , 38.0 ± 1.9 , 37.6±2.0, 36.4±2.3, and 32.6±2.1 Ma (Clemons, 1979, 1982b)

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TKs Lower Tertiary-Upper Cretaceous sedimentary rocks, undifferentiated-Gray, purplish-gray, maroon, tan, brown, and olive-brown sandstone, shale, and siltstone and pale-red mudstone derived in part from intermediate-composition volcanic rocks; also gray to reddish-gray cobble to boulder conglomerate consisting of clasts ranging from Precambrian to Lower Cretaceous; exposed in Mt. Riley-East Potrillo Mountains area; correlative with Love Ranch Formation and probably with basal parts of the Rubio Peak Formation; at least

Lobo Formation—Reddish-brown to grayish-red cobble-boulder conglomerate containing clasts of Paleozoic limestone, chert, and sandstone and Precambrian-Cambrian granite, syenite, and quartz syenite. Syntectonic to osttectonic with Laramide orogeny; approximately 100-500 ft (30-152 m) thick in Florida Mountains (Clemons and Brown, 1983; Clemons, 1995); TKI? in Tres Hermanas Mountains, approximately 250 ft (76 m) thick, probably is Lobo but could possibly be Lower Cretaceous Glance equivalent. At Granite Hill near Eagle Nest, the Lobo Formation consists of approximately 787 ft (240) m) of sandstone and conglomerate that overlies Precambrian granite noncon-

formably (Seager and Mack, 1990) Marbleized limestone—Formation unknown

ower Cretaceous rocks, undifferentiated—In the East Potrillo Mountains the section includes, from top down, Mojado(?) Formation: gray, crossbedded quartzite at least 50 ft (15 m) thick; U-Bar Formation: calcareous siltstone at least 50 ft (15 m) thick, gray, massive limestone 435 ft (133 m) thick calcareous siltstone and limestone 268-298 ft (81-90 m) thick, rudistid limestone 20-45 ft (6-13 m) thick, arkosic sandstone and pebble conglomerate 85-220 ft (25-66 m) thick, and oolitic limestone 45-110 ft (13-33 m) thick and Hell-to-Finish Formation: calcareous siltstone and fine sandstone 400-550 ft (119-168 m) thick and limestone cobble conglomerate 3-133 ft (1-40 m) thick. Total thickness is approximately 1,000 ft (305 m). Correlative partial sections include approximately 900 ft (274 m) of mostly U-Bar Formation at Eagle Nest and approximately 200 ft (61 m) of mostly Hell-to-Finish Formation at West Lime Hills. Strata are largely Aptian and Albian in age (Seager and

Yeso Formation-San Andres Formation, undifferentiated-Upper beds (San ndres) are massive to medium-bedded, gray limestone with local chert, pelecypods, crinoids, and corals. Lower beds (Yeso?) are light- to dark-gray, fineto medium-grained dolomitic limestone, with silty laminations, spar-filled yugs, and limestone breccia, interbedded with minor brown siltstone and yellow, fine-grained siltstone. Approximately 1,000 ft (305 m) thick in East Potrillo Mountains (Seager and Mack, 1994); a few hundred feet are exposed in the West Lime Hills, at Eagle Nest, and northwest of Camel Mountain

Bell Top Formation—Tbt, Tbs, Tbt₆, Tbl, Tbai; descriptions of individual units Ph Hueco Limestone—Medium- to thick-bedded, dark- to light-gray, fossiliferous o oolitic limestone with minor shale and, at the top in the Florida Mountains, approximately 30 ft (9 m) of red shale. In the Tres Hermanas Mountains the base is marked by approximately 10-15 ft (3-4.5 m) of red-brown chert pebblecobble conglomerate with similar conglomerates recurring higher in the section; algal limestone, crinoidal debris, and gastropod faunas are typical of Hueco beds; at least 550 ft (168 m) thick in the Tres Hermanas Mountains (Kottlowski and Foster, 1962) and at least 430 ft (131 m) thick in the Florida Mountains (Clemons and Brown, 1983)

> ennsylvanian rocks, undifferentiated—Light- to medium-gray, vellowishgray, and brownish-gray fossiliferous limestone and micrite with shale partings; 60 ft (18 m) of brown, fine-grained, crossbedded sandstone is prominent below middle of unit. Approximately 560 ft (170 m) thick in the Tres Hermanas Mountains (Kottlowski and Foster, 1962); absent in the Florida Mountains; Middle Pennsylvanian (Desmoinesian-Missourian) according to Kottlowski and Foster (1962) and Balk (1961)

ssippian and Devonian rocks-Rancheria Formation (Mississippian)—In the Florida Mountains the Rancheria includes interbedded fossiliferous limestone, micrite, shale, and chert, the latter comprising as much as 50% of the upper part of the formation; approximately 220 ft (67 m) thick in the Florida Mountains (Clemons and Brown, 1983). Escabrosa Group-In the northern Tres Hermanas Mountains the Escabrosa Group consists of an upper unit. the Hachita Formation, consisting of crinoidal limestone and a lower unit the Keating Formation, which is thin- to medium-bedded, gray fossiliferous limestone with abundant chert layers and nodules. The Escabrosa is approximately 360 ft (110 m) thick in the Tres Hermanas Mountains (Kottlowski and Foster, 1962). Percha Shale (Devonian)—Black to olive-gray shale with minor thin limestone beds; approximately 250 ft (76 m) thick in the Florida Mountains

Montoya Dolomite and Fusselman Dolomite-Fusselman Dolomite (Silurian)—Medium- to thick-bedded, light- to dark-gray, fine- to coarse-grained dolomite in alternating layers: 1.400-1.500 ft (427-457 m) thick in the Florida Mountains (Clemons and Brown, 1983). Montoya Dolomite (Upper Ordovician)-Upper 180 ft (55 m) of nearly chert-free, medium- to light-gray, finegrained Cutter dolomite and limestone; medial 125-200 ft (38-61 m) of cherty, gray Aleman limestone and dolomite; and basal 50-70 ft (15-21 m) of massive. dark-gray Upham dolomite (with basal Cable Canyon Sandstone)

Bliss Sandstone and El Paso Formation-El Paso Formation (Lower Or-

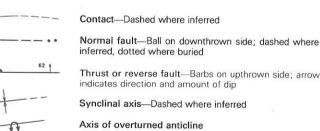
dovician)—Upper unit of thin- to medium-bedded, dark-gray, cherty limestone overlies a middle unit of thin- to medium-bedded, gray limestone and cherty limestone and a basal unit of dark-gray dolomite; 1,200 ft (365 m) thick in the Florida Mountains (Clemons and Brown, 1983). Bliss Sandstone (Cambrian-Ordovician)—Interbedded sandy limestone, black shale, and white quartzite grades down into a basal brown arkose and coarse sandstone; 50-185 ft (15-56 m) thick in the Florida Mountains (Clemons and Brown, 1983) ranite—Fine to coarse-grained, red to gray to brown granite containing, in

the southern Florida Mountains, as much as 50% hornblende and pyroxenehornfels xenoliths. U-Pb age of granite and hornfels in the Florida Mountains is 503 ± 10 Ma and 504 ± 10 Ma, respectively (Evans and Clemons, 1988) Clemons, 1995), but other published ages range from 371 Ma to 1,600 Ma, indicating uncertainty in the age of these rocks

syenite in the southern Florida Mountains containing many aplitic zones and large masses and xenoliths of hornblende and pyroxene hornfels. Pb-Pb and U-Pb ages on the syenite are 533 Ma and 460-471 Ma, respectively, but other reported ages range from 371 Ma to 1,600 Ma (Evans and Clemons,

Metamorphic rocks—Gneissic granite, gneiss, and biotite schist in the north vestern part of the Florida Mountains: unit also includes small exposures of cobble-boulder conglomerate described as diamictite by Corbitt and Wooderate of the Bliss Sandstone (Clemons, 1995)

MAP SYMBOLS



Axis of overturned syncline Strike and dip of bedding or of foliation in ash-flow

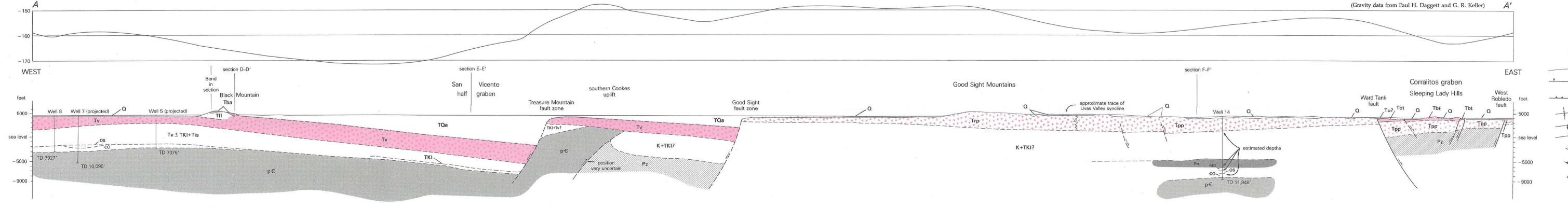
Strike and dip of foliation in flow-banded rhyolite Selected oil test hole—Number adjacent to symbol cor-

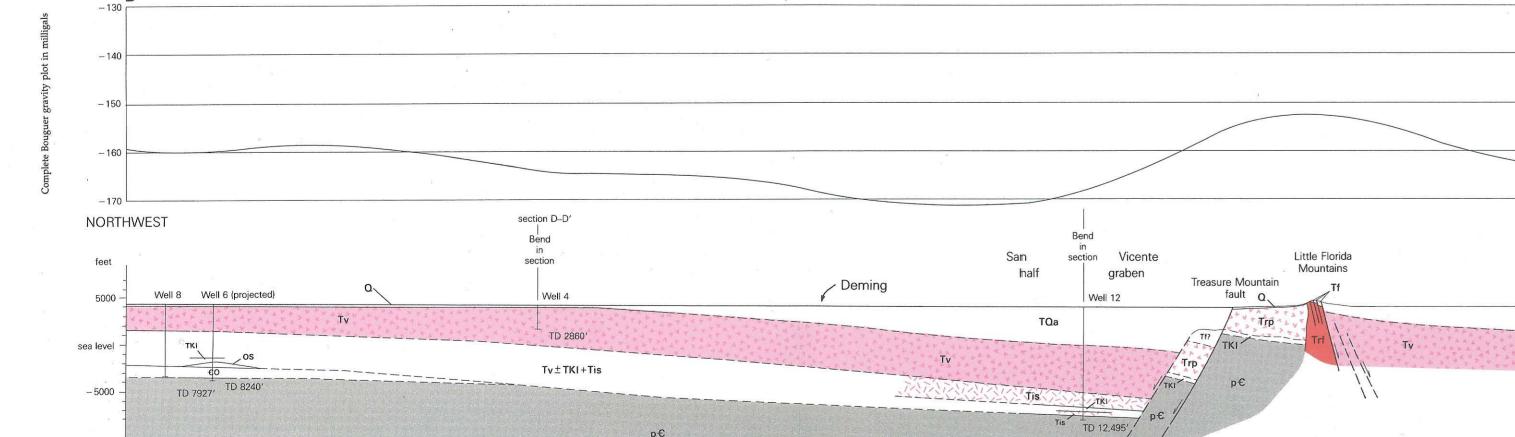
responds to well listed in Table 2 (Sheet 5); data from ottlowski et al. (1969) and Thompson (1982)

West Potrillo Mountains

Geologic sections and gravity profiles

through the southwest quarter of Las Cruces and northwest El Paso 1° × 2° sheets (scale 1:125,000), New Mexico by William R. Seager, 1995





of Camel Mountain