PREVIOUS WORK

The La Ventana quadrangle area was included in reconnaissance maps by Renick (1931) and Wood and Northrop (1946). The western part of the quadrangle was included in another reconnaissance map by Dane (1936). A map by Parry (1957) covers the northwestern part of the area. La Ventana Mesa was mapped by Bachman and others (1959) in connection with a study of uranium resources.

PRESENT WORK

Responsibility for present mapping is shown on inset map. The steeply dipping beds at the western edge of Sierra Nacimiento and the Precambrian rocks were mapped by Schumacher; the remaining part of the quadrangle was mapped by Woodward. The area around San Miguel mine was mapped by Woodward and Schumacher at 1:6,000 scale and transferred to this map.

STRATIGRAPHY

The crystalline rocks, overlain unconformably by Mississippian strata a few miles to the south (Fitzsimmons and others, 1956), are probably of Precambrian age as suggested by regional evidence (Muehlberger and others, 1967). Small inclusions of hornblende schist are present in both the gneiss and the granite. Only the largest inclusion is shown because most are too small to show at this scale. The gneiss, intruded by the granite, is the next younger unit.

The youngest Precambrian rocks are the mafic dike and the pegmatite-aplite bodies, but their age relations to each other are not known. The Madera Formation (Pennsylvanian) is the oldest Paleozoic unit. This formation thins northward and is absent about two miles north of this quadrangle where the Abo Formation (Permian) rests directly on the Precambrian. In this quadrangle the contact between Precambrian and vounger rocks is tec-

northern part of the area is not known. The Glorieta Sandstone and Bernal Formation, undivided, was not mapped north of San Miguel mine; however, some Glorieta may be present directly beneath the Agua Zarca (Triassic), with the Bernal missing. The lithologic similarity between the Agua Zarca and Glorieta makes their contact difficult to map. Thus, the Agua Zarca, as mapped north of San Miguel mine, may include some

Glorieta Sandstone at the base.

tonic; whether the Madera is tectonically or unconformably missing in the

being the Agua Zarca Member, and the upper, informally named "upper member." Where exposures are very good, the upper member can be subdivided into two units correlating with the Salitral Shale and Upper Shale Members of the Chinle to the north (Woodward and others, 1972).

Landslide deposits consist of blocks of sandstone derived from the Dakota Formation and the Mesaverde Group. Locally these deposits grade into colluvium.

STRUCTURE

Precambrian Deformation Small inclusions of hornblende schist in the gneiss and granite probably

here than in adjacent areas.

were metamorphosed prior to formation of the gneiss, but regional pattern of deformation cannot be discerned. Regional synkinematic metamorphism of the gneiss represents the oldest widespread deformation. The granite was emplaced later, possibly by dilation and stoping with minor assimilation. Pegmatite-aplite bodies and mafic dikes appear to have been emplaced by dilation.

Paleozoic Deformation

Isopach maps by Wood and Northrop (1946) indicate that the Nacimiento area was a positive structural element during Pennsylvanian time; and continued to show positive tendencies during Permian time, as evidenced by thinner strata

Laramide Deformation

The major structural features are Laramide in age. They consist of the Nacimiento uplift in the eastern part of the quadrangle and the San Juan basin to the west, separated by a belt of steeply dipping and complexly faulted beds at the western foot of Sierra Nacimiento (structure sections A-A' and B-B'). The Pajarito fault is the principal range-marginal fault and appears to be nearly vertical. The structural relief between the highest part of the uplift and the lowest part of the basin is at least 9,200 feet.

ment of the uplift and the range-marginal fault system. Evidence of right-shift between the uplift and basin during their early development is seen in several northwest-plunging en echelon folds along the eastern margin of the San Juan basin (Kelley, 1955; Baltz, 1967). In this quad-

Several small east-west trending faults appear to be related to develop-

tana Mesa; and a broad, open anticline having the same trend passes through the mesa northwest of the abandoned town of La Ventana. High-angle faults in the northwest part of the quadrangle trend north-northeast and have small stratigraphic separations. These faults may be tensional features related to the episode of right-shift noted above. The right-shift is early Laramide followed by the principal movement which was vertical.

ECONOMIC GEOLOGY

The San Miguel mine was worked extensively between 1884 and about 1900, producing about 5,000,000 pounds of copper (Lindgren and others, 1910, p. 147). Since 1900 the mine has been operated sporadically (Elston, 1967). Mineralization occurs in the Agua Zarca Member of the Chinle Formation. The ore occurs as chalcocite replacing, or in close association with, carbonaceous fossil plant material. Other minerals are malachite, azurite, and chrysocolla (?) disseminated interstitially in the clastic host rock. The ore reportedly carried 2½ to 3 oz/ton silver (Lindgren and others, 1910). Several other prospect pits (shown on map) in the Madera, Abo, and Agua

Zarca have minor copper mineralization. Uranium in coal, carbonaceous shale, and carbonaceous sandstone at the base of the La Ventana Tongue of the Mesaverde Group has been reported by Bachman and others (1959, p. 295) who estimated that 132,000 tons of coal

and carbonaceous shale containing an average of 0.10 percent uranium are present near La Ventana Mesa. The mineralized zone occurs in three lenticular beds and is several feet thick. Uranium content ranges up to 1.34 percent. A uranium prospect in the base of the Dakota Formation occurs near structure section A-A'. Lenticular beds of subbituminous coal up to six feet thick are present in the Menefee Formation of the Mesaverde Group. Detailed descriptions of the coal occurrences are given by Dane (1936, p. 142-150). At present coal is not

mined commercially in this quadrangle; locations of former mines are shown on

Gypsum is found in the Todilto Formation (Jurassic), but steep dip and thickness of overburden prevent large-scale open-pit mining.

facing and pit-run sub-base.

Tertiary-Quaternary terrace and pediment deposits are an excellent source of aggregate. The clasts consist mostly of Precambrian crystalline rocks and range up to 2 feet in diameter. Thickness of these deposits is mostly 2 to 10 feet, but locally may be up to 30 feet. These deposits have been used for road sur-

REFERENCES

Bachman, G. O., Vine, J. D., Read, C. B., and Moore, G. W., 1959, Uraniumbearing coal and carbonaceous shale in the La Ventana Mesa area, Sandoval County, New Mexico: U. S. Geol. Survey Bull. 1055-J, p. 295-307. Baltz, E. H., 1967, Stratigraphy and regional tectonic implications of part of Upper Cretaceous and Tertiary rocks, east-central San Juan basin, New Mex-

Dane, C. H., 1936, Geology and fuel resources of the southern part of the San Juan basin, New Mexico, in Part 3 – The La Ventana-Chacra Mesa coal field: U. S. Geol. Survey Bull. 860-C, p. 81-166. Elston, W. E., 1967, Summary of the mineral resources of Bernalillo, Sandoval,

ico: U. S. Geol. Survey Prof. Paper 552, 101 p.

Mexico: U. S. Geol. Survey Prof. Paper 68, 361 p.

and Santa Fe counties, New Mexico: New Mexico State Bur. Mines Mineral Resources Bull. 81, 81 p. Fitzsimmons, J. P., Armstrong, A. K., and Gordon, Mackenzie, Jr., 1956, Arroyo Peñasco Formation, Mississippian, north-central New Mexico: Amer. Assoc.

Petroleum Geologists Bull., v. 40, p. 1935-1944. Kelley, V. C., 1955, Regional tectonics of the Colorado Plateau and relationship to the origin and distribution of uranium: Univ. New Mexico Pub. in Geology, Lindgren, W., Graton, L. C., and Gordon, C. H., 1910, The ore deposits of New

Muehlberger, W. R., Denison, R. E., and Lidiak, E. G., 1967, Basement rocks in continental interior of United States: Amer. Assoc, Petroleum Geologists Bull., v. 51, p. 2351-2380. Parry, M. E., 1957, A sandstone channel in the Mesa Verde Group near Cuba,

New Mexico: Univ. New Mexico, unpub. M. S. thesis, 39 p. Renick, B. C., 1931, Geology and ground-water resources of western Sandoval County, New Mexico: U. S. Geol. Survey Water Supply Paper 620, 117 p. Wood, G. H., and Northrop, S. A., 1946, Geology of the Nacimiento Mountains, San Pedro Mountain, and adjacent plateaus in parts of Sandoval and Rio Arriba counties, New Mexico: U. S. Geol. Survey Oil and Gas Inv. Prelim.

Woodward, L. A., Kaufman, W. H., Anderson, J. B., and Reed, R. K., 1972, Geologic map and sections of the San Pablo quadrangle, New Mexico: New Mexico State Bur. Mines Mineral Resources Geol. Map 26.



