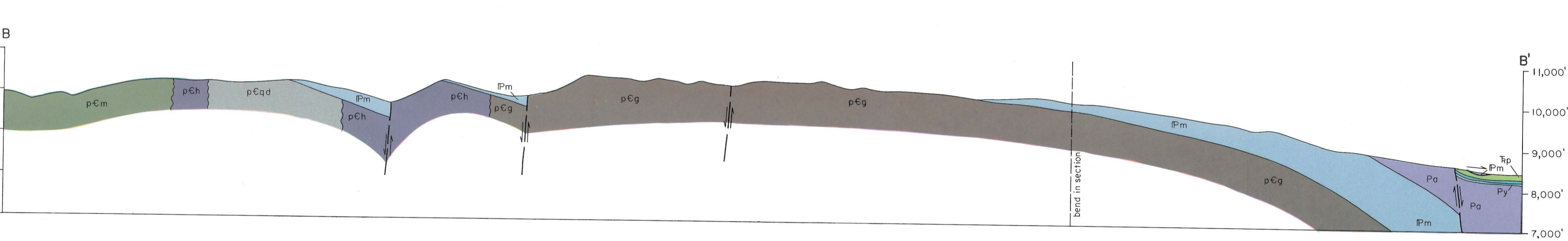
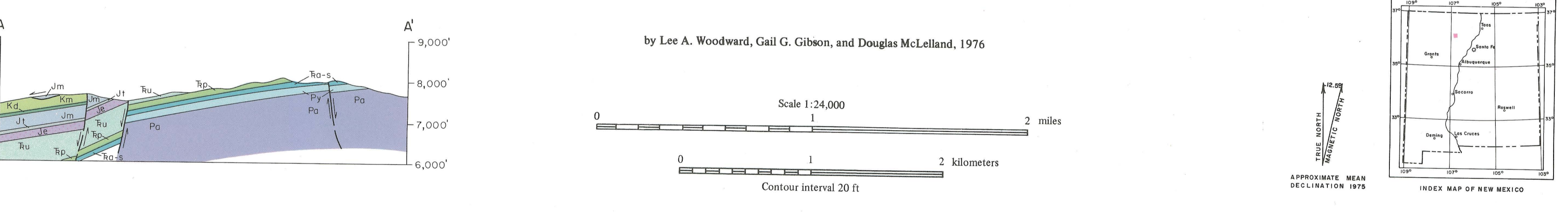


Geology by Lee A. Woodward, Gail G. Gibson, and Douglas McLelland, 1972, 1973, 1974, 1975



GEOLOGY OF GALLINA QUADRANGLE, NEW MEXICO



PREVIOUS WORK

A reconnaissance map of the Nacimiento Mountains and adjacent areas by Wood and Northrop (1946) included this quadrangle. The western part of the quadrangle was mapped by Hutson (1958) and the southeastern part by Acosta (1973) for master's theses.

ROCK UNITS

There are no radiometric ages reported for Precambrian rocks in this quadrangle. However, Brookins (1974) reported ages of 1,800 ± 50 m.y. and 1,840 ± 170 m.y. for metamorphic rocks in adjacent quadrangles to the south and to the west.

The suggested relative ages of the Precambrian rocks are shown in the map explanation. Relative ages of the granite bodies are not known, but it seems likely that they are contemporaneous and that the fine-grained granite is a facies of the more extensive muscovite-biotite granite.

The Madera Formation (Pennsylvanian) rests nonconformably on Precambrian rocks and may locally include beds of Mississippian age at the base (Fitzsimmons and others, 1956). We were not able to find any diagnostic fossils nor any lithologies similar to nearby known Mississippian rocks, and therefore included these strata resting on the Precambrian with the Madera Formation.

The upper contact of the Madera with the Abo Formation was placed at the top of the stratigraphically highest, thick bed of limestone containing marine fossils of Pennsylvanian age. There is a lateral intertonguing between the Madera

and the Abo insofar as the uppermost limestone beds of the Madera pinch out toward the southeast. Thus, the Madera-Abo contact is stratigraphically lower in the southeast part of the quadrangle.

The Yeso Formation (Permian) sits eastward and its presence along Cecilia Canyon is inferred from float.

Strata in the lower part of the Chinle Formation were mapped as Agua Zarca Sandstone Member-Salitral Shale member undivided because this interval does not consist of basal whitish sandstone (Agua Zarca) overlain by maroon shale (Salitral). Rather, the interval is comprised of intercalated shale and sandstone with laterally discontinuous conglomeratic sandstone lenses at various horizons. The Agua Zarca-Salitral is missing in the central part of the quadrangle where the Poleo Sandstone member lies directly upon the Yeso Formation.

The Morrison Formation thickens from 300 ft near Gallina Plaza to about 900 ft at Capulin Peak to the northeast.

STRUCTURE

The Gallina quadrangle is located at the junction of four major tectonic features: the Nacimiento uplift, the San Juan Basin, the Gallina-Archuleta arch, and the Chama basin (fig. 1). The northern end of the north-trending Nacimiento uplift is separated from the San Juan Basin by a monocline that is broken by a zone of steeply dipping faults that trend northeast near Gallina Plaza. There is about 8,600 ft of structural relief between the Nacimiento uplift and the San Juan Basin in this quadrangle. A low saddle separates the Nacimiento uplift from the structurally lower Gallina-Archuleta arch to the north.

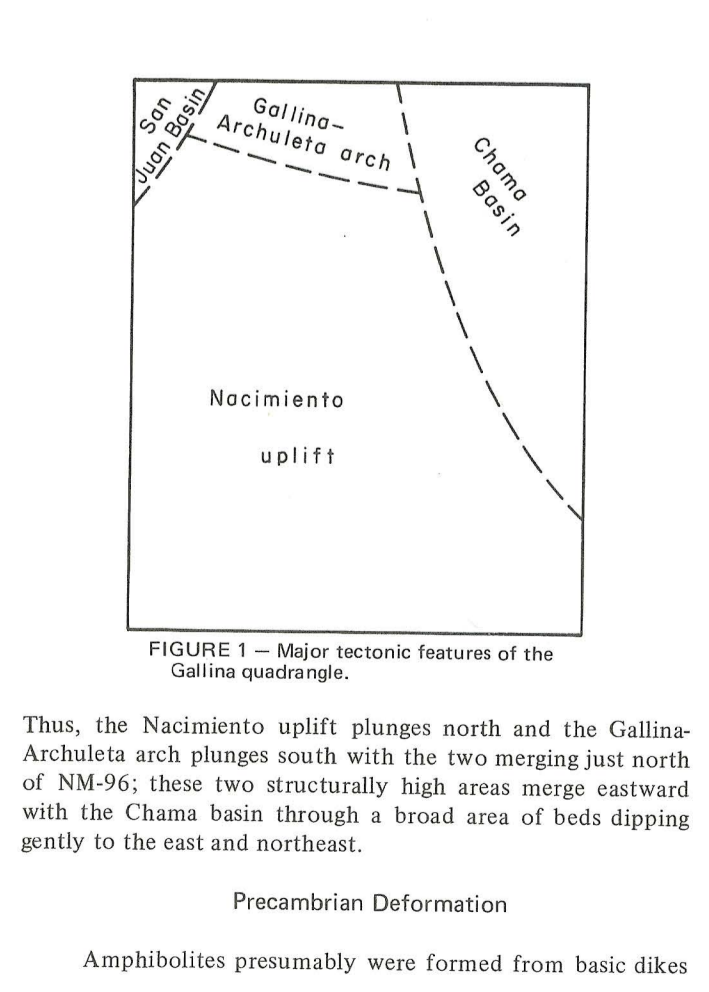


FIGURE 1—Major tectonic features of the Gallina quadrangle.

placed into a sedimentary and volcanic sequence. These rocks were later regionally and synkinematically metamorphosed in the upper greenschist and lower amphibolite(?) facies to form pCm and pCa.

The mechanisms of emplacement of the quartz diorite (pCqd) and the granitic rocks (pCg, pCgI, and pCgB) are not known.

Mesozoic Deformation

Absence of the lower members of the Chinle Formation (Triassic) along the present axis of the Nacimiento uplift (sec. 22, T. 23 N., R. 1 E.) suggests that the uplift was relatively positive during that part of Late Triassic time.

The rapid eastward thickening of the Morrison Formation (Jurassic) indicates differential subsidence, with the western part of the Nacimiento uplift undergoing less subsidence than to the west or east (Gibson, 1975).

Cenozoic Deformation

Development of the major tectonic features in the Gallina quadrangle began in early Cenozoic time. The Nacimiento uplift continued to rise at least into late Cenozoic time, as seen by tilting of the Pedernal Chert Member of the Abiquiu formation (Church and Hack, 1939) of probable Miocene age (Smith, 1938).

Tilted fault-blocks in the southern part of the quadrangle (section B-B') probably formed in response to stretching of the crest of the uplift as it was free to expand laterally over the San Juan Basin along a range-marginal upthrust on the west

side of the uplift (Woodward and others, 1972). The high-angle fault along Cecilia Canyon cuts a monocline marking the northeast margin of the Nacimiento uplift (section B-B').

Small, gravity-slide plates of Poleo Sandstone (secs. 23 and 24, T. 23 N., R. 1 E.), Madera limestone (sec. 8, T. 22 N., R. 2 E.), and Morrison Formation (secs. 8 and 18, T. 23 N., R. 1 E.) have the age relations of overthrusts, but appear to have formed by downslope gravitational movement after development of the major tectonic features (Nacimiento uplift, San Juan Basin, and Chama basin).

ECONOMIC GEOLOGY

Sandstone-type copper mineralization is present at several prospects in the Abo Formation in sec. 31, T. 23 N., R. 1 E.). These deposits do not appear to be economic due to small reserves. The genesis of these deposits has been discussed by Woodward and others (1974).

Gypsum in the Todilto Formation is exposed over about 300 acres and has a thickness of 90 ft at Cerro Blanco (secs. 4, 5, 8, and 9, T. 23 N., R. 1 E.). Distance from a market and low unit value make the gypsum uneconomic at this time.

Terrace and pediment deposits (QTlp) have been used extensively for aggregate. The clasts consist mostly of pebbles, cobbles, and boulders of Precambrian crystalline rocks.

Several uranium occurrences in the Abo Formation south of Gallina were noted by Chenoweth (1974), but we did not observe any uranium mineralization during our mapping.

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