# Geologic Map of the Fort Stanton Quadrangle, Lincoln County, New Mexico.

By

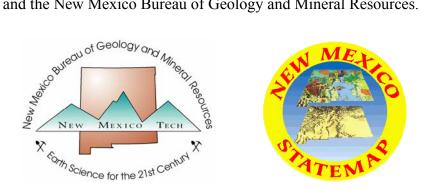
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### New Mexico Bureau of Geology and Mineral Resources Open-file Digital Geologic Map OF-GM 119

### Scale 1:24,000

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## Geologic Map of the Fort Stanton Quadrangle

### OF-GM-118

Mapped and compiled by Geoffrey Rawling

#### **EXPLANATION OF MAP UNITS**

#### **Anthropogenic Deposits**

af Artificial fill for stock tanks and highway embankments.

**daf** Heavily disturbed land and artificial fill. Mapped where extensive, underlying deposits are obscured, and/or geomorphic surfaces are extensively altered.

#### **Quaternary and Tertiary Surficial Deposits**

- **QHa Alluvium (Holocene to Historic)** Unlithified gravel and poorly to moderately sorted clay, silt, sand in active stream channels and ephemeral arroyos. Generally incised into **Qvf** and terrace deposits. Only mapped where extensive; unit is otherwise lumped with **Qvf**. Thickness: 0 to 4 (?) meters.
- **Qvf Valley fill (upper Pleistocene to Holocene)** Unlithified valley fill composed of poorly sorted clay, silt, and sand, commonly with angular to subrounded cobbles of local bedrock. Matrix material is light to dark brown. Grades into minor alluvial and

colluvial fans on toes of hillslopes. Anthropogenic disturbance common in developed areas. Generally incised by active drainages, floored by sand and cobble to boulder gravel of **QHa**. Thickness: 0 to 12 (?) meters

- Qaf Alluvial fan deposits (middle to upper Pleistocene) Alluvial fans composed of poorly sorted cobbles, boulders, sand, silt, and clay. Fans head in short, steep tributary canyons and interfinger with and/or spread out onto Qvf and terrace deposits. Stabilized by vegetation and apparently no longer active, and locally incised by drainages floored with QHa. Only mapped along major drainages where geomorphic expression is clear on aerial photos. Thickness: 0 to 8 (?) meters.
- **Rio Bonito terrace deposits** The Rio Bonito crosses a prominent diorite sill just west of the Fort Stanton Recreation Area boundary, forming a waterfall. Terraces with similar relative geomorphic positions upstream and downstream of this knickpoint may not be directly correlated, although they are identified as such on the map.
- **Qbt4 Lowest terrace deposit of Rio Bonito** (**Holocene**) Poorly to well-sorted alluvial deposits composed of interstratified fine to coarse tan sand and sandy cobble to boulder gravel. Deposit is approximately at grade of the Rio Bonito and generally only present within small meander bends of the active stream channel. Largely mapped from aerial photographs. Thickness: 0 to 2 (?) meters.
- **Qbt3 Intermediate terrace deposit of Rio Bonito (Uppermost Pleistocene to Holocene)** – Poorly to well-sorted alluvial deposits composed of interstratified fine

to coarse tan sand and sandy cobble to boulder gravel. Deposit is similar to **Qbt4** in geomorphic appearance and vegetative cover but has a terrace tread above stream grade. Largely mapped from aerial photographs. Thickness: 0 to 3 (?) meters.

- **Qbt2 Intermediate terrace deposit of Rio Bonito (Upper Pleistocene)** Poorly to well-sorted alluvial deposits composed of interstratified fine to coarse tan sand and sandy cobble to boulder gravel. Deposit forms a terrace whose tread is generally preserved and is 3 4 meters above present stream grade. Downstream (east) of the diorite sill which crosses the Rio Bonito west of the Fort Stanton Recreation Area the terrace is in part a strath terrace with discontinuous bedrock exposures along the terrace riser. Largely mapped from aerial photographs. Thickness: 0 to 4 meters.
- **Qbt1 Highest terrace deposit of Rio Bonito (Middle Pleistocene) -** Poorly to wellsorted alluvial deposits composed of interstratified fine to coarse tan sand and sandy gravel of rounded cobbles and boulders. Deposit forms a terrace whose tread is generally dissected and is 5 – 10 meters above present stream grade. Downstream (east) of the diorite sill which crosses the Rio Bonito west of the Fort Stanton Recreation Area the terrace is in part a strath terrace with discontinuous bedrock exposures along the terrace riser. Grades into and/or is locally overlapped by hillslope colluvium. Largely mapped from aerial photographs. Thickness: 0 to 6 (?) meters.

**Eagle Creek terrace deposits** 

- Qet2 Lower terrace deposit of Eagle Creek (Holocene) Poorly to well-sorted alluvial deposits composed of interstratified fine to coarse tan sand and sandy cobble to boulder gravel. Deposit forms a terrace whose tread is generally preserved and is within a meter of present stream grade. Becomes indistinguishable from Qet2 approximately 2 miles east of the western edge of the quadrangle; stream deposits in the Eagle Creek drainage are then mapped as Qvf. Largely mapped from aerial photographs. Equivalent to unit Qet2 of the Angus quadrangle (Rawling, 2004a). Thickness: 0 to 3 meters.
- Qet1 Upper terrace deposit of Eagle Creek (upper Pleistocene) Poorly sorted alluvial deposits composed of interstratified fine to coarse sand and sandy gravel of rounded cobbles and boulders. Deposit forms a terrace whose tread is preserved and is 3 – 4 meters above present stream grade. Becomes indistinguishable from Qet1 approximately 2 miles east of the western edge of the quadrangle; stream deposits in the eagle Creek drainage are then mapped as Qvf. Largely mapped from aerial photographs. Equivalent to unit Qet1 of the Angus quadrangle (Rawling, 2004a). Thickness: 0 to 5 (?) meters.

#### **Rio Ruidoso Terrace deposits**

**Qrt2 – Lower terrace deposit of Rio Ruidoso (upper Pleistocene)** - Poorly to wellsorted alluvial deposits composed of interstratified fine to coarse sand and sandy cobble to boulder gravel. Clasts are rounded intrusive and volcanic igneous rocks with lesser limestone and sandstone. Deposit forms a terrace whose tread is 5 - 6 meters above present stream grade, dissected, and generally affected by human disturbance. Mapped from aerial photographs. Equivalent to unit Qrt2 of the Ruidoso Downs quadrangle (Rawling, 2004c). Thickness: 0 to 6 (?) meters.

- **Qrt1 Upper terrace deposit of Rio Ruidoso (middle Pleistocene)** Poorly to moderately sorted alluvial deposits composed of interstratified fine to coarse sand and sandy cobble to boulder gravel. Clasts are rounded intrusive and volcanic igneous rocks with lesser limestone and sandstone. The surface of the deposit forms a terrace tread 12-15 meters above present stream grade, which is variably eroded, and strongly affected by human disturbance. Mapped from aerial photographs. Equivalent to unit Qrt2 of the Ruidoso Downs quadrangle (Rawling, 2004c). Thickness: 0 to 10 (?) meters.
- QTg Pediment gravel deposits (Pliocene (?) to lower Pleistocene) Moderately lithified, crudely bedded pebble to boulder gravel with reddish clayey sand matrix and local lenses of sand and sandy clay. Largest boulders are 80 cm in diameter. Clasts are dominantly Sierra Blanca volcanic rocks and associated intrusive igneous rocks, although angular clasts of Cretaceous Dakota Sandstone and Mesa Verde Group (Gallup?) sandstone become increasingly common near the base of the deposit in the eastern half of the quadrangle. Clasts on surface of deposit are heavily weathered and fractured. Surface is partly stripped. Contains stage III to III+ carbonate soil near surface and local strong carbonate buildup (equivalent to stage IV carbonate soil development) at base of unit. Extensive downslope colluvium makes

base of unit hard to define where not exposed in roadcuts. Deposited by steep gradient streams draining the Sierra Blanca, predating incision of modern drainages. Thickness: 0 to 30 (?) meters.

#### **Cenozoic Igneous Rocks**

- Tad Andesite/diorite and associated rocks, undivided (Oligocene) Aphanitic to very fine-grained phaneritic and phaneritic-porphyritic igneous rocks forming dikes, sills, and irregular intrusive masses. Generally dark gray on fresh surface and brown to black on weathered surfaces. When phaneritic, often has a "salt and pepper" appearance due to fine-grained equigranular white feldspar and black to brown augite (?). Phenocrysts include augite, hornblende, and tabular intermediate (?) plagioclase. Tabular plagioclase phenocrysts are up to 4 cm in diameter and are usually aligned with the dike margins. Thickness: dikes are < 1 to 5 meters wide.</p>
- Tml Monzonite/latite and associated rocks (Oligocene) Tan to brown aphanitic to very fine-grained phaneritic igneous rocks forming dikes, sills and irregular intrusive masses. Also forms a small laccolith in the northeast corner of the quadrangle. Typically composed of approximately equal amounts of white feldspar and tan to brown mafic minerals with little or no quartz. Feldspar is largely intermediate (?) plagioclase with lesser amounts of potassium feldspar and forms a felted network of interlocking crystals. Includes rocks ranging from syenite to diorite in composition. Thickness: dikes are < 1 to 5 meters wide; base and top of larger igneous masses and the laccolith are not exposed.</li>

Tcd – Diorite of Champ Hill (Oligocene) - Dark gray to grayish-brown fine-grained phaneritic diorite with 10-35% phenocrysts of augite up to several centimeters across. Named for Champ Hill, a stock which straddles the northeastern border of the quadrangle with the Angus quadrangle to the west (Rawling, 2004a). Rocks of similar appearance but varying phenocrysts content form numerous dikes, sills, and irregular intrusive masses, including the prominent northeast –striking sill which crosses the Rio Bonito. Thickness: dikes are < 1 to 5 meters; base and top of larger igneous masses are not exposed.

#### **Mesozoic Sedimentary Rocks**

Kmv - Mesa Verde Group, undivided (upper Cretaceous) – Pale lavender to tan sandstone, olive to gray sandy siltstone and siltstone, and dark gray, grayish purple, olive, and minor black shale. Sandstones are fine- to medium-grained, medium- to very thick-bedded, trough cross-bedded, and composed of subrounded to subangular grains. They are dominantly quartose and resistant, but locally arkosic and friable. Sandstones of the lower Mesaverde Group forming prominent cliffs in the northwestern corner of the quadrangle have been tentatively correlated with the Gallup Sandstone (Bodine, 1956; Arkell, 1983). Black shales are carbonaceous and fissile. Shaly sections are ubiquitously intruded by igneous dikes, sills and irregular masses. Adjacent to these bodies shales are often contact metamorphosed to low grade hornfelses and weather into gray angular chips rather than black flakes. Thickness: > 270 meters, top not exposed in quadrangle. Up to 425 meters thick in

the Angus and Ruidoso quadrangles to the west and southwest (Rawling, 2004a; Rawling, 2004b).

- Km Mancos Shale (middle to upper Cretaceous) Black to purplish gray laminated fissile shale. Black to dark gray to olive thin-bedded fine-grained sandstone and siltstone beds less than 0.5 meters thick are minor constituents. Thin- to mediumbedded gray limestone beds form up to 75% of the unit north of the Rio Bonito. Igneous intrusions are common. Generally only well exposed in stream cuts and manmade excavations. Thickness: ~ 335 meters.
- Kd Dakota Sandstone (lower to middle Cretaceous) Gray to tan to light purple sandstone and minor black shale. Sandstone is medium- to thick-bedded, trough to tabular cross-bedded, ripple-marked, and composed of subangular to subrounded vitreous quartz grains. Orange to rusty red liesegang bands are common on bedding planes and fracture surfaces. Sandstone is more resistant, forms more prominent outcrops, and weathers into more angular fragments than overlying Mesa Verde sandstones. Matrix-supported sandy chert pebble conglomerate is commonly present at the base of unit and as lenses throughout the unit. Thin discontinuous beds of black shale similar to the overlying Mancos Shale are sparsely distributed throughout the upper portions of the unit. Thickness: ~ 106 meters.
- **Fisr Santa Rosa Formation (upper Triassic)** Dark brownish-red fine-grained, micaceous, sandstone, siltstone, and conglomerate, and dark red, often massive and

bioturbated, mudstone. Very poorly exposed, except for outcrops on north side of Rio Bonito Valley in the vicinity of Fort Stanton. Generally only appears as dark red soils and subcrop. Reduction spots are common. Base of the unit is often marked by a gray, orange, or red medium-to thick-bedded quartzite and chert pebble conglomerate with a matrix of coarse chert-rich sand. Conglomerate is scoured into underlying Grayburg Formation. Unit is cut out by the basal Dakota unconformity in canyons on north side of Fort Stanton Mesa and locally reappears to the southwest in channels cut into the underlying Grayburg Formation. Units present are probably the Tecolotito and Los Esteros Members of the Santa Rosa Formation (Lucas et al., 2001). Thickness: 0 - 30meters.

#### **Paleozoic Sedimentary Rocks**

Pg - Grayburg Formation (upper Permian) – Gray, tan, and yellowish brown fine- to medium-grained sandstone and subordinate siltstone. Limy sandstone and dolomite beds are locally common. Intervals of thick-bedded, resistant, chert pebble to cobble conglomerate are prominent in exposures along the north side of Little Creek. Sandstones and siltstones are very thin- to thick-bedded, parallel to crossbedded, and composed of quartz. Red, generally massive, very fine sandstone, siltstone and mudstone with white reduction spots is locally present at base. This unit appears to be scoured into or deposited in karst depressions within the underlying San Andres Formation, e.g. at the northeast end of Fort Stanton Mesa and along Eagle Creek north of Gavilan Ridge. Unit is the lower member of the Artesia Group. Thickness: 90 – 152 meters.

- Psa San Andres Formation (middle to upper Permian) Light to dark gray and bluish-gray thin- to very thick-bedded limestone and dolomite. Limestones and dolomites are carbonate mudstones, wackestones, and grainstones. Freshly broken surfaces are darker gray than weathered surfaces and are commonly fetid. Silty and sandy beds common. Dark brown irregular chert nodules are sparse. Fossils are sparse and are dominantly crinoid stem fragments. Intraformational solution breccias and paleokarst features such as collapsed caves are common along faults and as isolated occurrences. They are characterized by red soil and red and yellow stained breccia fragments. Thickness: ~ 335 meters.
- Py Yeso Formation (middle Permian) Yellow to tan siltstone and fine sandstone, red to pink muddy siltstone and fine sandstone, gray to tan silty limestone and dolomite, and white to gray gypsum. Siltstone and sandstones are thin- to medium-bedded and friable. Muddy siltstones and sandstones are laminated to very thin-bedded and locally contain paleosol carbonate nodules in trains. Limestones are very thin- to thin-bedded, rarely medium- to thick-bedded. In general, they are thinner bedded than overlying basal San Andres beds. Meter scale interbedding of carbonate, siltstone, and sandstone is common. Bedding dips are chaotic due to dissolution of gypsum and (and carbonates?) and individual beds are generally not traceable laterally for more than a few 10s of meters. Natural exposures are poor except in stream cuts and very steep slopes and the upper contact is usually mantled by colluvium and/or landslides from the overlying San Andres Formation. Thickness:

Base not exposed; ~ 240 meters exposed along Rio Ruidoso Canyon in the Ruidoso Downs quadrangle (Rawling, 2004c).

#### Pyu - Yeso Formation and older rocks, undivided - cross-section only

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