



NEW MEXICO BUREAU OF GEOLOGY AND MINERAL RESOURCES A DIVISION OF NEW MEXICO INSTITUTE OF MINING AND TECHNOLOGY

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	the same and
0	Chief elements of the Grants SE quadrant Mesozoic rocks from uppermost Chinle C below Horace Mesa is comprised of land resistant base for the northern end of Las and the basal sandstones of the Dakota H ground left of center. Houses at the base
	<b>DE</b> <i>Note:</i> Soil descriptions after Birkeland (1
o	<b>Artificial fill</b> — Artificial fill. Compacted thick.
	Quaternary
Qsw Qdf	<b>Slopewash (Holocene)</b> — Sandy silt and <b>Deflationary pits (upper Pleistocene to</b> laminated silt and lesser clay with 5-15% 2 m thick.
Qc	<b>Colluvium (upper Pleistocene to Holoc</b> up to 2 m thick.
Qes	Landslides (upper Pleistocene to Holoc underlying hummocky terrain along slop Quaternary eolian sand (upper Pleistoc low dunes. Pale brown colors typical, wit Likely up to 2 m thick.
	Eolian sand overlying basalt (upper Pl Cibola ( <i>Qbc</i> ) or El Calderon ( <i>Qbh</i> ) basalt
	<b>Dune sand (upper Pleistocene to Holoc</b> development and sandsheets with mou weak soil development such as Stage I gy
Qae	<b>Alluvial and eolian sand, undivided (u</b> slopewash, and confined water flow. Var thick.
o Qah	<b>Historic alluvium, undivided (Holocen</b> historic transport. Deposits are typically Likely up to 1 m thick.
Qajh Qajhf Qajy Qajo	Historic alluvium associated with the I clays found along the current floodplain moderately sorted angular to rounded si Historic fan alluvium with the Rio San J from the termini of gullies and channels Younger alluvium associated with the treads up to 2 m above the active channel Stage I Bk development. Colors of 10YR S than 3 m thick. Older alluvium associated with the Rio the Rio San Jose at relatively high elevat granitic and siliceous pebbles that do not
	I as Va
00 Qaf Qafh Qafy Qafo	Alluvial fan material from Mezosoic-co silt derived from small drainages in the fi inferred from geomorphology and soil de Youngest alluvial fan material from Mezo fan surfaces inset upon those of <i>Qafo</i> and to 6 m thick. Older alluvial fan material from Mezo inset upon by those of Qafy and bearing Stage II Bk horizons and colors of 10YR 6 but this was not observed in outcrop. Like
	but this was not observed in outcrop. Lik
Qay	Younger alluvium of the El Malpais vall mainly along shallow, low gradient drain with clay as fine films on ped faces and b Calcareous alluvium and eolian materi
00	pebbles with rare dark gray organic mate carbonate. Fossils are mollusk shells, and thick.
Qazo	<b>Older alluvial fan material from Zuni C</b> from the Zuni Mountain interior, specific arkosic sandstone. Distribution suggests but probably after eruption of <i>Qbc</i> . Thick
	Qua
Qpy	Alluvium of the Zuni Mountains piedmo and soil development. Units are extended <b>Younger alluvium of the Zuni Mountai</b> $Qpy_{1\nu} Qpy_{2\nu}$ and $Qpy_3$ of Timmons and Ci Likely up to 3 m thick.
Qpy <sub>3</sub>	roungest alluvium of the Zuni Moun development. Contains sparse very fine 7.5YR 5/3 measured. 0-1 m thick.
Qpo	boulders and silty sands with strong soil of 5YR 5/4 to 8.5/2 measured, with color

Qbm	<b>McCartys flow (Hol</b> porphyritic, from <1% as subhedral, clear, <
Qbb	Bandera flow alluvit Consistent age estima
Qbz	Zuni Canyon flow (u translucent greenish reddish iddingsite(?).
Qbh	Hoya de Cibola flow eolian material and v
Qbc	El Calderon flow (up mm across of mainly olivine. Basalt surface breaks in slope, areas in results, from 34 to
Qbg	<b>Grants flow (upper F</b> by <i>Qbc</i> . To date, only Here, the contact is y

olocene) - Dark gray basalt flow largely unconcealed by eolian material or vegetation. Variably % to 20% phenocrysts, principally of subhedral, <1 mm across pyroxene, but also rare (<3%) plagioclase Likely 0 to 4 m thick of erosion, and areas with relatively high microrelief. Age of eruption uncertain due to a wide spread 130 ka (Table 1). Likely 0 to 7 m thick. Pleistocene) — Dark gray to black generally phenocryst-poor basalt underlying and largely concealed v conclusively distinguished as a separate flow by the paleomagnetic work of Cascadden *et al.* (1997). placed on a well-defined flow break along the Rio San Jose, and an inferred flow break along the western margin of the El Malpais valley; only the former contact is verified by paleomagnetic work. Rock mass is up to 4%

translucent greenish). Likely 0 to 7 m thick.

1 MILE

1 KILOMETER

associated with recent development may not be shown.

wells, buildings, roads, or other man-made structures.

mapping, and available geophysical (regional gravity and aeromagnetic surveys), and subsurface (drillhole) data. Cross-sections should be used as an aid to understanding the general geologic framework

Solid where exposed, dashed where approximately known, dotted where concealed. Tick gives attitude of fault plane, showing dip magnitude, arrow gives trend of fault plane lineations, showing plunge magnitude. Exposed monocline, showing direction of plunge. Anticlinal bend of a monocline, Solid where exposed, dashed where approximately known, dotted where concealed, queried where existence uncertain. 

Geologic contact. Solid where exposed or known, dashed

where approximately known, dotted where concealed; queried

Normal fault, ball-and-bar on downthrown side, shear arrows

MAP SYMBOLS

• 73 71 Normal tault, ball-and-ball on downlate interpretation suggests.

**A Location** of geologic cross section.

where uncertain.

Horizontal bedding.

\_\_\_\_\_

\_\_\_\_\_

where approximately known, dotted where concealed, queried where existence uncertain. Basaltic dike, unit Tib. Strike and dip of inclined bedding. Strike and dip of inclined bedding, remotely determined.

# **COMMENTS TO MAP USERS**

A geologic map displays information on the distribution, nature, orientation, and age elationships of rock and deposits and the occurrence of structural features. Geologic and fault contacts are irregular surfaces that form boundaries between different types or ages of units. Data depicted on this geologic quadrangle map are based on reconnaissance field geologic mapping, compilation of published and unpublished work, and photogeologic nterpretation. Locations of contacts are not surveyed, but are plotted by interpretation of the position of a given contact onto a topographic base map; therefore, the accuracy of ontact locations depends on the scale of mapping and the interpretation of the geologist(s). Any enlargement of this map could cause misunderstanding in the detail of mapping and may result in erroneous interpretations. Site-specific conditions should be verified by detailed surface mapping or subsurface exploration. Topographic and cultural changes The map has not been reviewed according to New Mexico Bureau of Geology and Mineral Resources standards. Revision of the map is likely because of the on-going nature of work in the region. The contents of the report and map should not be considered final and complete until reviewed and published by the New Mexico Bureau of Mines and Aineral Resources. The views and conclusions contained in this document are those of the authors and should not be interpreted as necessarily representing the official policies, either expressed or implied, of the State of New Mexico, or the U.S. Government. Crosssections are constructed based upon the interpretations of the authors made from geologic

of the map area, and not be the sole source of information for use in locating or designing



<1 mm across prismatic crystals. Age estimates range from 2.4 to 3.9 ka (Table 1). Likely 0 to 6 m thick. **ium (lower Holocene)** — Lithologically similar to *Qbm*, though with more fine-grained eolian cover. nates range from 9.17 to 12.5 ka (Table 1). Likely 0 to 6 m thick. upper Pleistocene) — Dark gray to black, generally phenocryst-poor basalt. Up to 3% phenocrysts of pyroxene and olivine, both <1/2 mm across, typically anhedral, and variably degraded to translucent (upper Pleistocene) — Basalt flows burying *Qbc*, and buried by *Qbm* and *Qbb*, locally concealed by vegetation though typically only partially concealed. Likely 0 to 5 m thick. pper Pleistocene) — Dark gray to black basalt with rare fine phenocrysts. Up to 5% phenocrysts <1/2  $\bar{y}$  subhedral translucent greenish pyroxene and lesser (up to 1% of rock) anhedral translucent greenish e is often largely concealed by eolian material and vegetation, with actual basalt cropping out mainly at

phenocrysts of pyroxene (up to 1 mm across, subhedral, translucent greenish) and <1-2% olivine (euhedral, <1 mm across,

ains piedmont, undivided (upper Pleistocene to Holocene) - Combined unit of Cikoski (2012), mapped where finer subdivision is not readily possible at this scale. ntains piedmont (Holocene) – Brown silty fine to coarse sand with little soil e carbonate nodules (Stage I carbonate horizon morphology). Colors of 5YR 6/4 to ns piedmont (upper Pleistocene) – Pale to medium reddish brown pebbles to development. Stage III carbonate horizon morphology common in outcrop. Colors controlled by abundance of carbonate. Likely up to 10 m or more thick. Quaternary and Tertiary Lava Flows and Intrusive Rocks

ically fine pebbles of granite, red siltstone, quartz sandstone, siliceous material, and s transport from Zuni Canyon into the El Malpais valley prior to the eruption of *Qbz* kness unknown, possibly tens of meters in the El Malpais graben. aternary Piedmont Alluvial Deposits Zuni Mountains nont, undivided. Typically subdivided based on age inferred from geomorphology ed from work to the west by Timmons and Cikoski (2012).

terial accumulations. Soil analyzes by White (1989) indicate deposits are up to 100% nd carbonate casts of grasses and roots are also locally common. Likely up to 3 m Canyon (Pleistocene) — Gravel, sand, and silt bearing clasts indicative of derivation

6/3-6/4 and 2.5Y 7/4. Evidence for a Stage III K horizon at the top of these deposits, ikely up to 10 m thick. El Malpais valley alluvium lley (upper Pleistocene to Holocene) — Fine sand and mud with local fine pebbles, inages. Active soils and buried soils are typically weak, A/Bw to A/Bt/Bk (Stage I), bridging grains. Likely up to 2 m thick. erial (upper Pleistocene to Holocene) - Fossiliferous sand, mud, and sparse fine

flanks of Las Ventanas Ridge and Horace Mesa. Typically subdivided based on age **esozoic-cored highlands (middle to upper Holocene)** — Likely up to 1 m thick. zosoic-cored highlands (upper Pleistocene to Holocene) - Alluvium underlying d bearing weak soils (A/Bw at most), with colors of 10YR 6/3 measured. Likely up cosoic-cored highlands (upper Pleistocene) - Alluvium underlying fan surfaces buried soils and evidence of a strong active soil. Buried soils have Bt and up to

**o** San Jose (upper Pleistocene) — Thin deposits of sand and gravel associated with ations along the active channel. Largely only identifiable by the presence of sparse ot occur in underlying strata. No signs of soil development observed. 0-1 m thick. entanas Ridge and Horace Mesa alluvium cored highlands, undivided (upper Pleistocene to Holocene) – Pebbly sand and

n of the Rio San Jose. No sign of soil development. Deposits are typically poorly to silt/sand grains and fine pebbles. Likely up to 1 m thick. Jose floodplain (Holocene) — Historic alluvium associated with thin alluvial fans s along the Rio San Jose floodplain. Likely up to 1 m thick. e Rio San Jose (upper Pleistocene to Holocene) – Alluvium underlying terrace el, with typical soil development including a weak reddened Bw horizon and up to 3.5/4 to 5/6, 7.5YR 5/3 to 5/4, and locally 5YR 4/4 and 7/4, measured. Likely not more

ene) – Boulders to fine sand along active stream channels exhibiting evidence of *y* poorly to moderately sorted angular to rounded silt/sand grains and fine pebbles. Rio San Jose alluvium Rio San Jose, undivided (Holocene) – Sand and silt with lesser gravel and rare

**Quaternary Alluvial Deposits** 

**cene)** — Clean to silty fine sand and fine sandy silt forming dune forms of variable nded microrelief. Variably vegetated, with the more vegetated dunes displaying ypsum accumulation. As much as 5 m thick. upper Pleistocene to Holocene) – Silty sand and sandy silt transported by wind, riably vegetated, typically does not exhibit signs of soil development. <1/2 to 2 m

opes below basalt-capped mesas. cene to Holocene) — Silty fine sand and fine sandy silt forming sand sheets and vith colors of 7.5YR 7/3 to 7/6 and N 9 measured. Little evidence of soil development. **Pleistocene to Holocene)** — Hybrid unit that locates areas of either the Hoya de t flows largely buried by thin (0-1 m thick) eolian sand cover.

Holocene) — Gray to brown (7.5YR 7/2 and 5YR 4/2 measured), generally well-6 fine sand, subangular to subrounded, accumulating in depressions. Likely up to **cene**) — Gravel, sand, and silt transported by mass wasting and sheet flow. Likely cene) — Slumps, slides, debris-flow deposits, and associated fine-grained material

ed gravel, sand, and mud underlying roads, railroads, and water tanks. Up to 5 m y Colluvium, Slopewash, and Eolian Deposits l silty sand accumulating in depressions in landslide areas. Likely up to 2 m thick.

Anthropogenic Deposits

**ESCRIPTION OF MAP UNITS** 1999), colors after Munsell Color (1994), and all proportions are visual estimates. **Cenozoic Erathem** 

Group sediments through the Crevasse Canyon Formation. The hummocky terrain dslides. In the foreground, pale yellowish sandstones of the Bluff Sandstone form a s Ventanas Ridge, and are capped by poorly exposed Morrison Formation sediments Formation at the top. The Holocene McCartys lava flow is apparent in the middle e of Horace Mesa lie within the eastern outskirts of the town of Grants.

ngle. In the background, Horace Mesa is underlain by Pliocene basalt flows capping

<sup>36</sup>Cl DP04 34°55.97'N 107°50.45'W Weighted mean of 0 mm/kyr and 5 mm/kyr erosion rate L94 35°00.32'N 108°04.35'W <sup>3</sup>He L94 34°59.64'N 108°05.22'W L94 34°59.57'N 108°05.48'W DP04 34°59.55'N 108°05.30'W Weighted mean of 0 mm/kyr and 5 mm/kyr erosion rate

L94 34°56.01'N 107°50.33'W

L94 35°05.16'N 107°46.52'W

 
 Table 1. Summary of geochronologic data for the Grants SE basalt flows.
 Latitude Longitude Comment

							assumptions		
	41	7	Ar/Ar	M94			Inconsistent age, possibly from excess argo		
Qbh	50	14	Ar/Ar	LWup			Uncertain age, possibly suffers from excess		
Qbc	128	33	K-Ar	CL88					
	54	50	K-Ar	L93	35°07.52'N	107°20.55'W			
	34.7	3.0	<sup>36</sup> Cl	DP04	35°04.49'N	107°45.29'W	Weighted mean of 0 mm/kyr and 5 mm/kyr		
							assumptions		
	115-120		PM	C97			Suggested to have erupted during the Blake		
							geomagnetic polarity event		
Ttbp	3238	85	K-Ar	L93	35°05.17'N	107°42.50'W	Lower lava flow from southeastern tip of H		
<ul> <li>PM – age from comparison of radiometric ages and paleomagnetic data to established paleomagnetic sequence; <sup>1</sup>C ages date material from immediately underlying the flows; <sup>3</sup>He and <sup>36</sup>Cl are cosmogenic surface exposure ages; <sup>40</sup>Ar/<sup>39</sup>Ar and K-Ar are crystallization ages.</li> <li><sup>4</sup>: C97 – Cascadden et al., 1997; CL88 – Champion and Lanphere, 1988; DP04 – Dunbar and Phillips, 2004; L93 – Laughlin et al., 1993; L94 1994; LWup – Laughlin and WoldeGabriel, unpublished data, cited in Laughlin and WoldeGabriel, 1997; M94 – McIntosh, 1994.</li> <li>Porphyritic trachybasalt capping Horace Mesa (Pliocene) — Medium gray basalt with up to 5% plagioclase (up to 6 mm across, translucent white to chalky white, lathe-like subhedral or fractured of rock mass) and pyroxene (&lt;1-6 mm across, black to dark translucent green, anhedral to subhedr. K-Ar age of 3.24 ± 0.09 Ma from Laughlin et al. (1993) thought to be from this flow. 15-25 m thick.</li> <li>Intrusive basalt (Pliocene) — Dark gray to dark brownish gray and black pyroxene-plagioclase b Rock mass is up to 10% greenish-translucent subhedral pyroxene, 1-3% clear to white subhedral reddish translucent crystals, possibly iddingsite replacing olivine. All phenocrysts &lt;1 mm across. D</li> </ul>									
	Mesozoic Erathem								
	Late Cretaceous Rocks								
Crevasse Canyon Formation									
(including associated tongue of the Mancos Shale)									
<b>Dalton Sandstone Member</b> — White to pale brown quartz-rich, variably silty fine to medium sattypically 10-60 cm thick, with common 1-5 mm thick cross-laminae. Colors of 2.5Y 8.5/1 and 2.5Y									

quadrangle. 75-80 m thick.

**GEOLOGIC CROSS SECTION** 

Js Jt

Jb

Kcda

Kcdi

Kgc

Kgm<sub>2</sub>

Kge

Kgm<sub>1</sub>

Kmrd

Kdt

Kmw

Kdc

**Moenkopi Formation** — Slope-forming red to reddish brown mudstone along the eastern edge of the quadrangle. Poorly exposed. Less than 5 m thick.

San Andres Formation - Light to dark gray limestone and dolomite. Exposed only along the eastern margin of the

**Paleozoic Erathem** 

Permian Rocks

Morrison Formation, undivided — Interbedded pale yellow to light brown sandstone, varicolored mudstone, and lesser conglomerate. Lower sandstones are of silty fine sand, typically pale yellow in color (2.5Y 8/3 measured), and are indistinctly bedded with cross-laminae. Upper sandstones and conglomerates are pale brown to white, of quartz-rich silty fine to coarse sand, generally in thin beds (5-8 cm thick) with common, well-expressed cross-laminae. Pebbles, only observed in the upper part of the unit, are up to 1 cm across, moderately sorted, subrounded to rounded, and mainly of siliceous material with rare granites (up to 2% of beds) and sparse white to gray quartzites (<<1%). Mudstones are clay to silt, with colors of 5YR 5/3, 5GY 8/1, and 5Y 8/2 measured. Use of *Jm* here differs from the previous maps of Thaden *et al.* (1967) and Maxwell (1986) in that the strongly cross-bedded sandstones and conglomerates between the varicolored mudstones and bioturbated lower Dakota sandstones are included in the Morrison and not considered a part of the Oak Canyon Member of the Dakota. Pinches out to the south. 0-30 m thick. Bluff Sandstone – Light gray to strong olive colored, strongly cross-bedded, dominantly eolian sandstone with local fluvial sandstone. Variably silty fine to medium sands, mainly of quartz. Thin to medium beds (up to 1.5 m thick observed), tabular and wedge-shaped, with steep, large-scale, eolian cross-laminae. Local fluvial beds are thinner, up to 20 cm thick, with low-angle, small-scale cross-laminae. Colors of 5Y 8/2 and 2.5Y 6/6 measured. Forms bold cliffs. 85-90 m thick. **Summerville Formation** — White to pale yellow, very poorly exposed sandstone, mudstone, and limestone conglomerate(?). One poor exposure suggests the presence of limestone pebble conglomerate channels, with subrounded to rounded, poorly sorted fine to medium pebbles in a silty fine to coarse sand matrix; given the poor exposure, however, it is possible this conglomerate is actually a caliche-cemented Quaternary deposit inset upon the Summerville sandstones. Colors of 2.5Y 8/1 to 7/3 measured. 13-16 m thick. Todilto Formation – Gypsiferous light to dark gray fine-grained limestone. Typically <5% of fresh faces is visible grains (locally up to 20%), with <1-2% fine (<1 mm across) circular fossils. Thinly bedded, 0.5 to 10 cm thick. Colors of 2.5Y 6/1 to /1 measured. 5-7 m thick. Entrada Sandstone – Pink to light reddish brown silty fine sandstone and siltstone. Grains are moderately to poorly sorted, very silty with rare medium sand grains. Indistinctly thinly bedded (2-10 cm thick), but well cross-bedded. Color of 2.5YR 7/4 measured. 45-50 m thick. **Triassic Rocks Friassic rocks, undivided** — Cross section only. Includes both **F***c* and **F***m*. 540-550 m thick. Chinle Formation, undivided – Poorly exposed red mudstone at the top, reddish brown sandstone and pebble conglomerate at the base, middle of unit is not exposed. Basal contact placed on the first pebble conglomerate.

laminated mudstone, commonly wavy, up to 4 mm thick. Local very fine sandstone beds up to 3 cm thick. Colors of 2.5Y 5/3, 7/2, and black (value <2.5, chroma ~1) measured. Unit includes a ~2 m thick sandstone interval dividing two mudstone intervals. 26-30 m thick. **Dakota Sandstone**, lower Oak Canyon Member – Pale brown to white, variably silty fine to medium sandstone and lesser pebbly sandstone and siltstone. Thin beds (1-10 cm thick beds) that are typically more massive at the base and internally planar and cross-laminated toward the top. Vertical and subhorizontal burrows are common. Locally carbonaceous. Pebbles at the base are up to ½ cm across, poorly sorted, generally well rounded, and principally of light to medium brown siliceous material with rare quartzite and sparse granite. 8-10 m thick. Jurassic Rocks

Dakota Sandstone, upper Oak Canyon Member – Black to pale brown gypsiferous claystone to siltstone. Strongly

where not bioturbated (30-50 cm thick where massive, 2-12 cm thick where cross-bedded). Vertical and subhorizontal burrows common. Unit includes a ~2 m thick medial section of mudstone, separating two sandstone intervals. 9-12 m thick.

sandstone. Colors of 2.5Y 7/3-8/3 measured on fresh faces. 9-12 m thick. Mancos Shale, Clay Mesa Tongue – Black to brown claystone to siltstone. Brown color of 10YR 5/4 measured; brown color may be an iron oxide stain. 18-24 m thick. Dakota Sandstone, Cubero Tongue – Pale brown to light yellow variably silty fine sandstone. Thinly bedded, particularly

Dakota Sandstone (including associated tongues of the Mancos Shale) base to 25-40 cm thick at top. Color of 2.5Y 7/4 measured on a fresh face. Gradational basal contact. 9-11 m thick. Mancos Shale, Whitewater Arroyo Tongue – Light gray to dark gray siltstone to claystone. Generally very poorly exposed. Color of 2.5Y 7/2 measured. 4-6 m thick Dakota Sandstone, Paguate Tongue – Light yellow to pale brown, variably silty fine to medium sandstone. A massive base grades up into a bedded but internally massive middle, followed by a bedded and internally cross-stratified top. Beds are 4-20 cm thick, cross-strata 1-12 mm thick. Unit tends to coarsen upsection, from silty very fine to clean fine-medium

Thinly bedded (4-30 cm thick), commonly massive with burrows at the base, grading up into cross-laminated (1-4 mm thick laminae). Colors of 2.5Y 7/1-8/3 measured. Designation as unit "F" after Molenaar (1983), previously referred to as "Gallup, lower tongue" by Thaden et al. (1967) and Zeigler et al. (2012). 12-15 m thick. Mancos Shale, equivalent to the Rio Salado and D-Cross Tongues – Light gray to medium gray to pale brownish gray, well-laminated gypsiferous siltstone to claystone and rare sparry gypsum beds. Thinly bedded, with beds up to 6 cm thick, commonly internally laminated, with laminae <1 to 1 mm thick. 45-55 m thick. **Dakota Sandstone**, **Two Wells Tongue** — Light yellow to white, quartz-rich, variably silty fine sandstone and local siltstone. Typically more massive and thicker bedded base, and more cross-bedded and thinner bedded top. Beds from 1 m thick at

and siltstone located stratigraphically between *Kgf* and *Kge*. 10-15 m thick. **Gallup Sandstone tongue "F"** – Pale brown to pale yellow to gray quartz-rich silty fine sandstone and fine sandy siltstone.

bedded (3-50 cm thick), massive with vertical tubular burrows to well cross-laminated, with 1-3 mm thick laminae. Color of 2.5Y 8/4 measured for a fresh face. Unit is capped by a distinctive brown, well indurated, calcite-cemented, muddy fine sandstone, with colors of 10YR 5/4 to 6/3. Designation as unit "E" after Molenaar (1983), previously referred to as "Gallup, upper tongue" by Thaden *et al.* (1967) and Zeigler *et al.* (2012). 15-25 m thick. Lower tongue of Mancos Shale intercalated with the Gallup Sandstone – Poorly exposed, thinly laminated mudstone

Gallup Sandstone (including associated tongues of the Mancos Shale) **Gallup Sandstone tongue "C"** – Pale brown to light yellowish gray quartz-rich silty fine sandstone. Thinly bedded (5-30 cm thick) with common internal cross-laminae 1 mm to 1 cm thick. Color of 10YR 8/4 measured for a fresh face. Designation as unit "C" after Molenaar (1983), previously referred to as "Gallup, main body" by Thaden et al. (1967) and Zeigler et al. (2012). 12-20 m thick. **Upper tongue of Mancos Shale intercalated with the Gallup Sandstone** – Poorly exposed, thinly laminated mudstone and siltstone located stratigraphically between *Kge* and *Kgc*. 30-36 m thick. Gallup Sandstone tongue "E" – Pale brown to light yellowish brown, quartz-rich silty-clayey fine sandstone. Thinly

faces. 34-40 m thick. Mulatto Tongue of the Mancos Shale- Yellowish brown mudstone, mostly siltstone. Strongly planar laminated/bedded mudstone, with wavy beds 2 to 30 mm thick. Colors of 2.5Y 7/4 to 8/3 measured for fresh surfaces, 10YR 7/6 measured for weathered surface. Includes a thin quartz-rich very fine sandstone interval. 50-60 m thick. "Stray" Sandstone Member – Pale brown quartz-rich clean fine sandstone. Thinly bedded, 10-40 cm thick, and massive to weakly internally cross-stratified. Colors of 2.5Y 8/3 to 7/4 measured on fresh surfaces. Includes a 2-3 m thick shale interval, as well as local mudstone interbeds (<5% of exposures). 20-25 m thick. **Dilco Coal Member** – Heterolithic unit of interbedded pale brown to gray siltstone, pale brown to tan sandstone, and black coal. Thin bedding to thick laminae; sandstone beds are typically 5-25 cm thick and massive or internally crosslaminated with 1-4 mm thick laminae, while mudstone beds are 3 mm to 3 cm thick, and coal beds are up to 1 cm thick and discontinuous. Sandstone color of 2.5Y 8/4 measured on a fresh face. 35-45 m thick.



### NMBGMR Open-file Geologic Map 241 **Last Modified April 2013**

## **CORRELATION OF MAP UNITS**



