

Geologic map of the Guaje Mountain quadrangle, Los Alamos and Sandoval Counties, New Mexico.

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COMMENTS TO MAP USERS

A geologic map displays information on the distribution, nature, orientation, and age relationships of rock and deposits and the occurrence of structural features. Geologic and fault contacts are irregular surfaces that form boundaries between different types or ages of units. Data depicted on this geologic quadrangle map may be based on any of the following: reconnaissance field geologic mapping, compilation of published and unpublished work, and photogeologic interpretation. Locations of contacts are not surveyed, but are plotted by interpretation of the position of a given contact onto a topographic base map; therefore, the accuracy of contact locations depends on the scale of mapping and the interpretation of the geologist(s). Any enlargement of this map could cause misunderstanding in the detail of mapping and may result in erroneous interpretations. Site-specific considerations should be verified by detailed surface mapping or subsurface exploration. Topographic and cultural changes associated with recent development may not be shown.

Cross sections are constructed based upon the interpretations of the author made from geologic mapping, and available geophysical, and subsurface (drillhole) data. Cross-sections should be used as an aid to understanding the general geologic framework of the map area, and not be the sole source of information for use in locating or designing wells, buildings, roads, or other man-made structures.

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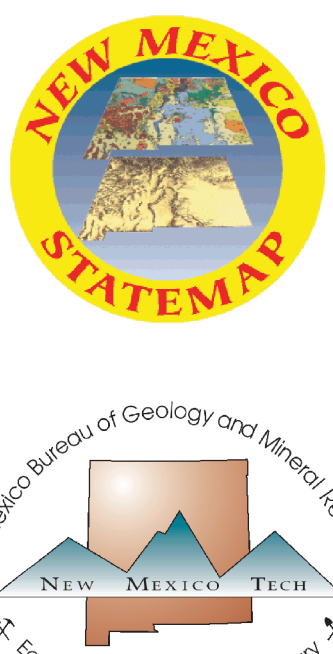
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Open-file Geologic Map 55

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Qal	Alluvium. Silt, sand and gravel deposited in modern drainages; may include terrace deposits. Streams draining the Sierra de Los Valles are particularly laden with sand and gravel of the rhyolitic or dacitic derived from the Tschicoma Formation; stream deposits within the Pajarito Plateau contain abundant clasts and phenocrysts grains of the Bandler Tuff. Holocene to Pleistocene in age (Phillips <i>et al.</i> , 1998; Reneau, 2000) with a maximum thickness of ~ 5 meters.
Qes	Eolian and sheet wash deposits. Tan to brown silt to clay-rich soil that develops primarily on Cerro Toledo interval on the east side of the map area. Soil material is transported in by the wind, but is reworked by slope wash. Holocene to Pleistocene in age with an average thickness of ~ 1 meter.
Qc	Colluvium. Talus debris that typically form wedge-shaped deposits adjacent to cliffs (in particular the Tshirege Member of the Bandler Tuff, QcB). Debris is unsorted to poorly-sorted, often obscuring faults and/or contacts between rock units. Although abundant, these deposits were mapped in only a few locations, particularly on north facing slopes. Holocene to Pleistocene in age with a maximum thickness of ~ 15 meters.
Qvcc	El Cajete Pyroclastic Beds, East Fork Member, Valles Rhyolite. White to beige pyroclastic fall deposits containing pumice, ash, crystals, and lithic fragments. The pumice contains ~ 10% phenocrysts of quartz sandstone, and biotite. The deposits in this are not primary, but are reworked. 50 to 60 ka (Reneau <i>et al.</i> , 1986) < 1 m thick.
Qg	Terrace gravel. Fluvial pebble to boulder-sized gravel composed primarily of Tschicoma Formation lava or Bandler Tuff, deposited along the edges of modern drainages at levels as much as 30 m above the present canyon floor. Terraces are particularly well developed in Rendija Canyon (Reneau and McDonald, 1996; Reneau <i>et al.</i> , 1996). Terraces are also well preserved in eastern Pueblo Canyon, Guaje Canyon, and Sawyer Canyon. El Cajete pumice sits on a low terrace on the south side of Pueblo Canyon near the eastern edge of the map. Holocene to Pleistocene in age (Reneau, 2000) with a maximum thickness of ~ 30 meters. In Pueblo Canyon Qg is the higher terrace level; QgT is lower. El Cajete pumice overlies QgT at the eastern edge of the quadrangle.
QgT	Alluvial fan deposits. Tan silt and gravel overlying Cerro Toledo deposits or Bandler Tuff east of the Sawyer Canyon and Pajarito fault zones in the eastern part of the map area. The gravel is derived from Pyge Formation eroding from the footwall of the faults. Pleistocene in age and ~ 1-3 meter thick.
Qvcc	Landslides. Two types: slump or block slides that remain nearly intact after detaching from a steep slope or cliff, or unsorted, chaotic debris embayed during a chaotic detachment event from a steep slope or cliff. The latter deposits are typically fan-shaped on the valley floor. Holocene to Pleistocene in age of highly variable thicknesses.
Qcta	Cerro Toledo deposits (see below) with younger alluvium and/or colluvium, undivided. These deposits typically occur in valleys where the upper portions of the underlying Cerro Toledo tephras have been reworked and are mixed with younger alluvial and colluvial material. Younger material typically includes reworked vitric pumices of Bandler Tuff (<i>Thakani</i> or <i>Thirge</i> Member) or clasts of devitrified Bandler Tuff. Alluvial material may also include younger terrace deposits, with rounded cobbles and boulders of Tschicoma lavas.
Qalv	Old alluvium. Clasts of rounded Tschicoma Formation dacite to rhyolite and rounded Bandler Tuff preserved in a southeast-trending paleochannel cut into Cerro Toledo deposits in upper Cabe Canyon.
Qalo	Older alluvium. In the town site of Los Alamos, this unit is largely composed of alternating gravel and sandstone layers of fluvial origin interbedded with beds of tephra. Some gravel lenses are dominated by pumice, but most lenses contain subequal amounts of pumice and Tschicoma Formation lava, and rare flow-banded rhyolite and obsidian granule to pebble-sized gravel. The tephras beds are < 1 m thick; the pumice is likely derived from the Cerro del Medio dome to the west of Los Alamos erupted at ~ 1.1 Ma (Spell and Harrison, 1993). The older alluvium is well exposed in a road cut just southwest of the Larry R. Walkup Aquatic Center on Canyon Road in Los Alamos. Northeast of Los Alamos, the unit is dominantly gravel, including clasts of Bandler Tuff; pumice is absent in these deposits. The older alluvium generally sits on Bandler Tuff. North of Los Alamos, Qalo contains abundant clasts of streaky, flow-banded rhyolite lava, black, gray tuff, and some obsidian. Qalo was deposited as fans shortly after the eruption of the Bandler Tuff prior to the development of the modern drainage system (Reneau. Pleistocene in age with a maximum thickness of ~ 30 meters.
Qlx	Upper Bandler Tuff, Tshirege Member. Beige to orange to gray, poorly-sorted to densely-welded ignimbrite containing abundant phenocrysts of sandine and quartz and trace amounts of clinopyroxene and hypersthene (Smith and Bailey, 1966). Sandine commonly displays black indurcescence. Accidental lithic fragments typically < 5% except in discrete lenses and lag horizons. Qlx is a compound flow unit with multiple flow units as described by Broxton and Reneau (1995) and Gardner <i>et al.</i> (1999, 2001), and is well exposed in canyons throughout the Pajarito Plateau. Surge deposits, typically < 1 meter thick, are common at the base of the unit and overlie a stratified pumice tephra (<i>Thakani</i> Member) that is typically ~ 1 meter thick. The ignimbrite was erupted from the Valles caldera at 1.25±0.01 Ma (Phillips, 2004) and filled valleys along the paleotopographically complex eastern flank of Sierra de Los Valles. The degree of welding decreases toward east and some of the upper cooling units pinch out toward the east. Maximum observed thickness is ~ 180 meters.
Qct	Cerro Toledo deposits. Tephras, reworked tephras and epiclastic alluvium deposited between eruptions of the Tshirege and Osovi members of the Bandler Tuff (Broxton and Reneau, 1995). Alluvial deposits include sand, silt, and conglomerates with rounded clasts of Tschicoma lavas. Aphyric obsidian common. Funglomerate alluvial deposits in Los Alamos Canyon record a major fluvial system during this interval. Tephras originated from rhyolite dome complex eruptions in the Toledo caldera and the Toledo embayment (Smith <i>et al.</i> , 1970; Heiken <i>et al.</i> , 1986). Tephras deposits typically contain pumices with rare phenocrysts of quartz, sandine and biotite. In the northern portion of the map area crystal-rich tephras occur near the top of the interval, with abundant phenocrysts of quartz and sandine. Radiometric ages of pumices and source domes range from 1.21 to 1.64 Ma (Zett <i>et al.</i> , 1981; Stix <i>et al.</i> , 1988; Spell <i>et al.</i> , 1996). Maximum exposed thickness is ~ 100 meters.
Qso	Lower Bandler Tuff, Osovi Member. White to beige poorly-welded ignimbrite with abundant phenocrysts of quartz and sandine (Bailey <i>et al.</i> , 1969). Abundant accidental lithic fragments of dark grey to red mafic rocks give outcrops a peppered appearance. Qso is primarily a slope-forming unit and was erupted from the Toledo caldera at 1.61±0.01 to 1.62±0.04 Ma (Zett and Orndorff, 1994; Spell <i>et al.</i> , 1996). Osovi Member ignimbrite was deposited primarily in the southeastern portion of the quadrangle and was not deposited north of Rendija Canyon. The Guaje Pumice Bed, a stratified pumice fall and surge deposit that preceded Osovi ignimbrite deposition, is exposed in isolated outcrops in Cabe Canyon (Qbp) and beneath the Osovi Member in Pueblo Canyon. The Guaje Pumice Bed, which is 2-3 m thick, also crops out the contact between the Pyge Formation and Cerro Toledo deposits in Guaje Canyon. Maximum exposed thickness is ~ 50 m.

TP	Pyge Formation. Light to dark gray fanglomerates derived from Tschicoma lavas (Griggs, 1964; Gardner <i>et al.</i> , 1986; Wanaschak, 1986; Turbell <i>et al.</i> , 1989). Silt and sand layers interbedded with cobble to boulder-sized conglomerates deposited as debris flows, lahars, and alluvial fans. Clasts are typically rounded to sub-rounded and include a wide variety of dacite and rhyolite lavas. Rhyolite of Rendija Canyon and dacite of Caballo Mountain clasts are common. Lower Pyge Formation deposits that overlie Santa Fe Group sediments east of the study area include clasts of rhyolite of Rendija Canyon, suggesting a major eastward propagation of Pyge Formation following the eruption of the Rendija Canyon rhyolite at ~ 5.0 Ma. Reneau and Dehner (1996) note that deposition of the Pyge Formation was probably controlled by blockage of the Rio Grande by volcanism in the Cerros del Rio volcanic field located to the southeast of the quadrangle. Age range of exposed Pyge Formation is ~ 2.4 to 5.3 Ma (WoldeGabriel <i>et al.</i> , 2001). Maximum thickness is ~ 100 meters.
TPs	Thin beds of tephras and/or ignimbrites (TPi) are interbedded with these sediments. In the middle reaches of Guaje Canyon, on the south side of the canyon, ~0.5 m layers of tephras and ignimbrite are interbedded, the entire deposit is 6 to 10 m thick. The contact relationships in the middle portion of Guaje Canyon are obscure, but this deposit appears to lie between Rendija Canyon rhyolite and Pyge Formation. The tan ignimbrite contains abundant gray pumice fragments, quartz, sandine, obsidian, and lithic fragments. A second set of exposures occurs on the south side of Guaje Canyon along the old road into the canyon. One tephra is exposed near the canyon bottom and one is near the top of the Pyge Formation. The upper tephra gives a 40Ar/39Ar age of 4.01 ± 0.22 Ma. The tephras are clay-supported pumice fall < 1 m thick. Clasts range in size from 0.5-10 cm. Clasts of large (1-3 mm), fresh, euhedral amphibole crystals are present throughout the canyon. Some clots include small, uncommon felsic minerals that are impossible to identify using a binocular microscope. Other clots have rare occurrences of biotite, but, for the most part, these clots are monomineralic. Lithics include light gray clasts of glassy and stony lava. Clay is pervasive as an alteration product, but a gritty texture suggests that primary glass may still be present. Maximum thickness is ~ 6 meters.

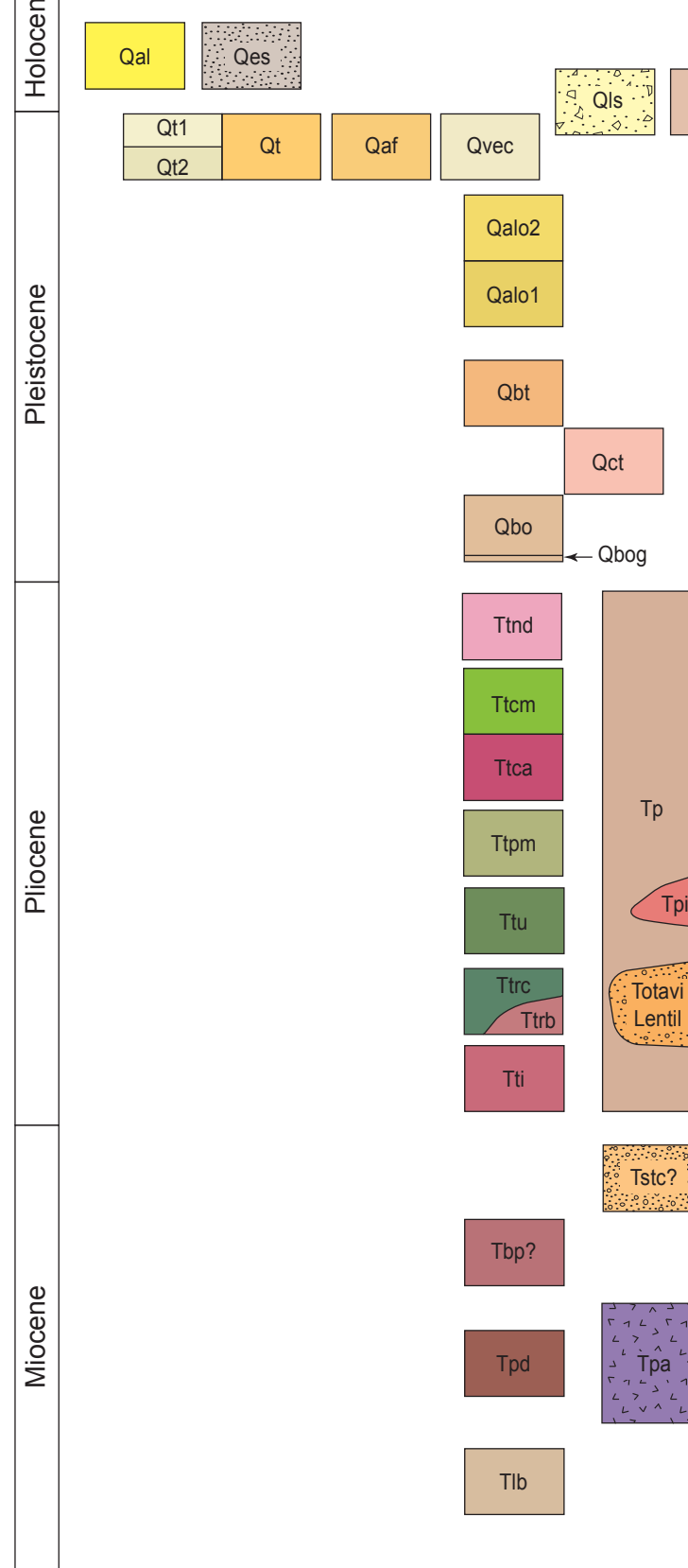
Tschicoma Formation (Pliocene)

“Northern dacites”, Tschicoma Formation (shown only on cross-section B-B’). Crystal poor dacite with 25% phenocrysts of plagioclase, clinopyroxene and orthopyroxene encountered in well H-19. 40Ar/39Ar ages of 2.36±0.54 and 2.49±0.23 Ma have been determined from similar lavas in a drillhole to the south of the quadrangle (Samuels *et al.*, 2007; Broxton *et al.*, 2007).

Dacite of Caballo Mountain, Tschicoma Formation. Lavender to dark gray flow-banded dacitic lavas and associated block and ash flows erupted from the vicinity of Caballo Mountain. Streaky, flow-banded lavender to red facies common. Lavas contain 10-20% phenocrysts of plagioclase and hornblende with subordinate pyroxene, biotite and quartz. At least three separate flows occur, the youngest containing more abundant hornblende. These flows extended in two main directions: one extending SSE of Caballo Mountain, capping the north canyon wall of Guaje Canyon, and one that flowed to the ENE of Caballo Mountain, extending to just west of Pine Springs. Sample from top of Caballo Mountain has age date of 3.03 ± 0.15 Ma (Broxton *et al.*, 2007).

Small volume tephras, ignimbrites, block and ash flows, and associated collied deposits. Outcrops are limited to exposures ~ 0.5 km NNW of Pine Springs. Glassy lava fragments in the block and ash flows appear dacitic, with phenocrysts of plagioclase, amphibole, pyroxene and biotite. Stratified tephras and ignimbrite facies dip ~ 15° in an easterly direction. These deposits appear to be overlain by Caballo Lavas and are therefore likely to be > 3.0 Ma. Maximum thickness is ~ 30 meters.

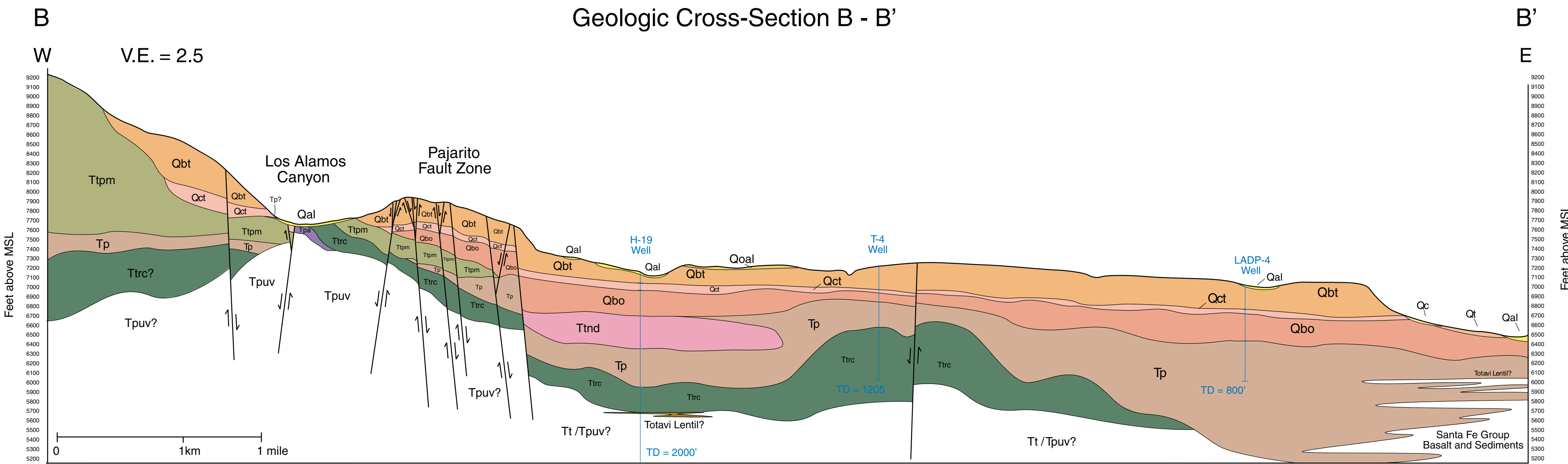
Correlation of Map Units



Legend

- Volcanic vent
- Well
- Cross section line
- Normal Fault - solid where exposed or known; dashed where approximately located; dotted where concealed; queried where uncertain. Bar and ball on downthrown side, tic showing dip.
- Lineament
- scarp
- Contacts - solid where exposed or known;

Geologic Cross-Section B - B'



Geologic Cross-Section A - A'

