

Pennsylvanian System of Chloride Flat Grant County, New Mexico

by D.V. LeMone, W.E.King, J.E.Cunningham



New Mexico Bureau of Mines & Mineral Resources

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ABSTRACT

The Oswaldo Formation (Pennsylvanian) is exposed in faulted, eroded remnants immediately north of Silver City, Grant County, New Mexico. The Oswaldo unconformably overlies the Lake Valley Formation (Mississippian). The Beartooth Quartzite (Upper Cretaceous) unconformably overlies the Oswaldo Formation. The fauna of the Oswaldo consists primarily of foraminifera, rugosan and tabulate corals, crinoids, brachiopods, bryozoa, gastropods, ostracods, some sponge spicules, and trilobite fragments. The fusulinids are represented in three reasonably well-defined zones (in ascending order): *MillerellaEostaffella* zone (Morrow), *Profusulinella* zone (lower Atoka), and *Fusulinella* zone (upper Atoka). Three biohermal to biostromal zones of *Chaetetes milleporaceus* were observed. The algal flora is represented by dasycladacean and phylloid algae. Petrographic analysis of the Oswaldo carbonate sequence indicates a shallow water, normal marine, low to moderate energy, open shelf environment. Silicification, apparently confined to the surface, is recorded. The upper surface of the Oswaldo Formation has a zone of limonite weathering. Patchy recrystallization is noted throughout the sequence. The presence of this Morrow-Atoka sequence of carbonates prompts two interesting paleogeographic speculations. The Oswaldo Formation at Silver City could represent a sag in the positive Florida Islands-Zuni arch axis. The Silver City sag could be an interconnecting link between the marine sediments of the Florida shelf of the Pedrogosa basin and the Robledo shelf of the Orogrande basin. The Silver City sag could also represent the westernmost known Robledo shelf (Lower Pennsylvanian sequence).

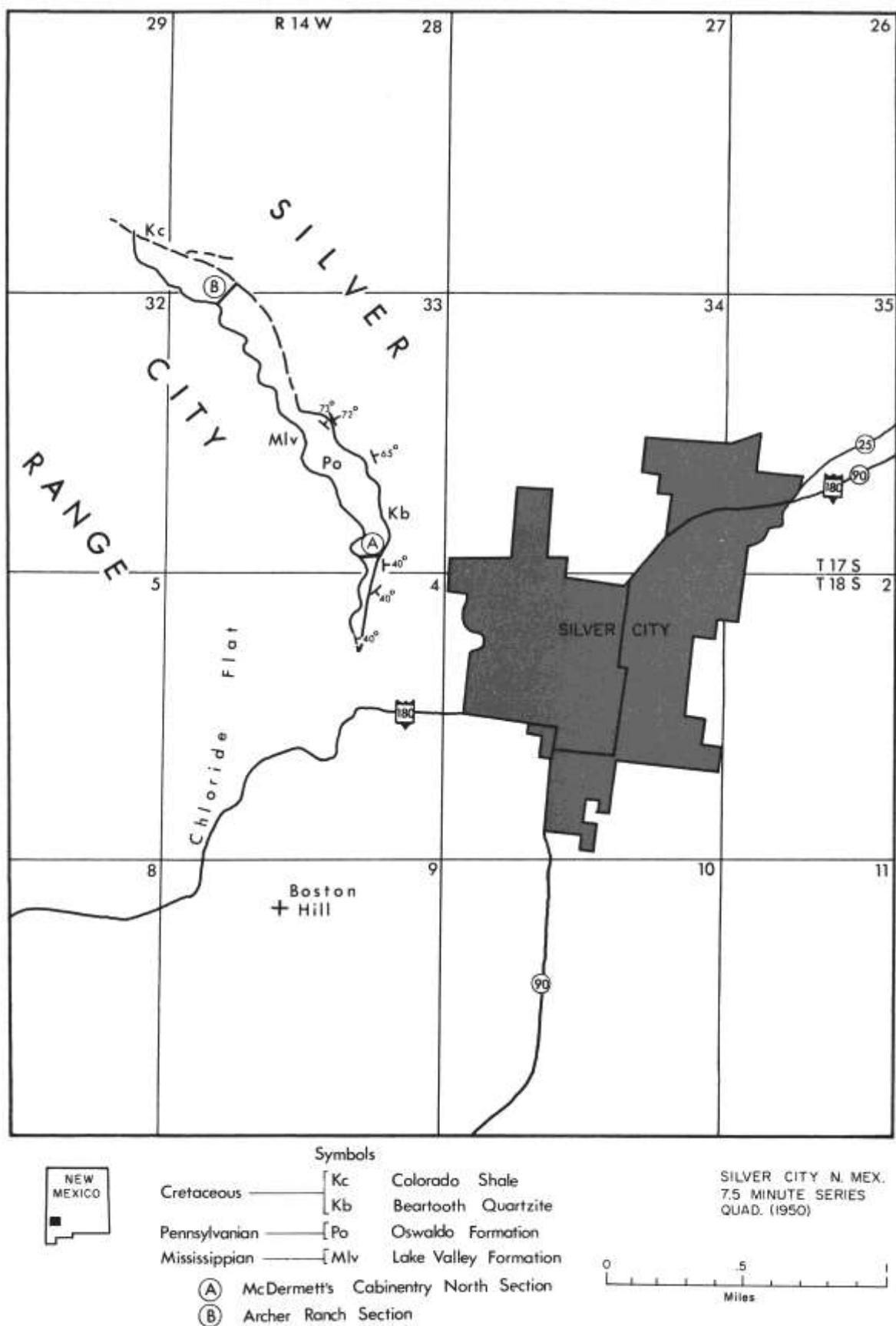


FIGURE 1—Location of the Oswaldo Formation outcrop in the Chloride Flat fault block, Grant County, New Mexico.

INTRODUCTION

Stratigraphic studies completed in 1970 revealed the presence of structurally uncomplicated sequence of the Oswaldo Formation in the Chloride Flat fault block of the southern part of the Silver City Range (fig. 1). Chloride Flat is a valley formed on the easily-eroded Percha Shale at the south end of the Silver City Range, just north of the town of Silver City. The Silver City Range is an eroded east-dipping monocline which exposes a sequence of rocks from Precambrian to Mesozoic age.

Entwistle (1944, p. 23) observed 89 ft of Pennsylvanian beds beneath the Beartooth Quartzite in the Bear Mountain area 6 miles northwest of Silver City. The earlier map of Paige (1916) does not separate the Pennsylvanian. Four known areas of occurrence of Pennsylvanian rocks are now recorded in the Silver City quadrangle.

The 80 Mountain Area has Pennsylvanian strata in the SW/4 of section 19 and the SE/4 of section 8, T. 15 S., R. 14 W. The base of the Pennsylvanian is not exposed in the 80 Mountain Area. The Cretaceous-Pennsylvanian unconformity and a partial section are reasonably well exposed in this area. Pennsylvanian is also recorded in the Cleveland mine area.

The Cleveland mine area is too structurally complex to construct a section. Pennsylvanian rocks underlying the Cretaceous are present in center sec. 2, T. 17 S., R. 14 W.

Structurally, the Bear Mountain fault block is somewhat complex, having been involved with local intrusive activity. The stratigraphic sequence of the Pennsylvanian is not well exposed. Exposures are noted in the Little Walnut Creek area (center S/2 sec. 2, T. 17 S., R. 15 W.) and the South Fork of the Little Walnut Creek (SE sec. 12 and the NE sec. 13, T. 17 S., R. 15 W.). These two exposures exceed the previously reported 89 ft. Possibly the Syrena Formation (Pennsylvanian) and the Abo Formation (Permian) both may be represented in the Little Walnut Creek area.

The most accessible and best Pennsylvanian exposures are observed in the Chloride Flat block (fig. 1). Here the Oswaldo Formation (Morrow-Atoka) is an eroded remnant unconformably underlain by the Lake Valley Formation (Osage) and unconformably overlain by the Beartooth Quartzite or the Colorado Shale (both Gulf). The two sections listed in the Appendix were measured in this block. The southern section is designated the McDermett's Cabin-entry North (SE/4, sec. 33, T. 17 S., R. 14 W.) and consists of 186 ft. of Oswaldo Formation. The Archer Ranch section was measured on the south side of Railroad Canyon (NW/4, sec. 33, T. 17 S., R. 14 W.) and consists of 145.5 ft of Oswaldo Formation. The Dunham carbonate classification was used in megascopic descriptions of lithology. Thickness of rock units were measured in a standard manner with a Jacob's staff. Sediment color is in accordance with the Rock-Color Chart published in 1951 by the Geological Society of America. Thin-section analyses of floral and

faunal content were made (tables 1-6). Carbonates were cross-classified in accordance with Dunham and Folk nomenclature (tables 7-9). (*Tables in appendix.*)

Paige (1916) is the best prior reference for a general geology map of the Silver City quadrangle. Spencer and Paige (1935), Jones and others (1967), Entwistle (1944), and Thompson (1942) are among the major contributors to the Pennsylvanian stratigraphy in the area. The studies of Kottowski (1960, 1963, 1970), Wengerd (1969, 1970), Meyer (1966), and Wilson and others (1969) are important regional references on the Pennsylvanian of southwestern New Mexico. The Silver City quadrangle has recently been remapped by one of the co-authors, John E. Cunningham (to be published by The New Mexico Bureau of Mines and Mineral Resources).

The present study has been supported by the New Mexico Bureau of Mines and Mineral Resources and the University Research Institute of the University of Texas at El Paso.

STRATIGRAPHY

The late Paleozoic rocks of the Silver City area are underlain unconformably by the Lake Valley Formation (Osage) which is a coarsely crystalline, gray, crinoidal limestone (Lochman-Balk, 1965, p. 98). The late Paleozoic rocks are subdivided into 3 distinct formations (fig. 2), in ascending order: the Oswaldo Formation (Morrow-Missouri), the

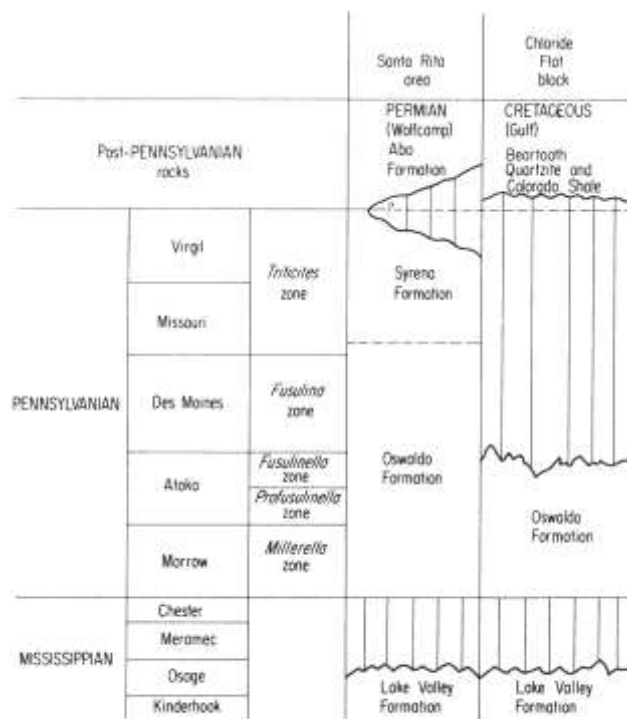


FIGURE 2—Nomenclature chart for the Pennsylvanian rocks of Grant County, New Mexico.

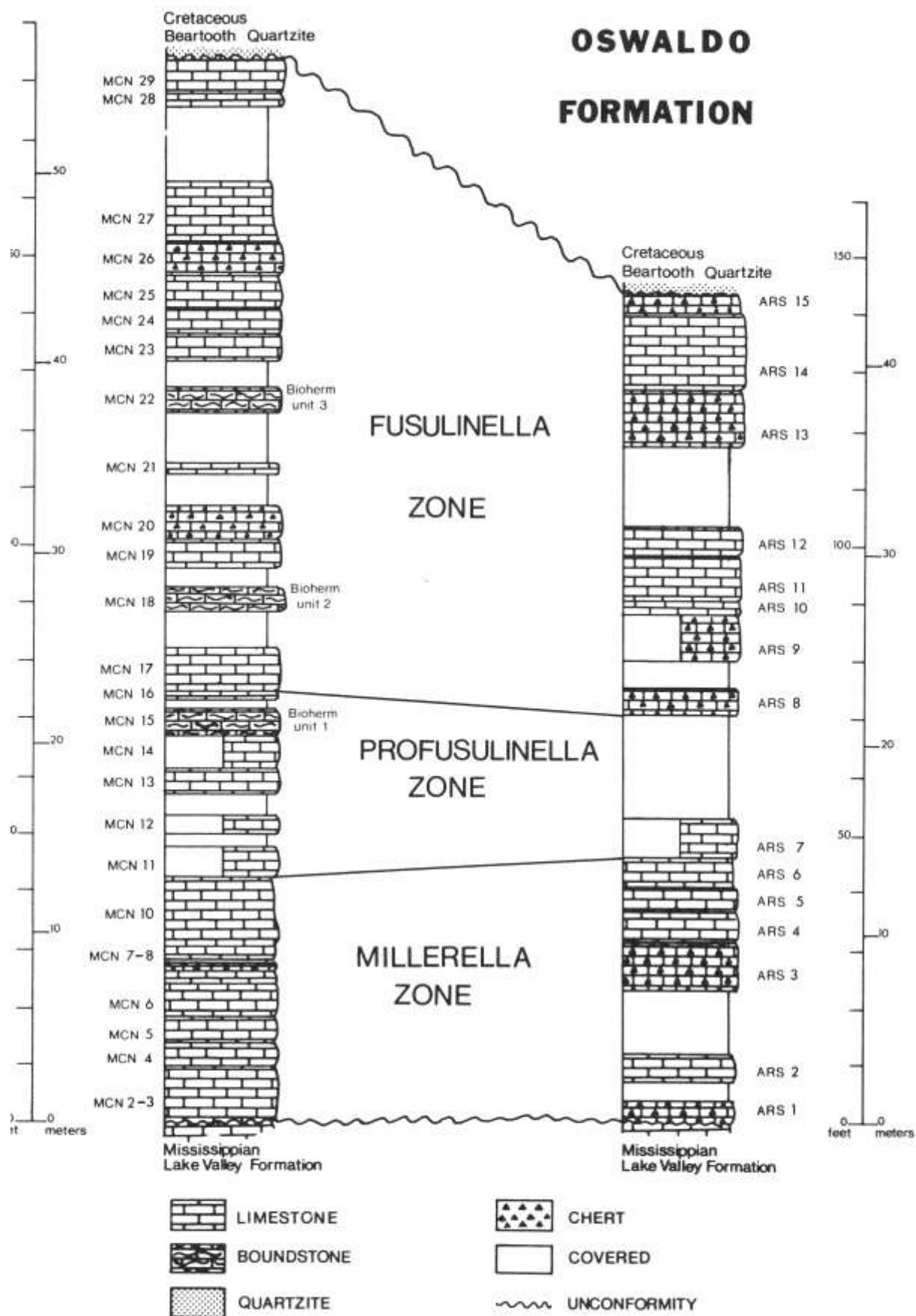


FIGURE 3—Stratigraphic sections of the Oswaldo Formation in the Chloride Flat fault block, Grant County, New Mexico.

Syrena Formation (Missouri-Virgil), and the Abo Formation (probably Wolfcamp). The late Paleozoic rocks are unconformably overlain by either the Beartooth Quartzite (Cretaceous) or Colorado Shale. The Beartooth Quartzite (Probably Gulf) is a fine- to medium-grained quartzite. The Colorado Shale (Gulf) consists largely of limy shales in the lower part of the section.

The Chloride Flat fault block, unconformably overlain by the Colorado Shale and the Beartooth Quartzite, preserves only the lower part of the Oswaldo Formation. The overlying upper Oswaldo, Syrena, and Abo Formations were probably deposited in the Chloride Flat area and were removed by post-Wolfcamp to pre-Gulf erosion. The upper surface of the Oswaldo is iron stained. The lithology of the Oswaldo consists of a sequence of fine- to medium-grained limestone; the upper surface in contact with the overlying Beartooth Quartzite is highly altered by weathering.

The carbonate petrography of the Oswaldo reveals a sequence composed primarily of biomicrites and biomicrudites with sporadic occurrences of intraclastic material (wackestone to packstone in the Dunham classification; see tables 7-9 for petrography). Recrystallization to micro-spar and pseudospar is slight to moderate. Virtually no dolomitization is observed in the sequence. Extensive fracturing of one or more generations is noted in many samples. The fractures are normally filled with secondary sparitic calcite. Pressure solution is demonstrated between grains (photo 1), and by the development of stylolites (photo 2). Sporadic corrosion of individual grains is noted, probably due to coating by a blue-green algae such as *Girvanella*. Silicification, which is slight to moderate, is probably developed as a weathering phenomena.

Spicular biomicrites (photo 3) are recorded in the sequence. Boundstones (biostromal to biohermal in shape and averaging 5 ft in thickness) of *Chaetetes milleporaceus* are recorded in 3 intervals in the Atoka portion of the Oswaldo in the McDermett's Cabinentry North section (photo 4). *Chaetetes* was observed in the talus on the north slope of Railroad Canyon. The *Chaetetes* boundstone units were not located on the south slope of Railroad Canyon, in the Archer Ranch section.

PALEONTOLOGY

The megafauna of the Oswaldo Formation was not studied in detail. The rich fauna includes crinoid ossicles, corals (tabulate and rugosan), echinoid spines, brachiopods (largely productoid and spiriferoid, photo. 5), bryozoans (fistulopodid and fenestrate), gastropods, trilobite exuviae, fusulinids, and pelecypods. The megafauna consists of poorly preserved phylloid algae recorded in up to 7 zones in the Morrow and the Atoka sequences. Stromatolitic layers may also be present. Thin-section analysis reveals an equally varied microscopic fauna and flora (tables 4-6).

The microscopic identifications of the carbonates reveal

an extensive foraminifera fauna that includes *Millerella* (photo. 2), *Eostaffella* (formerly *Paramillerella*) (photo. 6), *Saccaminopsis*, endothyrids, ophthalmid foraminifera (?), *Bradyina*, *Profusulinella* (photo. 7), *Fusulinella* (photo. 8), *Tetrataxis* (photo. 9), *Climacamina* (photo. 10), and *Eoschubertella* (see tables 1-3).

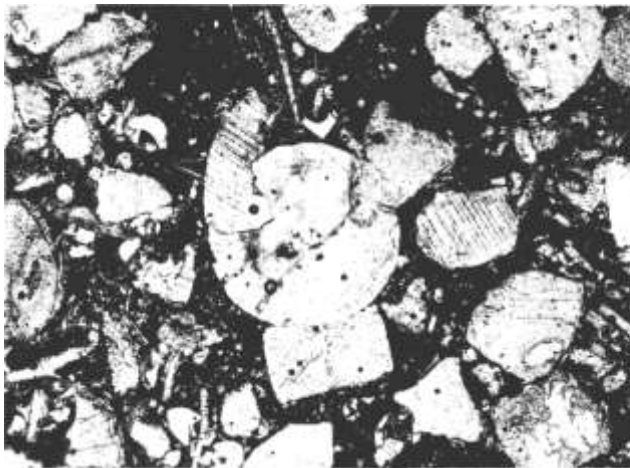
The fusulinid sequence enables precise zonation of the Oswaldo Formation (fig. 3). Random orientation of fusulinid specimens in thin section makes specific identification impossible and generic identification difficult. In poorly oriented random sections, advanced *Profusulinella* is difficult to distinguish from *Fusulinella*. In each case where *Fusulinella* is designated, a clear and distinct diaphanoteca was observed.

The subdivisions revealed by thin section are, in ascending order: the *Millerella-Eostaffella* zone (Morrow), *Profusulinella* zone (lower Atoka), and the *Fusulinella* zone (Atoka). The McDermett's Cabinentry North section develops a 42-ft *Millerella* zone (MCN 2-10), a 31 1/2-ft *Profusulinella* zone (MCN 11-16), and a 112 1/2-ft eroded *Fusulinella* zone (MCN 17-29). The Archer Ranch section to the north closely approximates these thicknesses which are a 46-ft *Millerella* zone (ARS 1-6), a 25-ft *Profusulinella* zone (ARS-7), and a 74 1/2-ft eroded *Fusulinella* zone (ARS 8-15). Variations in thickness can be attributed to position of a sample in the unit used for thin section and the normal variations in sedimentology that would be recorded on an open shelf environment. All of the megafauna discussed are represented in thin section examples. The flora represented in thin section includes recrystallized phylloid algal blades (photo. 11), "*Girvanella*" corrosion, and dasycladacean algae. The presence of Morrow and Atoka phylloid algal blades is noteworthy, being the second recorded occurrence of Morrow phylloid algae in the region. The other occurrence is in the Franklin Mountains north of El Paso, Texas in the La Tuna Formation (Mims, 1971). The overlying Berino Formation (Atoka) in the Franklin Mountains also contains phylloid algae. The traditional exploration for phylloid algae has been limited to rocks ranging in age from Des Moines to Wolfcamp. The possibility of developing phylloid banks in Morrow and Atoka rocks now needs to be seriously reexamined in current exploration programs throughout the Southwest and elsewhere.

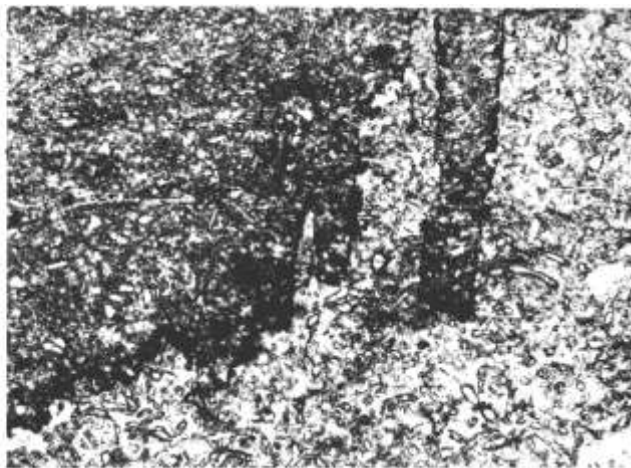
PALEOGEOGRAPHY AND PALEOECOLOGY

The paleoecology of the Oswaldo Formation of the Chloride Flat fault block reveals an environment that would be normal marine, warm shallow water with low to moderate energy on an open shelf.

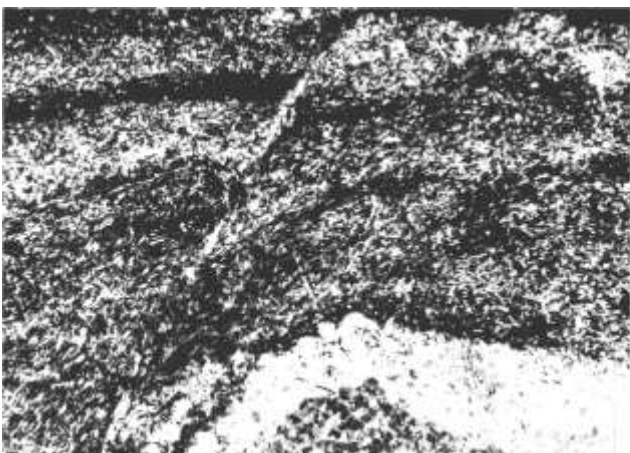
The paleogeographic implications of this interpretation indicate that by position this section would be very close to the crest of a Defiance-Zuni arch-Florida axis positive element if they formed a continuous NNW-SSE axis (fig. 4). Two conclusions are possible with this interpretation. The



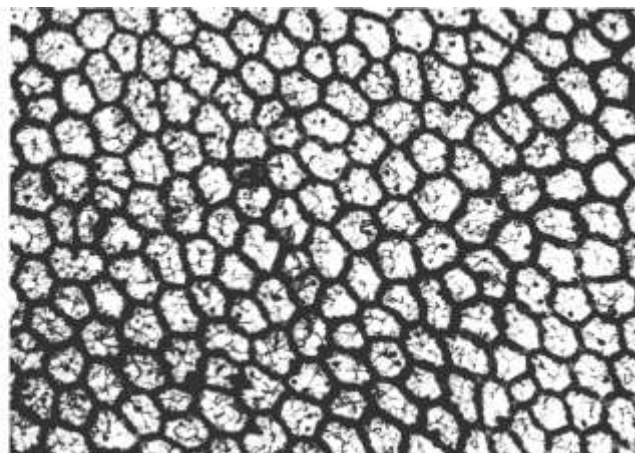
PHOTOMICROGRAPH 1—*Crinoidal* biomicrudite demonstrating recrystallization of crinoidal ossicles and embayed grain formation by pressure solution (MCN-3). x 11



PHOTOMICROGRAPH 2—Stylolite formation and sections of *Millerella* capped by limonite (MCN-5). x 11



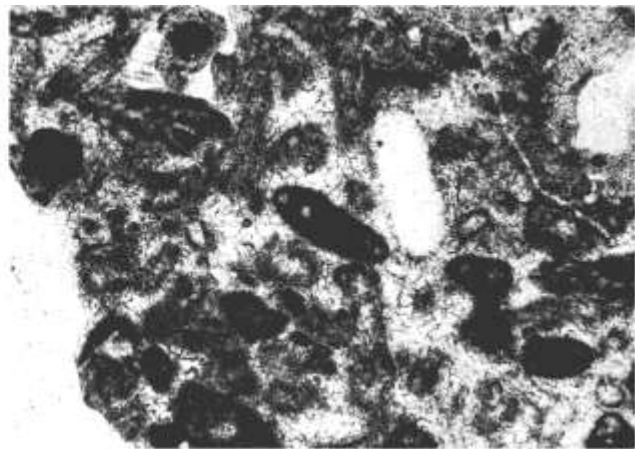
PHOTOMICROGRAPH 3—Spiculitic biomicrite (MCN-5). x 11



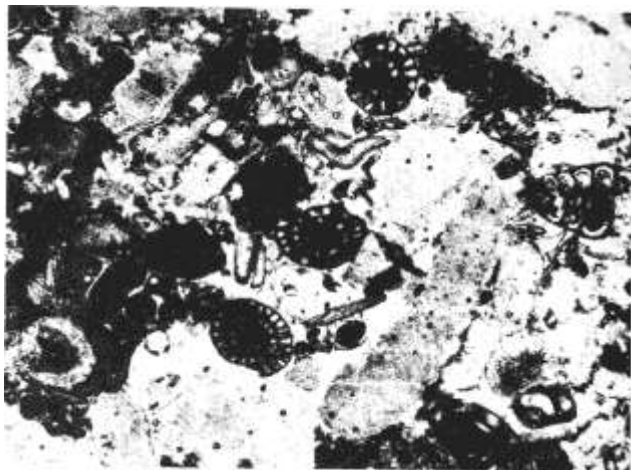
PHOTOMICROGRAPH 4—*Chaetetes* milleporaceous boundstone (MCN-22). x 11



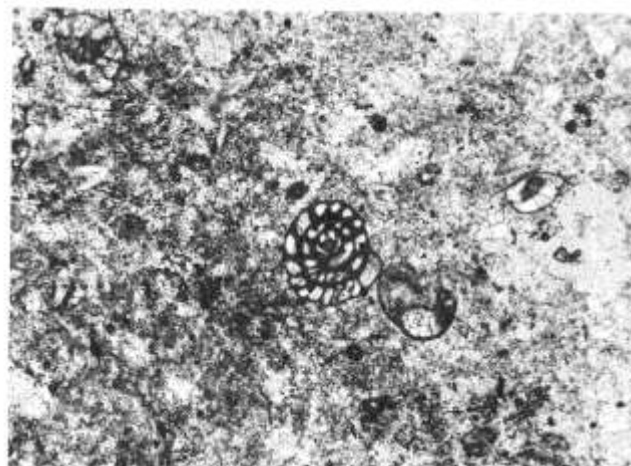
PHOTOMICROGRAPH 5—Biomicrudite with productid brachiopod demonstrating geopetal structure (ARS-10). x



PHOTOMICROGRAPH 6—*Eostaffella* sp. in a recrystallized biomicrite 33X (MCN-4). x 11



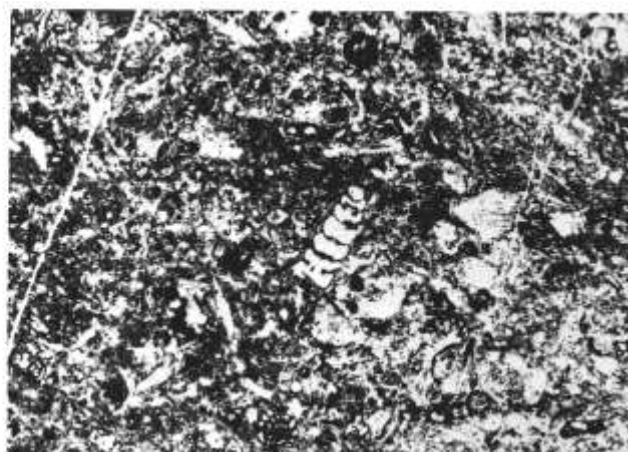
PHOTOMICROGRAPH 7—*Profusulinella* sp. with a *textularid* foraminifera (MCN-13). x 11



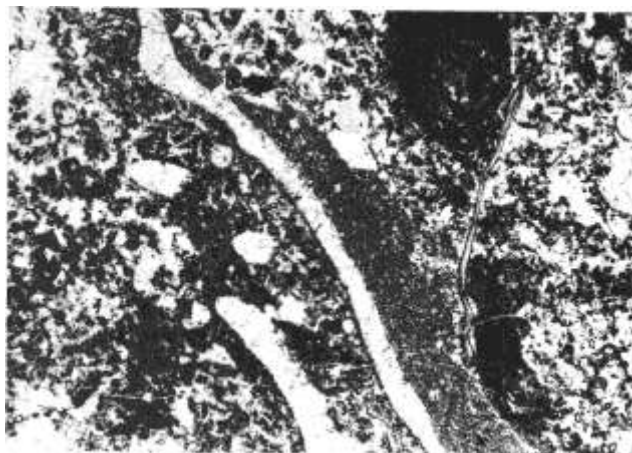
PHOTOMICROGRAPH 8—*Fusulinella* sp. in a recrystallized biomicrite (MCN-132). x 11



PHOTOMICROGRAPH 9—*Textaxis* sp. in a recrystallized biomicrite (MCN-152). x 11



PHOTOMICROGRAPH 10—*Climacammina* sp. in a recrystallized biomicrite (MCN-21). x 11



PHOTOMICROGRAPH 11—Recrystallized phylloid algal blade (MCN-9). x 11

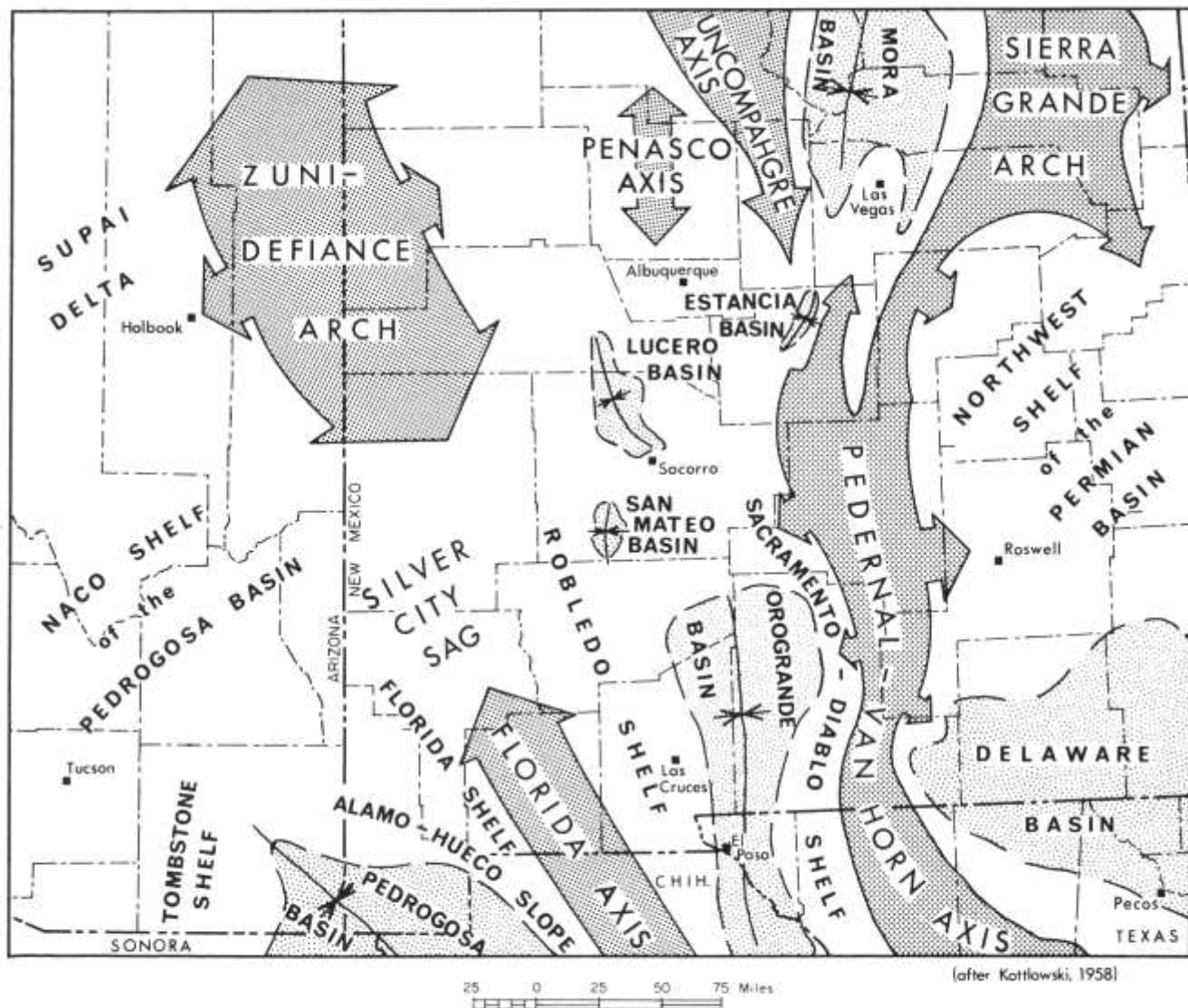


FIGURE 4—Pennsylvanian paleotectonic and paleogeographic map of the southwestern New Mexico region.

Oswaldo Formation (of the Chloride Flat fault block) represents the far western extension of the Robledo shelf open-shelf marine sequence or the far eastern extension of the Florida shelf open-shelf marine sequence. The term Robledo shelf (Meyer 1966) is equivalent to the term Potrillo shelf of Wengerd (1969, 1970).

The second interpretation is that the Defiance-Zuni arch

and Florida axis do not represent one continuous positive element but instead are separated by a sag. The name Silver City sag is proposed for this paleogeographic feature. Probably this area remained a shallow water open-marine region connecting the western Florida shelf sequences with the eastern Robledo shelf sequence.

CONCLUSIONS

The results of this study are particularly interesting because we have observed "*Ivanovia*"-like phylloid algae in rocks bearing Morrow and Atoka fusulinids. The fusulinids assigned for dating are referred to Morrow-Atoka age and the term Derryan is disregarded. The eroded Oswaldo sequence contains a fauna and flora that indicates an open-shelf sequence of low to moderate energy in a normal marine environment.

The occurrences of phylloid algae in Morrow and Atoka sediments should be reexamined regionally to determine whether or not phylloid algae and their biohermal bank structures, which are so commercially feasible, are present and acting as stratigraphic reservoirs for petroleum. This data will require a reevaluation of the petroleum possibilities of the Pennsylvanian system in southwestern New Mexico.

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APPENDIX

TABLE 1 – Archer Ranch series (ARS) foraminiferal distribution

[illegible]

TABLE 2—MCN series foraminiferal distribution

[illegible]

TABLE 5 – MCN series faunal distribution

	SPONGE SPICULES	CORALS	CHAETETES	FENESTRATE BRYOZOA	FISTULOPORID BRYOZOA	BRYOZOA (INDET.)	BRACHIOPOD SPINES	BRACHIOPOD VALVES	GASTROPODS	OSTRACODS	TRILOBITE FRAGMENTS	CRINOIDAL DEBRIS	ECHINODERMATA (OTHER)	DASYCLADACEANS	PHYLLLOID ALGAE	ALGAL CORROSION
X MCN-29-186							X	X				X	X			
MCN-29-183							X					X				
MCN-29-181				X			X		X	X		X				
X MCN-28-173																
MCN-27-160	X									X						
MCN-27-158						X	X	X		X		X				
X MCN-27-153																
X MCN-26																
X MCN-25																
X MCN-24																
MCN-23-135					X							X	X		X	
MCN-23-132				X				X								
MCN-22			X													
MCN-21							X	X				X				
MCN-20								X				X				
MCN-19								X		X						
MCN-18		X	X						X			X				
MCN-17								X	X			X				

TABLE 6 – MCN series faunal distribution

	SPONGE SPICULES	CORALS	CHAETETES	FENESTRATE BRYOZOA	FISTULOPORID BRYOZOA	BRYOZOA (INDET.)	BRACHIOPOD SPINES	BRACHIOPOD VALVES	GASTROPODS	OSTRACODS	TRILOBITE FRAGMENTS	CRINOIDAL DEBRIS	ECHINODERMATA (OTHER)	DASYCLADACEANS	PHYLLLOID ALGAE	ALGAL CORROSION
MCN-16						X	X			X		X	X			
X MCN-15																
MCN-14																
MCN-13					X			X	X				X	X		
MCN-12						X	X		X	X		X		X	X	X
MCN-11						X	X	X		X		X		X	X	
MCN-10						X	X	X				X	X			
MCN- 9												X			X	
MCN- 8										X		X	X			
X MCN- 7												X				
MCN- 6					X			X		X						
MCN- 5	X															
MCN- 4						X		X				X				
MCN- 3						X	X	X		X		X		X		
MCN- 2								X				X				
MCN- 1					X	X			X		X	X	X			

TABLE 7 – Archer Ranch series (ARS) petrography

	ORTHOQUARTZITE	MUDSTONE	MACKESTONE	PACKSTONE	BIOMICRITE	BIOMICRUDITE	INTRACLASTIC BIOMICRITE	INTRACLASTIC BIOMICRUDITE	INTRACLASTS	PELLETS (MINOR)	RECRYSTALLIZED	IRON OXIDE REPLACEMENT	FRACTURED (RESEALED WITH CALCITE)	FRACTURED (LIMONITIC FILLED)	LIMONITIC STAINING	SILICIFIED FAUNA			
ARS-16	X																		
ARS-15			X			X													
ARS-14			X		X	X													
ARS-13				X	X						X								
ARS-12				X			X		X		X		X						
ARS-11			X	X	X						X								
ARS-10			X			X					X				X	X			
ARS-9			X	X	X						X			X					
ARS-8			X		X						X		X						
ARS-7				X		X				X			X						
ARS-6		X	X		X								X						
ARS-5				X	X						X								
ARS-4			X	X	X						X		X	X					
ARS-3			X	X			X		X				X						
ARS-2			X	X	X	X						X	X						
ARS-1			X		X							X	X	X					

TABLE 8 – MCN series petrography

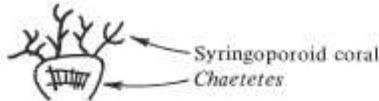
	ORTHOQUARTZITE	BOUNDSTONE	MUDSTONE	MACKESTONE	PACKSTONE	GRAINSTONE	BIOLITHITE	BIOMICRITE	BIOMICRUDITE	BIOSPARITE	INTRACLASTIC OOLITIC BIOSPARITE	MICRITE	INTRACLASTIC BIOSPARITE	INTRACLASTS	OOLITES	RECRYSTALLIZED	FRACTURED (RESEALED WITH SPARITE)	ALGAL CORROSION	STYOLITES	LIMONITE STAINING
BO - 1	X																			X
X MCN-29-186																				X
MCN-29-183				X				X												X
MCN-29-181				X				X								X				
X MCN-28-173																				
MCN-27-160			X	X				X									X			
MCN-27-158				X													X			
X MCN-27-153																				
MCN-27-152				X				X								X				
X MCN-26-150																				
X MCN-25-145																				
X MCN-24-130																				
MCN-23-135					X								X	X		X				
MCN-23-132					X			X								X				X
MCN-22		X					X													
MCN-21				X	X			X												
MCN-20				X	X			X								X				
MCN-19				X				X									X		X	X
MCN-18				X				X								X	X			
MCN-17				X	X			X								X	X			
MCN-16				X				X												
X MCN-15		X					X													
MCN-14			X									X					X			

Located in NW/4, sec. 33, T. 17 S., R. 14W., Silver City Range. In Railroad Canyon the Oswaldo Formation has been recognized above the Tierra Blanca member of Lake Valley Formation (Mississippian). Fusulinids and *Chaetetes milleporaceus* were collected on the north slope of the canyon. Measurements in the arroyo of the canyon indicate a minimum of 62 feet of Oswaldo. The Colorado Shale (Cretaceous) is observed beyond the overlying 38 feet. A fault zone is strongly indicated. Section (ARS series) was taken on the better exposures on the south canyon wall.

Mudstone, wackestone, packstone, grainstone, and boundstone refer to Dunham's carbonate classification scheme which is adequate for megascopic identification. All carbonates are limestones.

Sample No.	Description	Thickness (ft)		
ARS-16	Beartooth Quartzite (Cretaceous): Float of questionable fault gouge and Beartooth Quartzite. Iron staining extensive Fresh—grayish orange 10YR7/4, grayish red 10R4/2		ARS-8	gray, fresh surface dark gray to black. Contains forams (?) and gastropods. White chert with fusulinids. Fair to poor exposure Weathered—medium gray N5 Fresh—medium dark gray N4 Covered
				8 5
ARS-15	Oswaldo Formation (Pennsylvanian): <i>Fusulinella</i> zone Recrystallized limestone with crinoid columnals and solitary corals. Mottled. Chert make-up 30 to 55% of the total volume Weathered—medium gray N5 Fresh—brownish gray 5YR4/1, grayish red 10R4/2, moderate yellowish brown 10R5/4	145½ 4½		Wackestone with a silicified fauna contains fusulinids. Minor chert light gray in color, with a purplish tinge on the fresh surface Weathered—medium light gray N6 Fresh—brownish gray 5YR4/1
ARS-14	Wackestone containing large rugosoid corals, fusulinids, brachiopods, and questionable fragmental phylloid algae. At 135 ft wackestone has a banded look on the weathered surface Weathered—light brownish gray 5YR6/1 Fresh—medium dark gray N4		ARS-7	<i>Profusulinella</i> zone Covered with rubble Encrinal wackestone and packstone. Poor exposures as in the entire section. Red siliceous zone at the top followed by covered interval. Thinner bedded than ARS-6 Weathered—medium gray N5 Fresh—medium dark gray N4
ARS-13	Dark-gray wackestone containing fusulinids, crinoid ossicles, and other fine faunal debris, with banded dark-brown chert. This unit grades up into a recrystallized, mottled purple and gray, medium-grained limestone, which has features of soft sediment deformation Weathered—medium light gray N6 Fresh—medium dark gray N4	13 10	ARS-6	<i>Millerella</i> zone Largely mudstone with some encrinal grainstone. Mudstone contains irregular beds of graded crinoidal debris and phylloid algae. Cross sections of brachiopods. Thin-bedded but weathers massive. Forms a rubble slope Weathered—light gray N7 Fresh—olive gray 5Y4/1
ARS-12	Wackestone and packstone containing fusulinids, crinoids, and irregularly bedded (graded) phylloid algae. Upper ft contains a chert and limestone unit. White chert has distinct fusulinids. Intraclasts of one inch. Iron stained Weathered—medium light gray N6 Fresh—medium dark gray N4 Covered with rubble	5 14	ARS-5	Largely mudstone with some fine-grained calcarenitic stringers. Contains large crinoid columnals. No slabby surfaces Weathered—light gray N7 Fresh—medium dark gray N4
ARS-11	Medium-bedded wackestone with forams (?), crinoidal debris, and fragments of unidentified fusulinids Weathered—medium light gray N6 Fresh—medium dark gray N4	8	ARS-4	Gray mudstone with encrinal packstone and grainstone stringers. Considerable hematite. Poor exposure. Contains large crinoid columnals Weathered—light gray N7 Fresh—medium gray N5, grayish red 10R4/2
ARS-10	Limestone. Medium to coarse grained in a mottled unit. Silicified fauna (crinoidal). Wackestone in part with good graded faunal fragments in micrite. Weathered—medium light gray N6 Fresh—medium dark gray N4	5 2	ARS-3	Silicified, excellently preserved brachiopod in a mudstone matrix, gastropods, phylloid algal debris. Poor exposure. Some encrinal debris with brachiopod fragments that indicates reworking. Contains red brown flattened lenticular chert Weathered—medium light gray N6 Fresh—medium gray N5 Covered
ARS-9	Mudstone to wackestone weathers distinct light		ARS-2	Alternating mudstone and encrinal packstone with gastropods and solitary corals (7-8.5). 8.5-11 ft covered. Flocculent hematite in medium grained limestone (11-12). Questionable microfauna. Phyl-
				8 11

MCN-20	Fresh—medium dark gray N4	2	MCN-11	Weathered—light gray N7	
	Covered	5		Fresh—light olive gray 5Y5/2	5
MCN-19	Gray wackestone, micritic. Some lens-like irregular chert, tan in color at 103 ft. Lower 3 ft black on fresh surface and with fusulinids very abundant. Upper 3 ft more gray red with fusulinids very abundant		MCN-10	Limestone. Poorly exposed. Fusulinids, forams crinoids, appears to be fine-grained packstone. Unidentifiable fragments	
	Weathered—light gray N7			Weathered—medium light gray N6	
MCN-18	Fresh—medium dark gray N4	6	MCN-9	Fresh—pale yellowish orange 10YR8/6	5
	Wackestone with visible fossils, micritic. With fusulinids, mostly fine faunal hash. Distinct 1 to 3 inch limonite-stained zone at 100 ft			<i>Millerella</i> zone	
MCN-17	Weathered—medium light gray N6		MCN-8	Limestone as in MCN-9, micritic, much purer looking. More tan on fresh surface, break in slope on this unit. Fusulinids, crinoids, dasy-cladacean algae?	
	Fresh—olive gray 5Y4/1	5		Weathered—grayish orange 10YR7/4, dark yellowish brown 10YR4/2	
MCN-16	Covered	3½	MCN-7	Fresh—medium light gray N6	5
	Biohermal unit number 2. Massive bench of boundstone— <i>Chaetetes milleporaceus</i> with syringoporoid corals. Exposure good, as in the first exposures the hair corals and syringopora are silicified (weathers tan). High-spined gastropods			Limestone. Forms massive cliff, considerable limonite replacement after crinoids, questionable phylloid algae, brachiopods, first fusulinids (observed megascopically). Micritic	
MCN-15	Weathered—light olive gray 5Y6/1		MCN-6	Weathered—light gray N7	
	Fresh—light olive gray 5Y5/2	3½		Fresh—medium gray N5, dark yellowish orange 10YR6/6	7
MCN-14	Covered rubble	6½	MCN-5	Limestone. Basal 2½ ft massive with 6 x 3 inch chert nodules at 2 ft, dark gray. Top 1½ ft nodular limestone with crinoidal debris and brachiopods, gastropods with 4 x 47 mm echinoid spines. Dark gray at 30 ft	
	Wackestone, micritic (top 6 inches very cherty). At 80 ft syringoporoid coral colony—(6 inches by 3 to 4 inches). Brachiopods, crinoids. Fusulinids. Unit is more or less massive covered with rubble			Weathered—light brownish gray 5YR6/1	
MCN-13	Weathered—light gray N7		MCN-4	Fresh—medium dark gray N4	5
	Fresh—olive gray 5Y4/1	8		Limestone, with 2- to 3-inch bands of gray-brown chert, at base of massive, limestone gray brown on fresh surface, unit micritic like the underlying MCN-6	1½
MCN-12	Covered (Shifted section laterally)		MCN-3	Limestone. Micritic, weathered and fresh surface gray, questionable microfauna	
	<i>Profusulinella</i> zone			Weathered—medium light gray N6, moderate yellowish brown 10YR5/4	5
MCN-11	Wackestone. Replete with small fauna across surface (crinoid, etc.) large brachiopods		MCN-2	Chert and limestone. Chert is laminated cream, red and brown and make in excess of 50% of unit; unit would make an excellent keybed. Limestone micritic or recrystallized micrite. No fauna observed. Limestone weathered and fresh surface gray.	
	Weathered—light brownish gray 5YR6/1	1		Weathered—grayish orange 10YR7/4, light olive gray 5Y5/2	
MCN-10	Fresh—brownish gray 5YR4/1	1½	MCN-1	Fresh—grayish orange 10YR7/4, olive gray 5Y5/2	4½
	Covered	5		Limestone. Darker than the underlying unit contains brown chert (minor in extent), darker in color than MCN-3, crinoid debris. Jasperoid material in great abundance, does not affect the fossil material, noticeable in the micritic matrix. Rich zone at the very top of section	
MCN-9	Biohermal unit number 1. Boundstone, lower unit micritic with abundant brachiopods. Light tan in color. <i>Chaetetes milleporaceus</i> . Colonies 12 inches by 12 inches. Syringoporoid colonies also, some in association with <i>Chaetetes</i> . 8 inch gray units alternating with ¼ inch tan units, fetid odor, micritic. Small fusulinids at 70-71		MCN-0	Weathered—medium gray N5	
				Fresh—dark gray N3	4
MCN-8			MCN-0	Limestone. Crinoidal coquina, apparently cemented with sparite, reflects probably high energy environment. Crinoids and (?) foraminifera (not fusulinids)	
				Weathered—moderate brown 5YR3/4	
MCN-7			MCN-0	Fresh—medium gray N5	5
				Limestone. Large crinoid columnals 8 x 12 mm, some brachiopods, small crinoidal debris in fine-grained matrix. Much less in fossils than MCN-1, weathers with pinpoint porosity	
MCN-6			MCN-0	Weathered—pale yellowish brown 10YR6/2	
				Fresh—light olive gray 5Y6/1	5
MCN-5			MCN-0	Lake Valley Limestone (Mississippian):	



Syringoporoid coral
Chaetetes

MCN-1	Base of section in resistant dip slope-former (approximately 6 inches at base) limestone. Surface silicification of crinoid columnals (weathers out tan, encrinal coquina with trilobites, crinoids, ostracods, and brachiopods. Encrinal bioclastic with sparite cement Weathered—light gray N7 Fresh—yellowish gray 5Y7/2
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Type Faces: Camera-ready copy composed on IBM MT
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Subheads 11 pt. Press Roman
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Presswork: Text printed on 38" single color Miehle
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Paper: Cover on 65# antique white
Body on 70# Mountie Matte

