

THE ANCHA FORMATION: TEXTURAL SUBDIVISIONS, LOWER CONTACT, AND HYDROGEOLOGIC IMPLICATIONS



Daniel Koning and Peggy Johnson

New Mexico Bureau of Geology and Mineral Resources

March, 2005

INTRODUCTION AND PURPOSE

The Ancha Formation is the uppermost basin fill unit in the Santa Fe embayment. It consists of gravel, sand, and silt derived from the southwestern flank of the Sangre de Cristo Mountains. Although there are some compositional differences in the Ancha Formation that reflect compositional heterogeneity of the crystalline basement, the gravel is generally dominated by granite or gneissic granite, with minor amounts of amphibolite, quartzite, and schist. This Plio-Pleistocene deposit is mostly non-cemented and weakly consolidated. It unconformably overlies the Tesuque Formation (Miocene) in the Santa Fe embayment, north-central New Mexico (Spiegel and Baldwin, 1963). The Santa Fe embayment is bounded by the Sangre de Cristo Mountains to the east, Galisteo Creek to the south, the Cerrillos Hills to the southwest, basalt-capped mesas of the Cerros del Rio volcanic field to the northwest, and the Santa Fe uplands underlain by the Tesuque Formation north of the Santa Fe River. The Ancha Formation extends under the Cerros del Rio basalts westward towards the Santa Fe River, although here it is not as thick as in the center of the embayment – probably because of footwall uplift associated with the La Bajada fault (Koning et al., 2002b).

The main purpose for studying the geologic characteristics of the Ancha Formation is to gain insight into its hydrogeologic properties. For example, it has been demonstrated that effective grain size and sorting relates to hydraulic conductivity (Hazen, 1911; Shepard, 1989). Non-cemented, relatively coarse-grained channel deposits in the Santa Fe Group have been shown to have higher hydraulic conductivity values than finer deposits (Sigda and Paul, Appendix F, this report). The general sorting of sand or gravel within a bed also influences the hydraulic parameters of an aquifer, with poorly sorted textures tending to have lower hydraulic conductivity and porosity than well sorted textures (Fetter, 1988). It is thus important to estimate the percentage of coarse channel deposits in a clastic hydrogeologic unit, such as the Ancha Formation, in addition to describing the architecture and connectivity of these channels (Fogg, 1986; Tyler and Finely, 1991; Davis et al., 1993; Dreyer et al., 1993; Gaud, 2002). Finer-grained deposits, particularly floodplain clay and mud beds, can create confined or semi-confined conditions in an aquifer. Describing their thickness and lateral extent is also useful in assessing the groundwater conditions in a hydrogeologic unit.

The basal contact of the Ancha Formation corresponds with an angular unconformity in all observed outcrops, and could potentially affect vertical groundwater flow between the Ancha Formation and underlying Tesuque Formation, particularly where there is also a grain size distinction between the two formations. Noting such features as weathering and pedogenic activity along the basal contact of the Ancha Formation, or changes in the dip of the underlying Tesuque Formation, could also have important implications for assessing groundwater flow between this and underlying hydrogeologic units.

The following presents results from a detailed geologic examination of the Ancha Formation. This work builds on previous study of the Ancha Formation summarized in Koning et al. (2002b). In particular, this study differentiates ancestral Santa Fe River deposits from alluvial slope deposits derived from smaller catchments, approximately delineates gradational textural trends (at a scale of 1:50,000), and examines the basal contact.

PREVIOUS WORK

Spiegel and Baldwin (1963) were the first to formally apply the name Ancha Formation to the Pliocene-Pleistocene gravel, sand, and silt that rest with angular unconformity on the Tesuque Formation. They originally defined a partial type section for the Ancha Formation using a 49 m-thick exposed interval of arkosic, weakly consolidated sediment on the southwest slope of Cañada Ancha, 18 km northwest of Santa Fe. Later work restricted the Ancha Formation to the upper 12 m of Spiegel and Baldwin's type section, with the underlying strata assigned to a coarse unit of the Tesuque Formation (Koning et al., 2002b). Koning et al. (2002b) recognized that south of Interstate 25 the Ancha Formation becomes significantly finer to the west. They thus proposed subdividing the Ancha Formation into a fine alluvial member to the west and a coarse alluvial member to the east. Relatively coarse sediment in the general area between the modern Santa Fe River and La Cienega Creek was interpreted to have been deposited by an ancestral Santa Fe River (Koning et al., 2002b).

With three exceptions, geologic mapping over the past three decades (Bachman, 1975; Johnson, 1975; Booth, 1977; Kelley, 1978; Lisenbee, 1999; Read et al., 1999 and 2000; Koning and

Hallet, 2000; Maynard and Lisenbee, 2002) generally agrees with the mapping of Spiegel and Baldwin (1963). One exception is for detritus shed from the Cerrillos Hills. Because this sediment is compositionally and temporally similar to the Tuerto Formation to the south, Koning and Hallet (2000) include it with the Tuerto Formation rather than the Ancha Formation. A second exception is for the area north-northeast of Cañada Ancha. Here, Spiegel and Baldwin (1963) included granite-rich sandy gravel and gravelly sand found in the western Horcado Ranch and northern Agua Fria quadrangles with the Ancha Formation on their map and type section. However, subsequent geologic mapping by Kelley (1978) and Koning and Maldonado (2001), together with dating of tephras, confirm that the Ancha Formation here is restricted to a 10-20 (?) m-thick sand and gravel deposit immediately beneath flows and tephras of the Cerros del Rio volcanic field (Koning et al., 2002b). Lastly, the correlation of the Ancha Formation to high-level gravel deposits in the northeastern Espanola Basin (Miller et al., 1963) was rejected by Manley (1976 and 1979).

Past studies of the Tesuque Formation in the east-central Española Basin may also have relevance to the Ancha Formation, since both were emplaced in the same general depositional environment. Cavazza (1986) subdivided Tesuque Formation strata in this part of the basin into two lithosomes based on composition and paleocurrent data. Sediment associated with lithosome A was derived from the granite-dominated Sangre de Cristo Mountains, and deposited by westward-flowing streams. The Ancha Formation would fall under lithosome A in the nomenclature of Cavazza (1986). Past studies (Smith, 2000, and Smith and Kuhle, 2001, in particular, but also Koning and Maldonado, 2001, Koning, 2002, Koning et al., 2002a, and Koning, 2003) have interpreted an alluvial slope depositional environment for lithosome A of Cavazza (1986). Because of their similarities, information regarding textural and hydrogeologic characteristics of alluvial slopes, such as those derived from study of lithosome A of the Tesuque Formation (Gaud, 2002), would be applicable to much of the Ancha Formation as well. Lithosome B of Cavazza (1986) is not readily applicable to the Ancha Formation because it was deposited by a relatively large, south-southwest flowing river sourced in Paleozoic sedimentary strata and Proterozoic quartzite rocks. It is also worth noting that other lithosomes have been proposed for the Tesuque Formation: lithosome S for ancestral Santa Fe River deposits of

Oligocene-Miocene age (Koning et al., 2004a), in addition to lithosome E for volcanic-bearing sediment derived from erosion of the Espinazo Formation (Daniel Koning, unpublished data).

THICKNESS

The Ancha Formation generally ranges from 10 to 90 m thick in the Santa Fe embayment based on geologic map, drill-hole, and seismic data (Koning et al., 2002b; S. Biehler, personal commun., 1999). Locally it may be slightly thicker than 90 m, such as where it buries paleovalleys developed on the upper Tesuque Formation contact. The Ancha Formation is 19-26 m thick on the western edge of the Caja del Rio Plateau north of the Santa Fe River, where it is appreciably thinner than in the center of the Santa Fe embayment to the east.

AGE

The base of the Ancha Formation is diachronous, and ranges from 2.7-3.5 Ma (?) in the western Santa Fe embayment to ~ 1.6 Ma in the eastern embayment near the Sangre de Cristo Mountains (Koning et al., 2002b). Sedimentation near the Pliocene-Pleistocene boundary appears to have been concentrated in the eastern embayment, and may reflect changes in discharge and sediment supply at that time.

ANCHA FORMATION DEPOSITS AND HYDROGEOLOGIC IMPLICATIONS

The Ancha Formation was deposited on a streamflow-dominated piedmont (i.e., alluvial slope) in the Santa Fe embayment; most of these streams were probably ephemeral. In and southwest of Santa Fe, much of the Ancha Formation corresponds to a fluvial facies of a Plio-Pleistocene Santa Fe River. Although this ancestral Santa Fe River was also deposited on an alluvial slope, it drained an appreciably larger area than streams associated with the alluvial slope to the south or north. Consequently, it was likely perennial and had relatively high flow energy.

South of the ancestral Santa Fe River, the alluvial slope deposits consist of coarse-grained channel facies interbedded with noticeably finer-grained sediment herein called extra-channel sediment, following Koning (2003) and Koning et al. (2004b). The proportion of extra-channel sediment significantly diminishes as one moves towards the mountain front, which probably imparts trending anisotropy to this hydrostratigraphic unit. These particular piedmont deposits

are similar to alluvial slope sediment described in the Tesuque Formation by Smith (2000). Alluvial slope deposits generally lack the tabular, planar-bedded couplets of relatively coarse- and fine-grained sediment diagnostic of sheetflood deposits, which are characteristic of alluvial fans (Bull, 1972; Blair, 1987 and 1999; and Blair and McPherson, 1994). The preserved constructional surface of the alluvial slope sediment away from the mountain front is relatively flat, and constitutes the top of the Ancha Formation. This surface has been designated as the Plains surface by Spiegel and Baldwin (1963).

Within about 3 km west of the mountain front south of the ancestral Santa Fe River, gravelly sediment comprises greater than a third of the estimated sediment volume, and lobate, fan-like geomorphic features are recognized on the present land surface. This sediment is dominated by very thin to medium, lenticular beds that probably reflect amalgamated erosionally truncated channels. Locally, about 10-20% of the section consists of poorly sorted deposits interpreted as being laid down as hyperconcentrated flows (including debris flows). Possible sheetflood deposits are very minor (trace to 2%); however, limited outcrop width makes it difficult to verify whether planar, very thin beds of alternating coarse-fine, sand-gravel actually fill broad channels, rather than being unconfined sheetflood deposits. This area adjacent to the mountain front may perhaps be thought of as a gradation between alluvial slope to alluvial fan deposits, but for the purposes of this report we treat them as alluvial slope deposits.

The Ancha Formation north of the ancestral Santa Fe River was deposited on a gently sloping (~1-2 degrees) alluvial slope on the western flank of the Santa Fe uplands. Drainages and their associated channels on this alluvial slope were relatively small, as were their corresponding feeder canyons in the Santa Fe uplands. Eolian sedimentation was significant here.

The interpreted depositional environment has important hydrogeologic implications. Channel avulsion on the alluvial slope south of the Santa Fe River was probably common, and the resulting distribution of channels in an outcrop appears random. The ancestral Santa Fe River probably did not significantly meander, but it may have been more likely to shift back and forth in a continuous fashion because its coarse-grained channels are more laterally continuous and connected than in alluvial slope sediment to the south.

Below, we describe the alluvial slope and ancestral Santa Fe River deposits in detail. Although technically the ancestral Santa Fe River sediments were deposited on an alluvial slope, we will distinguish them separately from alluvial slope deposits north and south of this river.

Stratigraphic sections representing these two deposits, and various textural subdivisions within them, are presented in Appendix 1. These stratigraphic sections were measured and described by Daniel Koning except for the Galisteo 2 section, which was described by Sean Connell and Frank Pazzaglia (unpublished data). Section locations are shown on Plate 3 of this report.

Alluvial Slope Deposits (Map Unit QTaas)

General description of alluvial slope deposits south of the ancestral Santa Fe River. The alluvial slope deposits of the Ancha Formation south of the ancestral Santa Fe River can be subdivided into extra-channel and channel facies. The extra-channel facies is characterized by a poorly to moderately sorted, silty or muddy sand (mostly very fine- to medium-grained sand with subordinate coarse- to very coarse-grained sand) that contains minor, scattered pebbles. Some of the poorly sorted beds with scattered pebbles may represent hyperconcentrated flow deposits. The sediment is well-consolidated and commonly weakly cemented by calcium carbonate. The beds in this facies are medium to thick, tabular to broadly lenticular, and internally massive or bioturbated. Scattered very thin to thin lenses of coarse sand and pebbles may be present in sparse quantities. The sediment of this facies is interpreted to have been deposited in very broad channels or swales, or as small depositional lobes on an alluvial slope.

The other sedimentary facies consists of channel deposits of gravelly sand, sandy gravel, and medium- to very coarse-grained sand. These coarse channels are lenticular to broadly lenticular in form and commonly medium to thick. The channels are generally 2-30 m in width. Within a channel, the sand is commonly planar-laminated and the pebbles are in very thin to medium, lenticular beds; local tangential cross-stratification or trough-cross-stratification is present but generally less than 50 cm thick. Gravel is clast-supported, moderately to poorly sorted, and subrounded (more subangular towards the mountain front). Clasts consist of granite, foliated granite, and granitic gneiss with subordinate (3-15%) amphibolite or amphibolite-gneiss; locally,

there are trace intermediate-felsic volcanic rocks presumably derived from reworking of the Espinazo Formation or Bishops Lodge Member of the Tesuque Formation. Quartzite clasts comprise less than 2% of the total gravel fraction south of Interstate 25, but become more abundant (1-16%) between the Galisteo River and Gallina Arroyo. Channel sediment is generally loose to weakly consolidated and non-cemented. However, locally there is moderate to strong cementation, especially at the base of the Ancha Formation where it overlies the Galisteo Formation and Mesozoic strata.

The Plains surface contains compound soils that locally exhibit <25 cm-thick, clay-rich Bt or Btk horizons underlain by 50 to >100 cm-thick calcic and siliceous Bk or Bkq horizons with stage II to III+ pedogenic carbonate morphology (Koning et al., 2002b). Below the soils associated with the Plains surface, buried soils are not common. Where exposed, these buried intraformational paleosols are characterized by clay-rich Bt horizon(s) overlying paler-colored calcic horizon(s) with stage II to III pedogenic carbonate morphology (Koning et al., 2002b).

General description of alluvial slope sediment between the ancestral Santa Fe River and upper Cañada Ancha.

The Ancha Formation alluvial slope deposits between the ancestral Santa Fe River and upper Cañada Ancha, in particular its tributary of Alamo Creek, are generally very pale brown to light yellowish brown, silty, very fine- to medium-grained sand. They contain 1-20% channel deposits of medium- to very coarse-sand and sandy gravel. The channels are generally in very thin to medium, lenticular beds and only 1-5 m wide, but locally are as much as 2 m thick and 30 m wide. The relative abundance of these coarse channel deposits increases towards the east. The gravel in the channels are pebble- to fine-cobble in size, poorly sorted, subrounded to subangular, and consist of granite with 10-35% quartzite, 0.5-1% chert, 1-15% Paleozoic clasts of limestone and siltstone, 2-5% amphibolite, and 3% micaceous gneiss and schist. The composition of the gravel, combined with the relative small size of the channels, indicates the sediment was locally reworked from the Tesuque Formation in the Santa Fe uplands located north of Santa Fe. Most of the silty very fine- to medium-grained sand is likely eolian in origin but may have been reworked by local slopewash processes. Buried soils are locally present and may be vertically spaced on the scale of 1-5 m.

The fine-grained sediment is interbedded with subordinate, very thin to thick beds of phreatomagmatic deposits; these generally consist of medium- to very coarse-grained sand (basalt with various proportions, but generally <50%, of arkosic sand) and minor very fine to medium basaltic pebbles. Phreatomagmatic deposits are found throughout the Ancha Formation in the northern Santa Fe embayment, but are most abundant northwest of Arroyo Calabasa (Ralph Shroba, 2004, written communication). These are partly indurated (with only minor effervescence in dilute hydrochloric acid) or well consolidated.

Hydrogeologic implications of features in alluvial slope deposits. The extra-channel facies of the Ancha Formation probably has lower hydraulic conductivity values compared to the coarse channel facies. This inference is based on comparison to air-permeameter measurements on lithosome A extra-channel facies in the Tesuque Formation (John Sigda, 2003, unpublished data). The extra-channel sediment is generally poorly to moderately sorted, with appreciable amounts of silt or mud in the interstices of the sand grains. These fines would be expected to reduce porosity. Strong cementation observed in coarse channels of the Ancha Formation near its base in the southern part of the Santa Fe embayment (see below discussion of the basal contact for this part of the embayment) may dramatically reduce the hydraulic conductivity. Relatively indurated and well consolidated phreatomagmatic deposits, which are particularly common north of the Santa Fe River, may potentially influence unsaturated flow in the Ancha Formation in this area.

Textural-based subdivisions. The Ancha Formation varies considerably in texture, with the general trend of becoming finer to the west (Koning et al., 2002b). We have subdivided the alluvial slope deposits into four units based on the proportions of coarse-grained channels to finer-grained extra-channel sediment. The four textural units corresponding to this subdivision are shown on the compilation map. The boundaries between the units are very gradational in a lateral sense (3 to 4 km in width). We approximated the unit boundaries based on inspection of outcrops and available records of well cuttings. It should be noted that the outcrops south of Interstate 25 are small and of poor quality. Here, most outcrops of the Ancha Formation expose only 1 to 2 m below the ground surface, an interval which has incorporated much eolian silt and fine sand during Pleistocene time. This eolian sediment commonly makes the Ancha Formation

appear finer-grained near the surface than it actually is at depth. Consequently, in this textural subdivision we place little weight on small outcrops at the modern surface, especially when they overlie well-developed calcic soils. The four textural subdivisions are described below, in decreasing order of inferred groundwater-resource potential, as well as inferences regarding hydrogeologic characteristics:

>35% coarse channels: More than 35% of this unit contains coarse-grained channel deposits probably associated with ephemeral avulsing streams. Channels are probably interconnected, and there is likely to be greater than 50% coarse channels in most places. The inferred high degree of channel interconnection could make this unit a potentially useful aquifer zone, with relatively high overall hydraulic conductivity. However, most of this textural subdivision is not within the saturated zone.

25-60% coarse channels: This unit is differentiated only in the southernmost Santa Fe embayment, and contains about 25-60% coarse-grained channels. Most channels are probably interconnected, and so the overall hydraulic conductivity of the unit would be inferred to be moderate to high. However, this textural subdivision is generally not located within the saturated portion of the Ancha Formation.

15-60% coarse channels: This unit contains 15-60% coarse-grained channels scattered in finer extra-channel deposits. Some to most coarse channels are interconnected. Overall hydraulic conductivity of the unit is inferred to be moderate to high. Aquifer transmissivity measured in wells completed across this unit and into the underlying Tesuque or Espinaso Formations ranges from 70 to 980 ft²/day.

1-30% coarse channels: Channels are probably not significantly interconnected. Overall hydraulic conductivity of the unit is generally inferred to be relatively low, but locally may be moderate. Aquifer transmissivity in wells completed across this unit and into the underlying Espinaso Formation ranges from 15 to 280 ft²/day.

Ancestral Santa Fe River Deposits (Map Unit QTasr)

General description. Sediment deposited by an ancestral Santa Fe River during Pliocene and early Pliostocene times contains sandy gravel in thin to thick, lenticular to broadly lenticular to channel-shaped beds; locally, there is planar- or tangential- cross-stratification up to 90 cm thick. The gravel is generally clast-supported and consists of pebbles with 30 to 50% cobbles; clasts are subrounded, poorly sorted, and composed of granite with 1 to 6% quartzite and 1 to 3% amphibolite. There are minor beds of silty or clayey very fine to fine sand that correspond to floodplain sediment, plus minor beds of extra-channel sediment (muddy very fine to very coarse sand with 1 to 15% pebbles) similar to that seen in alluvial slope deposits to the south. Extra-channel sediment become more abundant (30 to 50% of sediment volume) near the southern and northern margins of the ancestral Santa Fe River deposits, likely because of interfingering with alluvial slope deposits. The sediment is weakly to moderately consolidated, and not cemented.

Hydrogeologic implications of features in the ancestral Santa Fe River deposits. The overall coarseness of this deposit, in particular the predominance of sandy gravel channel deposits, is strongly suggestive of relatively high hydraulic conductivity values. In general, the floodplain beds do not appear sufficiently laterally extensive as to form confined conditions within the unit. However, locally there may be laterally extensive floodplain beds which could act as aquitards to groundwater flow. The aquifer transmissivity measured in wells completed in this unit ranges from 470 to 7750 ft²/day, and reflects the highest transmissivity values measured in the Southern Española Basin.

Rate of Clast Size Diminishment Away From the Mountain Front

Because grain size is known to relate to hydraulic conductivity (Hazen, 1911), a survey of gravel clast size was conducted across the Santa Fe embayment from the mountain-front to the western edge of the Caja del Rio Plateau. Clast diameters in the ancestral Santa Fe River deposits were generally measured from vertical outcrop faces, but many measurements in the alluvial slope deposits to the south were taken from clasts lying on the surface that did not show evidence of recent spallation. At a given locality, the longest and intermediate axes of the 5 to 7 largest clasts were recorded and an arithmetic mean was calculated. The mean intermediate (b) axis of gravel is plotted against distance from the mountain front (Figure 1). On the graph, the data are

differentiated into two groups that correspond to the ancestral Santa Fe River and alluvial slope deposits south of the Santa Fe River.

The data plotted in Figure 1 show an appreciable clast size decrease westward away from the mountain front, and there is more variation in the average maximum clast sizes in the alluvial slope deposits compared to the ancestral Santa Fe River deposits of the Ancha Formation. This figure allows some inferences to be made concerning the competency of the drainages that deposited the alluvial slope sediment versus that of the ancestral Santa Fe River. Stream competency is the largest particle that a stream can move under a given set of hydraulic conditions (Ritter, 1986). The data in Figure 1, in particular the difference in slopes of the two data sets, indicates that the Santa Fe River was able to maintain higher competency away from the mountain front than the smaller alluvial slope drainages. Buried clast sizes revealed in good exposures representing both data sets mimic this trend, and so it is difficult to argue that the two slopes in Figure 1 are primarily due to weathering effects such as spallation.

The westward decrease in stream competence indicated in Figure 1 may be due to reduction in original channel depth (lessening the shear stress on a given clast) or reduction in depositional slope in that direction. Loss of discharge due to evaporation or infiltration is also a factor in the westward decrease of the maximum clast sizes. The lesser clast size range and lower regression line slope of the ancestral Santa Fe River data may possibly be related to the more perennial stream discharge in that area. The riparian vegetation associated with perennial streams would likely create more resistance to bank erosion (Smith, 1976), and lead to higher channel depth to width ratios. One would also expect less infiltration under a perennial stream due to relatively constant saturated conditions. Thus, channels associated with the ancestral Santa Fe River may have maintained their depths and a given stream discharge volume transferred farther westward from the mountain front, creating a more constant flow regime and competence. However, possible weathering and spallation effects on surface clasts, in addition to the more numerous measurements, might also contribute to the larger spread of clast size data in alluvial slope deposits south of the ancestral Santa Fe River.

BASAL CONTACT OF THE ANCHA FORMATION

Description

Exposures of the basal contact of the Ancha Formation were described as part of this study. All exposures show an angular unconformity at this contact. Because there are no exposures in the center of the Santa Fe embayment, we cannot be absolutely certain if the basal contact here represents an angular unconformity, disconformity, or conformable transition with the underlying Tesuque Formation. Because of a 5 to 6 Myr age difference between the Tesuque and Ancha Formations along Cañada Ancha (Koning et al., 2002b) it is unlikely that the Ancha-Tesuque Formation is conformable. The Ancha Formation is generally undeformed where exposed, whereas the underlying Tesuque Formation has been mildly to moderately tilted. Thus, it is likely that this unconformity is angular throughout the Santa Fe embayment. In the following, we describe and discuss the basal Ancha Formation contact at various places in the Santa Fe embayment.

Northeastern Santa Fe embayment. The Ancha Formation (both the alluvial slope and ancestral Santa Fe River deposits) overlies the Tesuque Formation with an angular unconformity. The contact is well-exposed in numerous locations in the eastern Santa Fe embayment between the Santa Fe River and Arroyo Hondo. At these localities, there is no cementation or even significant discoloration associated with the Ancha Formation contact. Moreover, the basal contact here corresponds to a scour surface with meter-scale relief, and no noticeable soils were observed in the Tesuque Formation immediately below the contact. Photos and descriptions of the contact and overlying Ancha Fm are shown in Figures 2 and 3.

Northern and northwestern Santa Fe embayment. The base of the Ancha Formation is readily observed on the western edge of the Caja del Rio escarpment, under the basalt flows of the Cerros del Rio volcanic field. Here, the basal contact is an angular unconformity, with the Ancha Formation overlying a distinctive volcanoclastic unit of the Tesuque Formation (lithosome E) whose gravel fraction consists of reworked detritus from the Espinazo Formation. No cementation or discoloration was observed at the contact, nor was there a soil developed in the uppermost Tesuque Formation.

Under the flows of the Cerros del Rio volcanic field near the big bend of Cañada Ancha (near Portales Pond), the Ancha Formation overlies lithosome S and lithosome A of the Tesuque Formation with angular unconformity. Because the Tesuque Formation is only shallowly dipping (less than 6 degrees), this apparent dip discrepancy is generally not obvious in the small exposures of the area. At one locality in the “big bend” (site WA-88; UTM coordinates: 3951930 N, 402025 E, zone 13; NAD 27) the uppermost 60 cm of Tesuque Formation has a calcic soil with a stage III pedogenic carbonate morphology. East of the latter location, the contact is not well-exposed. Here, there may possibly be a soil in the Tesuque Formation immediately below the contact, as seen in WA-88, but there is nothing to suggest development of a petrocalcic horizon with stage IV carbonate morphology.

North of the prominent eastward bend of Cañada Ancha, the Ancha Formation overlies the coarse upper unit of the Tesuque Formation (lithosome A). Because of the low dips of the Tesuque Fm (1 to 3 degrees), there is no obvious discrepancy in bedding dips between the Ancha and Tesuque Formations on either side of the unconformity.

In the modern Santa Fe River west of Agua Fria, the base of the Ancha Formation is well-exposed. Here, one can observe the basal Ancha contact from northeast to southwest, where a change from locally reworked Tesuque Formation material in the Ancha Formation to ancestral Santa Fe River gravel is well exposed (Figures 4 and 5). In the modern river bed, there is a prominent angular unconformity at the base of the Ancha Formation. However, there is no cementation or discoloration associated with the contact. Furthermore, there is no significant soil development at the top of the underlying Tesuque Formation (lithosome S).

North of the Santa Fe River and south of Cañada Ancha, the relatively fine Ancha Formation overlies lithosome S of the underlying Tesuque Formation in an angular unconformity. No cementation, discoloration, or soils were observed at the contact in this area.

La Cienega area. Although the basal Ancha contact projects to the ground surface at La Cienega, poor exposures precluded an evaluation of this boundary. In the lower part of the

Ancha Formation described in the La Cienega stratigraphic section (Appendix 1), much of the lower units are weakly cemented by CaCO_3 , and unit 4 has 5% very thin to thin, wavy beds of strongly cemented very fine to fine sand. The lower Ancha Formation may be somewhat more cemented here than to the east, perhaps due to upwelling of groundwater east of the structurally high, and relatively impermeable, Espinazo Formation and Oligocene intrusions at La Cienega.

Bonanza Hill and Turquoise Hill area. In cuttings logs for water wells 1 to 4 km east of Bonanza Hill and Turquoise Hill, there is a 3 to 9 m thick interval at the base of the Ancha Formation that is cemented by calcium carbonate (commonly called “caliche” or “hard white shale” in the driller logs). Such beds of dense to powdery calcium carbonate have been noted locally on the surface in the area (e.g., Spiegel and Baldwin, 1963; Koning and Hallet, 2000, for the Plio-Pleistocene sediment re-designated as the Tuerto Formation). It was noted that these exposed “limy” deposits were developed in the Ancha Formation immediately above its contact with underlying, significantly less permeable rocks (Spiegel and Baldwin, 1963). This also appears to be the case for the subsurface basal Ancha Formation 1 to 4 km east of Bonanza Hill and Turquoise Hill, where the Ancha Formation overlies the Galisteo and Espinazo Formations (Koning and Hallett, 2000). This calcium carbonate may have accumulated in the distalmost part of the Ancha Formation early in its deposition, as this sediment on-lapped onto paleo-topographic highs of the La Bajada fault footwall. The actual process of calcium carbonate enrichment may be due to evaporation of shallow groundwater mounding up against these lower-permeability rocks.

Southern Santa Fe embayment. Where the Ancha Formation overlies the Galisteo Formation, there locally is a 1.5 to 2 m-thick zone at the base of the Ancha Formation, with induration of at least some of the beds (Figure 6). This degree of strong cementation was not seen where the Ancha Formation overlies the Espinazo Formation (see section SEO-A87 in Appendix 1), although only a few exposures were accessible for observation. In the latter there were local zones of weak to moderate cementation in the lower units of the Ancha Formation (manifested as white bands); however, the cementation did not result in induration. Eastward towards Lamy, where the Ancha Formation overlies Mesozoic strata, there is significant cementation of the lower 4 to 10 m of the Ancha Formation (see Galisteo 02 section of Appendix 1; Figure 7). Near

the syncline axis in the southern part of the Santa Fe embayment, where the Ancha Formation is interpreted to overlie the Tesuque Formation, driller's cuttings logs generally do not mention a cemented zone at the inferred base of the Ancha Formation.

Hydrogeologic Implications

The basal contact of the Ancha Formation corresponds with an angular unconformity in all observed outcrops, and could potentially affect vertical groundwater flow between the Ancha and underlying Tesuque Formation, particularly where there is also a grain size distinction between the two formations. Anisotropy related to sedimentary stratification has been shown to decrease vertical hydraulic conductivity up to an order of magnitude relative to horizontal hydraulic conductivity (Domenico and Schwartz, 1990; Davis, 1969). The effects of an unconformable contact could presumably be even more pronounced as the weathering and pedogenic activity that is likely to occur on an exposed surface over geologic time may lead to an increase in fines or cementation just below the exposed surface (Birkeland, 1999). After burial, any increase in fines and development of cementation associated with the exposed surface would likely reduce vertical hydraulic conductivity across the basal contact. Noting such features as weathering and pedogenic activity, or changes in the dip of underlying formations, along the basal contact of the Ancha Formation could have important implications for assessing groundwater flow between this and underlying hydrogeologic units.

The strong degree of cementation of the basal Ancha Formation in much of the southeast and southwest parts of the Santa Fe embayment appears to coincide with localities where the Ancha Formation overlies Mesozoic strata, the Galisteo Formation, or locally, the Espinazo Formation. Such strong cementation would likely retard vertical water flow through this basal zone. Here one might expect perched aquifers in the Ancha Formation, or local confined conditions in underlying hydrostratigraphic units. This cementation could also impede downward flow into underlying rock units. Immediately east of La Cienega, where the Ancha Formation overlies the Tesuque Formation, the base of the Ancha Formation may possibly be somewhat more cemented than closer to the mountain front.

In the northwest and northeast parts of the Santa Fe embayment, in particular where the Ancha Formation overlies the Tesuque Formation north of La Cienega Creek, there is no evidence of a significant barrier to groundwater flow at the base of the Ancha Formation. Likewise, in the southern embayment near the syncline axis there is no evidence of significant cementation where the Ancha Formation overlies the Tesuque Formation. It is very likely that vertical flow between the Ancha Formation and the underlying hydrogeologic units is significantly greater where Ancha overlies the Tesuque Formation than in the southern part of the embayment where the Ancha Formation overlies pre-Tesuque hydrogeologic units.

ACKNOWLEDGMENTS

This report was improved greatly by reviews and comments of Sean Connell of the New Mexico Bureau of Geology and Mineral Resources, and of Ralph Shroba of the U.S. Geological Survey.

REFERENCES

- Bachman, G.O., 1975, Geologic map of the Madrid quadrangle, Santa Fe and Sandoval Counties, New Mexico: U.S. Geological Survey, Geologic Quadrangle Map GQ-1268, scale 1:62,500.
- Birkeland, P.W., 1999, Soils and geomorphology: New York, Oxford University Press, 430 p.
- Birkeland, P.W., Machette, M.N., and Haller, K.M., 1991, Soils as a tool for applied Quaternary geology: Utah Geological and Mineral Survey, miscellaneous publication 91-3, 63 p.
- Blair, T.C., 1987, Sedimentary processes, vertical stratification sequences, and geomorphology of the Roaring River alluvial fan, Rocky Mountain National Park: *Journal of Sedimentary Research*, v. 57, p. 845-862.
- Blair, T.C., 1999, Sedimentary processes and facies of the waterlaid Anvil Springs Canyon alluvial fan, Death Valley, California: *Sedimentology*, v. 46, p. 913-940.
- Blair, T.C., and McPherson, J.G., 1994, Alluvial fans and their natural distinction from rivers based on morphology, hydraulic processes, sedimentary processes, and facies assemblages: *Journal of Sedimentary Research*, v. A64, p. 450-489.

- Booth, F.O., III, 1977, Geologic map of Galisteo Creek area, Lamy to Cañoncito, Santa Fe County, New Mexico: U.S. Geological Survey, Miscellaneous Field Studies Map MF-0823, scale 1:12,000.
- Bull, W.B., 1972, Recognition of alluvial-fan deposits in the stratigraphic record, *in* Rigby, J.K., and Hamblin, W.K., ed., Recognition of Ancient Sedimentary Environments: Soc. Econ. Paleont. Mineral. Spec. Publ., v. 16, p. 63-83.
- Cavazza, W., 1986, Miocene sediment dispersal in the central Española Basin, Rio Grande rift, New Mexico, USA: Sedimentary Geology, v. 51, p. 119-135.
- Davis, S. N., 1969, Porosity and permeability of natural materials, *in* R.J.M. DeWiest, ed., Flow Through Porous Materials: Academic Press, NY, p. 54-89.
- Davis, M.J., Lohmann, R.C., Phillips, F.M., Wilson, J.L., and Love, D.W., 1993, Architecture of the Sierra Ladrones Formation, central New Mexico – depositional controls on the permeability correlation structure: Geological Society of America Bulletin, v. 105, p. 998-1007.
- Domenico, P.A. and Schwartz, F.W., 1990, Physical and Chemical Hydrogeology: John Wiley & Sons, 824 p.
- Dreyer, T., Falt, L.M., Hoy, T., Knarud, R., Steel, R., and Cuevas, J.L., 1993, Sedimentary architecture of field analogues for reservoir information (SAFARI) – a case study of the fluvial Escanilla formation, Spanish Pyrenees: International Association of Sedimentologists, Special Publication 15, p. 57-80.
- Fetter, C.W., 1988, Applied hydrogeology: Don Mills, Ontario, Collier Macmillian Canada, Inc., 592 p.
- Fogg, G.E., 1986, Groundwater flow and sand body interconnectedness in a thick multiple-aquifer system: Water Resources Research, v. 22, p. 679-694.
- Gaud, M.N., 2002, Outcrop investigations of the permeability and spatial distributions of alluvial-slope lithofacies near Española, New Mexico [M.S. thesis]: Albuquerque, New Mexico, University of New Mexico, 127 p.
- Hazen, A., 1911, Discussion: Dams on Sand Foundations. Transactions, American Society of Civil Engineers, v.73, 199 p.
- Johnson, R.B., 1975, Geologic map of the Galisteo quadrangle, Santa Fe County, New Mexico: U.S. Geological Survey, Geologic Quadrangle Map GQ-1234, scale 1:24000.

- Kelley, V.C., 1978, Geology of the Española Basin, New Mexico: New Mexico Bureau of Mines and Mineral Resources, Geologic Map GM-48, scale 1:125,000.
- Koning, D.J., 2002, Geologic map of the Española 7.5-minute quadrangle, Santa Fe County, New Mexico: New Mexico Bureau of Geology and Mineral Resources, Open-file Geologic Map OF-GM-54, scale 1:24,000.
- Koning, D.J., 2003, Geologic map of the Chimayo 7.5-minute quadrangle, Rio Arriba and Santa Fe counties, New Mexico: New Mexico Bureau of Geology and Mineral Resources, Open-file Geologic Map OF-GM-71, scale 1:24,000.
- Koning, D.J., and Hallett, R.B., 2000, rev. 2003, Geology of the Turquoise Hill 7.5-min. quadrangle, Santa Fe County, New Mexico: New Mexico Bureau of Mines and Mineral Resources, Open-file Geologic Map OF-GM 41, scale 1:24,000.
- Koning, D.J., and Maldonado, F., 2001, Geologic map of the Horcado Ranch quadrangle, Santa Fe County, New Mexico: New Mexico Bureau of Mines and Mineral Resources, Open-file Geologic Map OF-GM-54, scale 1:24,000.
- Koning, D.J., Nyman, M., Horning, R., Eppes, M., and Rogers, S., 2002a, Geology of the Cundiyo 7.5-min. quadrangle, Santa Fe County, New Mexico, New Mexico Bureau of Geology and Mineral Resources, Open-file Geologic Map OF-GM 56, scale 1:24,000.
- Koning, D.J., Connell, S.D., Pazzaglia, F.J., and McIntosh, W.C., 2002b, Redefinition of the Ancha Formation and Pliocene-Pleistocene deposition in the Santa Fe embayment, north-central New Mexico: *New Mexico Geology*, v. 24, no. 3, p. 75-87.
- Koning, D.J., Smith, G., Lyman, J., and Paul, P., 2004a, Lithosome S of the Tesuque Formation: hydrostratigraphic and tectonic implications of a newly delineated lithosome in the southern Española Basin, New Mexico [abstract]: Hudson, M.R., *ed.*, Geologic and Hydrogeologic Framework of the Española Basin -- Proceedings of the 3rd Annual Española Basin Workshop, Santa Fe, New Mexico, March 2-3, 2004, U.S. Geological Survey Open-File Report 2004-1093, p. 17.
- Koning, D.J., Smith, G.A., and Aby, S., 2004b, Third-day supplemental road log, Penasco to Espanola: New Mexico Geological Society, 55th Field Conference, Guidebook, 99-107.
- Lisenbee, A.L., 1999, Geology of the Galisteo 7.5-min. Quadrangle, Santa Fe County, New Mexico, New Mexico Bureau of Mines and Mineral Resources, Open-file Geologic Map OF-GM 30, scale 1:24,000.

- Manley, K., 1976, The late Cenozoic history of the Española basin, New Mexico [Ph.D. thesis]: Boulder, Colorado, University of Colorado, 171 p.
- Manley, K., 1979, Tertiary and Quaternary stratigraphy of the northeast plateau, Española Basin, New Mexico: New Mexico Geological Society, Guidebook 30, p. 231-236.
- Maynard, S.R., and Lisenbee, A.L., 2002, Geology of the Picture Rock 7.5-min. quadrangle, Santa Fe County, New Mexico, New Mexico Bureau of Geology and Mineral Resources, Open-file Geologic Map OF-GM 51, scale 1:12,000.
- Miller, J.P., Montgomery, A., and Sutherland, P.K., 1963, Geology of part of the southern Sangre de Cristo Mountains, New Mexico: New Mexico Bureau of Mines and Mineral Resources, Memoir 11, 106 p.
- Read, A.S., Rogers, J., Ralser, S., Ilg, B., Kelley, S., 1999, Geology of the Seton Village 7.5-min. quadrangle, Santa Fe County, New Mexico: New Mexico Bureau of Mines and Mineral Resources, Open-file Geologic Map OF-GM 23, scale 1:12,000.
- Read, A.S., Koning, D.J., Smith, G.A., Ralser, S., Rogers, J., Bauer, P.W., 2000, Geology of the Santa Fe 7.5-min. quadrangle, Santa Fe County, New Mexico, New Mexico Bureau of Mines and Mineral Resources, Open-file Geologic Map OF-GM 32, scale 1:12,000.
- Ritter, D.F., 1986, Process Geomorphology: Dubuque, Iowa, Wm. C. Brown Publishers, 579 p.
- Shepard, R.G., 1989, Correlations of Permeability and Grain Size, Ground Water, v. 27, no. 5, p. 633-638.
- Sigda, J., 2003, Preliminary hydraulic properties of Tesuque Formation sediments, Buckman area, White Rock quad, poster presentation at the second annual Española Basin Technical Advisory Group workshop, Santa Fe, NM, 4-5 March 2003.
- Smith, D.G., 1976, Effect of vegetation on lateral migration of anastomosed channels of a glacial meltwater river: Bulletin of the Geological Society of America, v. 86, p. 27-43.
- Smith, G.A., 2000, Recognition and significance of streamflow-dominated piedmont facies in extensional basins: Basin Research: v. 12, p. 399-411.
- Smith, G.A., and Kuhle, A.J., 2001, Alluvial-slope deposition of the Skull Ridge Member of the Tesuque Formation, Española Basin, New Mexico: New Mexico Geology, v. 23, no. 2, p. 30-37.
- Spiegel, Z., and Baldwin, B., 1963, Geology and water resources of the Santa Fe area, New Mexico: U.S. Geological Survey Water-Supply Paper 1525, 258 p.

Tyler, N., and Finley, R.J., 1991, Architectural controls on the recovery of hydrocarbons from sandstone reservoirs, *in* Miall, A.D., and Tyler, N. (eds.), The three-dimensional facies architecture of terrigenous clastic sediments and its implications for hydrocarbon discovery and recovery: Society for Sedimentary Geology (SEPM), Concepts in Sedimentology and Paleontology, v. 3, p. 1-5.

FIGURES

Figure 1

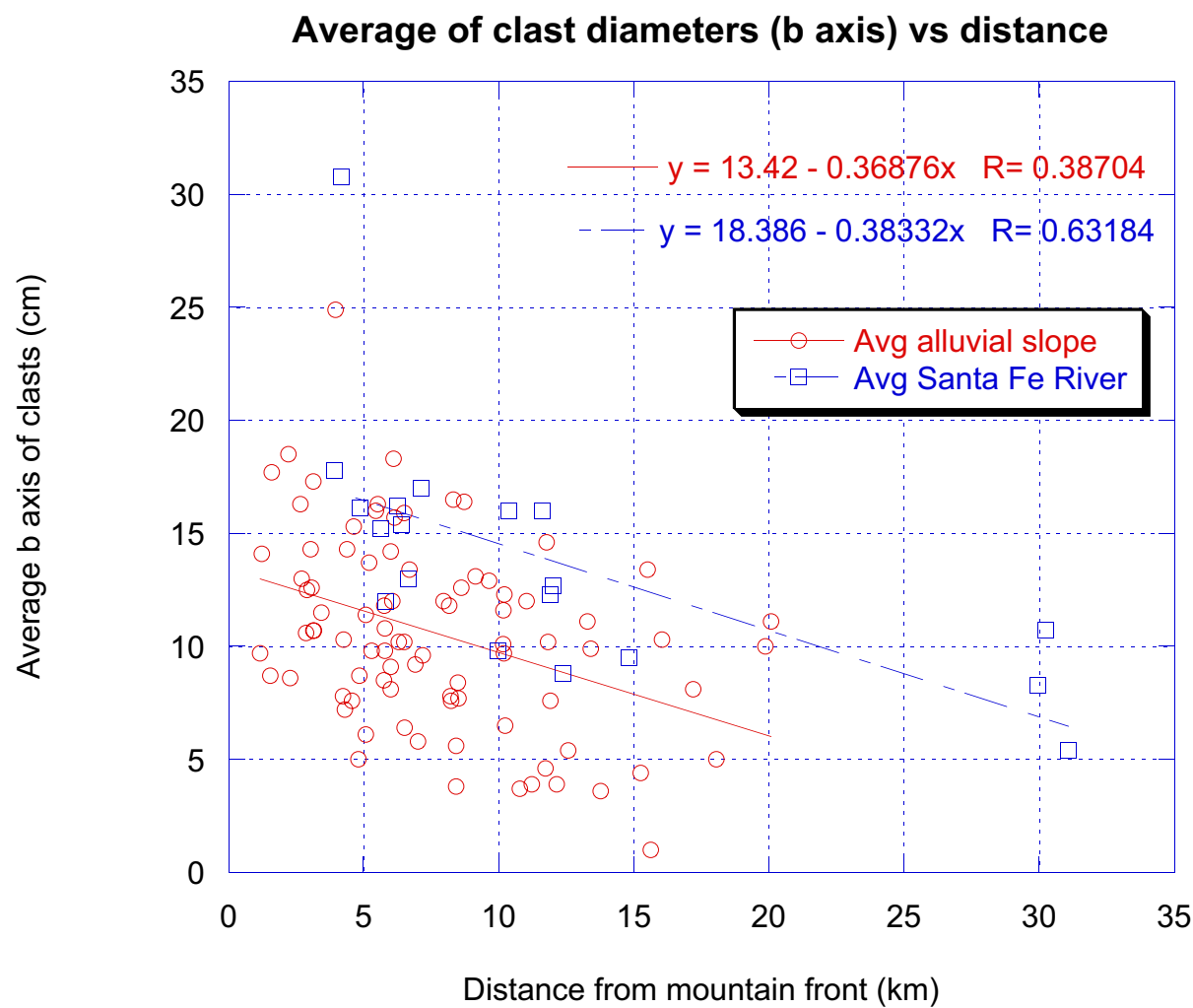




Figure 2. Brownish sandy gravel of the Ancha Formation (12 m thick) overlying reddish lithosome S of the Tesuque Formation (4.5 m thick); head of rock hammer is on the contact. Note the lack of discoloration, cementation, or paleosols at the contact. Photo taken at site SEO-A3 in Arroyo de los Chamisos, about 0.4 km east-northeast of Santa Fe High School. UTM coordinates: 3945309N, 412777E (NAD 27, zone 13).



Figure 3. Close-up view of the Ancha Formation in Figure 2. The Ancha Formation here is a sandy gravel interpreted to have been deposited by an ancestral Santa Fe River (unit QTasr). Beds are vague, thin to medium, and broadly lenticular. The sediment is weakly to moderately consolidated and not significantly cemented.

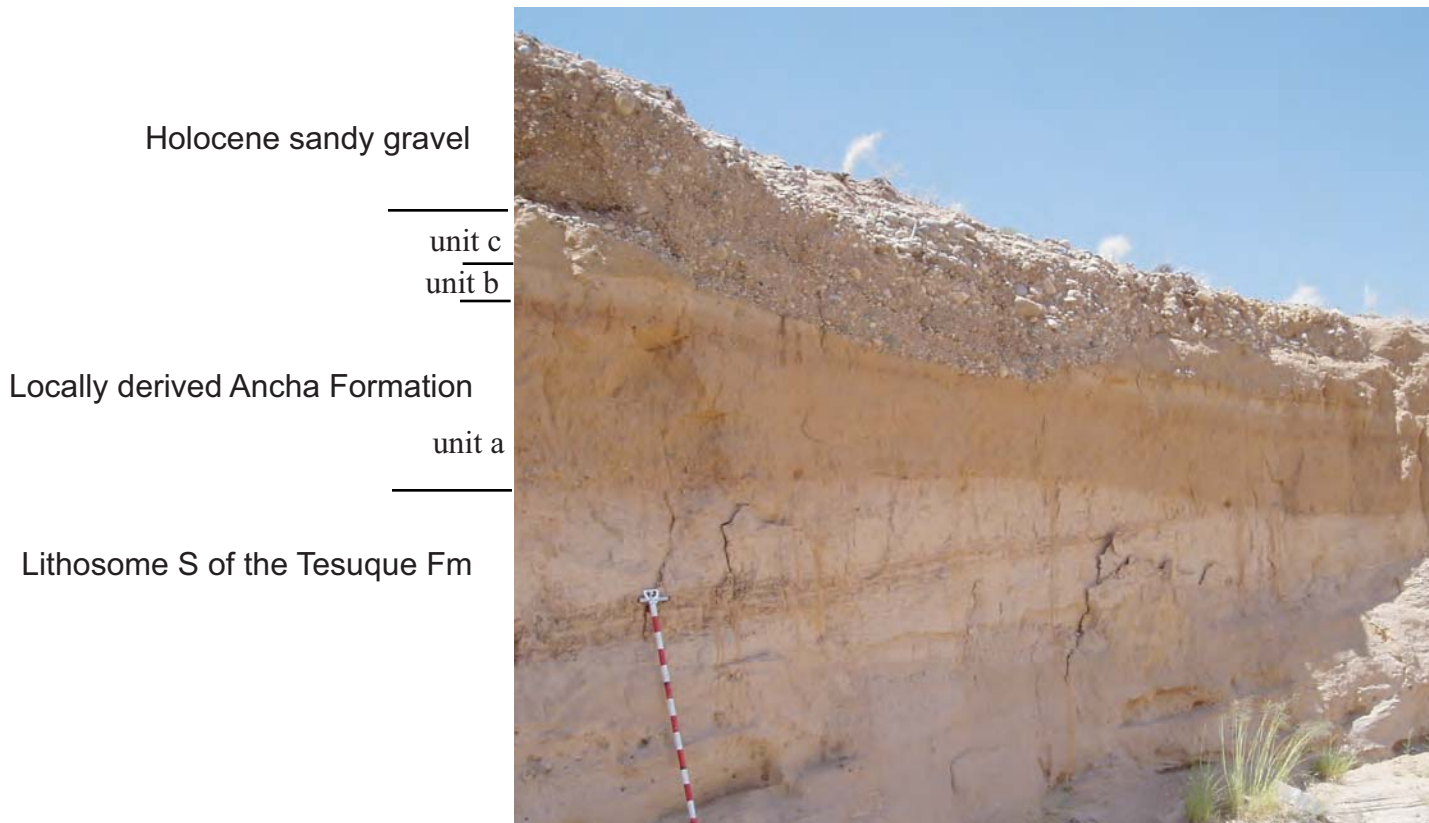


Figure 4. Angular unconformity under the Ancha Formation at site SEO-A76. The Ancha Formation here consists of three units, from base to top: a) 95 cm of a fining-upward, light brown, slightly clayey, very fine to very coarse sand with 3-5% very fine to fine pebbles; lowermost 10 cm of unit a consists of very fine to very coarse pebbles and minor fine cobbles; a reddish Bt soil horizon has formed on upper 11-16 cm of the unit; b) 35 cm of light yellowish brown, very fine to very coarse sand with 7-10% very fine to coarse pebbles; c) 0-95 cm of yellowish brown, clayey-muddy, very fine to very coarse sand (mostly mU-cU) that is in vague, thin to medium, lenticular to broadly lenticular beds. Clast count of basal gravel (unit a, n=106) gives: 39% granite, 34% quartzite, 9% vein quartz, 8% cherty quartzite, 5% altered, epidotized-chloritized granite, and 1% muscovite. This clast composition is very similar to the underlying Tesuque Fm and indicates that the Ancha Fm here was locally reworked from the Tesuque Fm. Site is located 260-270 m southeast of the County Road 62 bridge over the modern Santa Fe River (UTM coordinates: 3946077N, 0407098E; NAD 27, zone 13).

SOUTHWEST

NORTHEAST

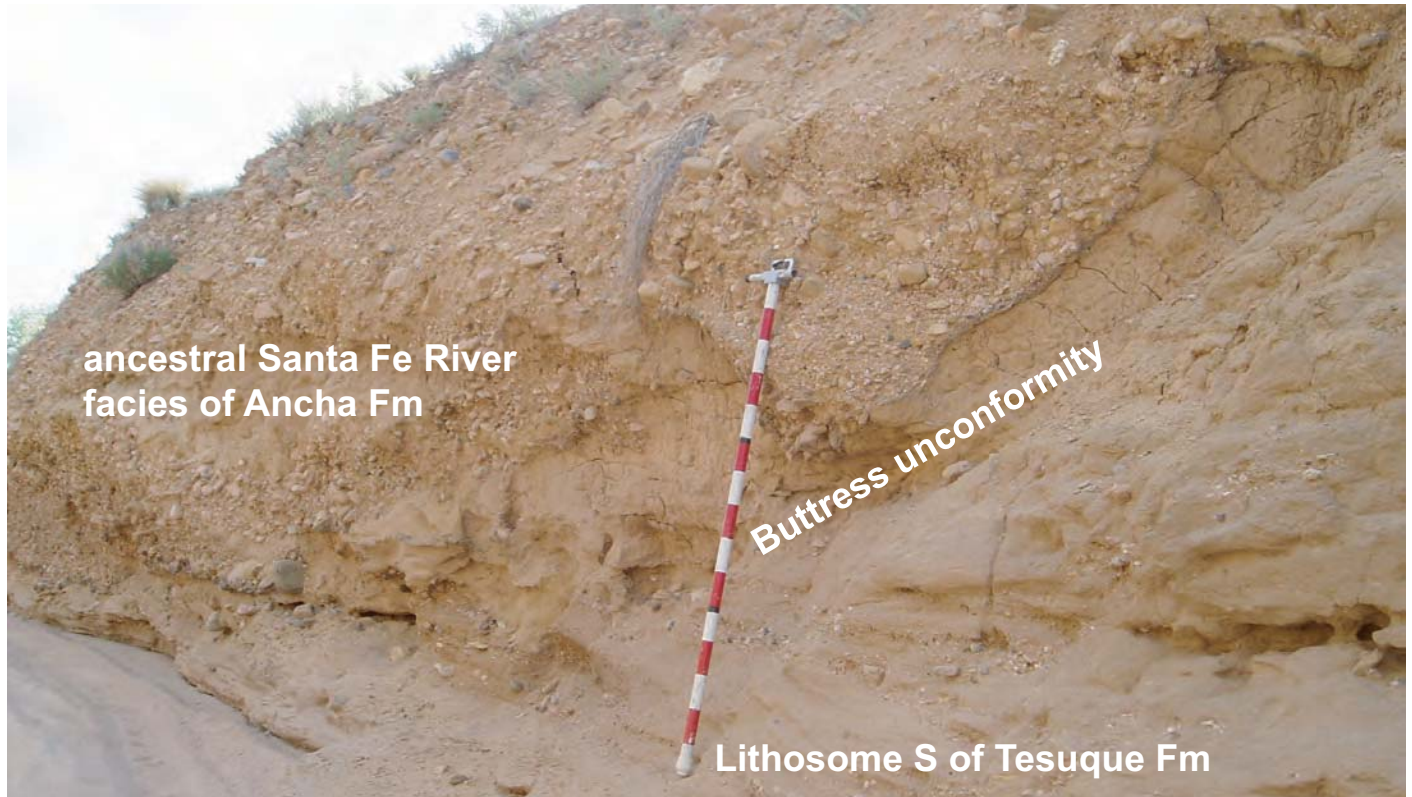


Figure 5. Buttress unconformity under the ancestral Santa Fe River facies of the Ancha Formation at site SEO-A78 (also refer to associated stratigraphic section in Appendix 1). UTM coordinates: 3945850N, 0406730E (NAD 27, zone 13).

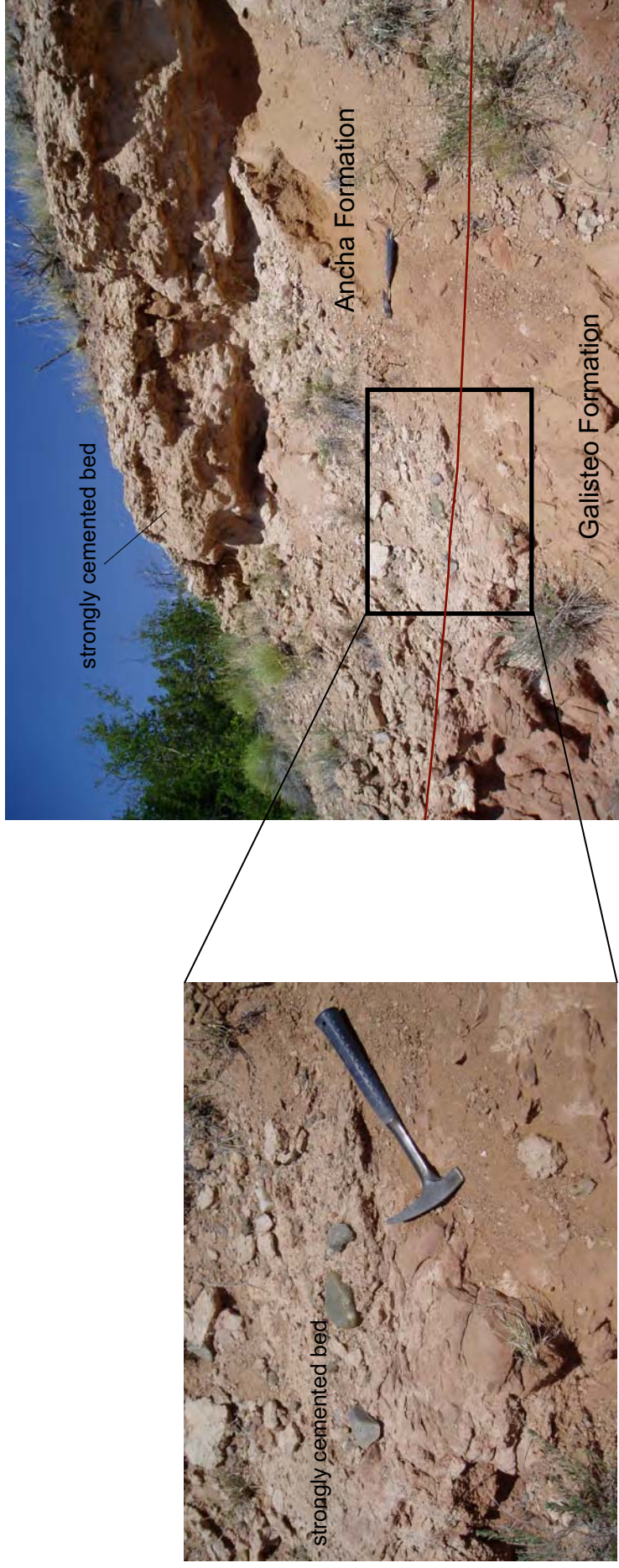


Figure 6. Unconformable contact between the Ancha Fm (top) and sandstone of the Galisteo Fm (base) in the southern Santa Fe embayment. Note the two strongly cemented beds in the basal Ancha Fm, with a non-cemented interval between them. Located at base of stratigraphic section SEO_A80 (west side of Highway 14 about 1.78 km southwest of San Marcos Spring; UTM coordinates: 3923262 N, 402244E, NAD 27, zone 13).

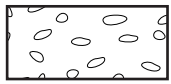


Figure 7. Strongly cemented zone at base of Ancha Formation in the southeastern Santa Fe embayment. Here, it is 7-8 m thick. Location is in lower part of Galisteo 02 stratigraphic section (UTM coordinates of base of section: 3927340 N, 417770 E (NAD 83, zone 13)). This basal cemented zone is also observed in well data at Eldorado 5-6 km to the north.

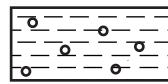
APPENDIX 1

STRATIGRAPHIC SECTIONS

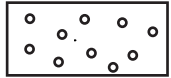
EXPLANATION FOR STRATIGRAPHIC SECTIONS



Sandy gravel



Silt, mud, or very fine to fine sand with minor pebbles



Gravelly sand



Pumiceous (lapilli-size) sand



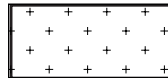
Medium to very coarse sand



Muddy-sandy phreatomagmatic deposit



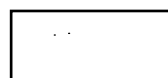
Very fine to medium sand



Ashy sand



Clay-silt and very fine to fine sand



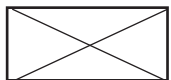
Cross-stratified sand



Clay-silt



Strongly cemented



No exposure



Soil

Miscellaneous abbreviations

ang = angular

subang = subangular

subrnd = subrounded

rnd = rounded

eff = effervescence

lt = light

Sand max grain diameter abbreviations

vfL = very fine-lower; 62-88 μ

vfU = very fine-upper; 88-125 μ

fL = fine-lower; 125-177 μ

fU = fine-upper; 177-250 μ

mL = medium-lower; 250-350 μ

mU = medium-upper; 350-500 μ

cL = coarse-lower; 500-710 μ

cU = coarse-upper; 710-1000 μ

vcL = very coarse-lower; 1000-1410 μ

vcU = very coarse-upper; 1410-2000 μ

Pebble max diameter abbreviations

vf = very fine; 2-4 mm

f = fine; 4-8 mm

m = medium; 8-16 mm

c = coarse; 16-32 mm

vc = very coarse; 32-64 mm

Cobble max diameter abbreviations

f = fine= 64-128 mm

c = coarse; 128-256 mm

Soil development abbreviations (from Birkeland et al., 1991, and Birkeland, 1999)

Soil structure: 1 = weak, 2 = moderate, 3 = strong; f = fine (0.5-1 cm), m = moderate (1-2 cm), c = coarse (2-5), vc = very coarse (>5 cm); sbk = subangular blocky, abk = angular blocky.

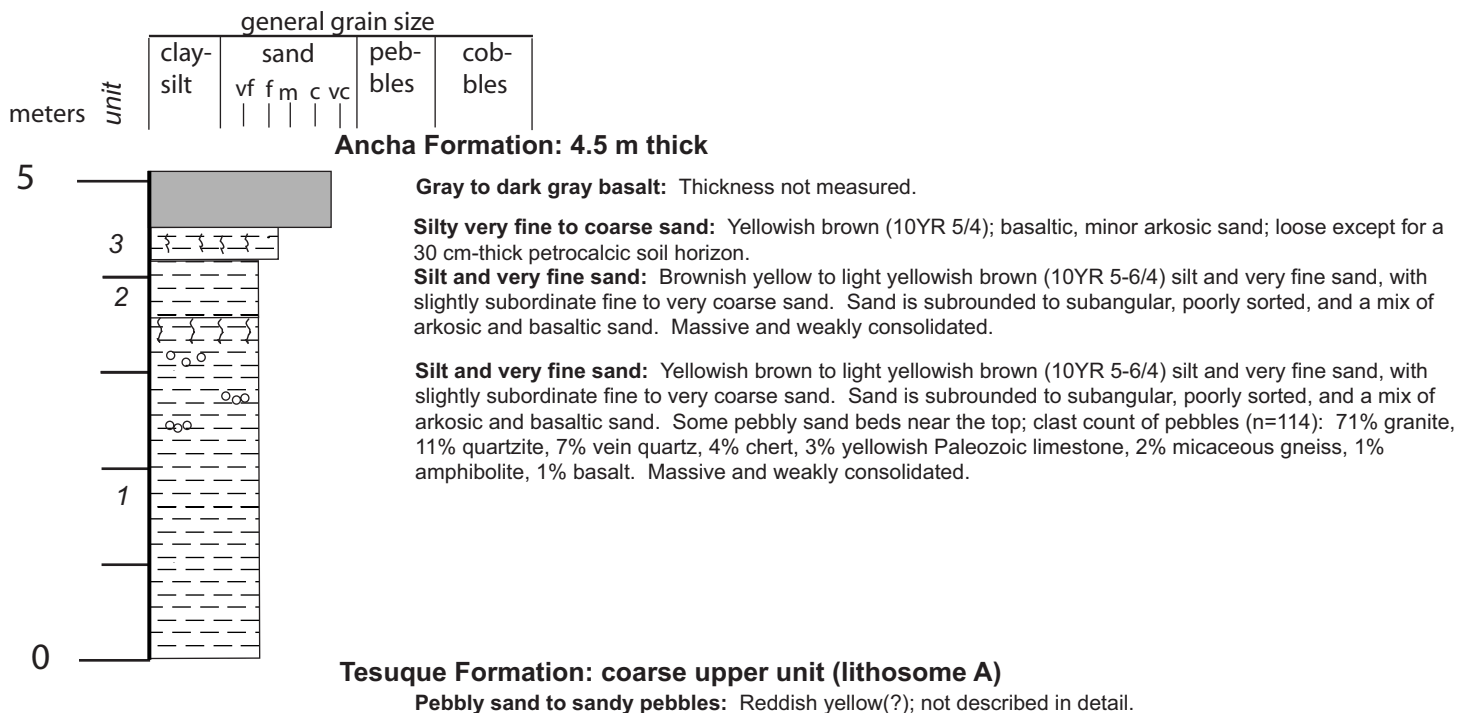
Clay films: v1 = <5% of total surface area, 1 = 5-25% of total surface area, 2 = 25-50% of total surface area, 3 = >50% of total surface area; f = faint, d = distinct, p = prominent; pf = clay films on ped faces, po = clay films line tubular or interstitial pores, br = clay bridges, co = colloid coats mineral grains, cobr = coats and bridges are present.

Soil horizon designations follow those of the Soil Survey Staff (1992) and Birkeland (1999).

Staff in photos is 1.5 m long, with red-white color increments of 10 cm each.

WA-81

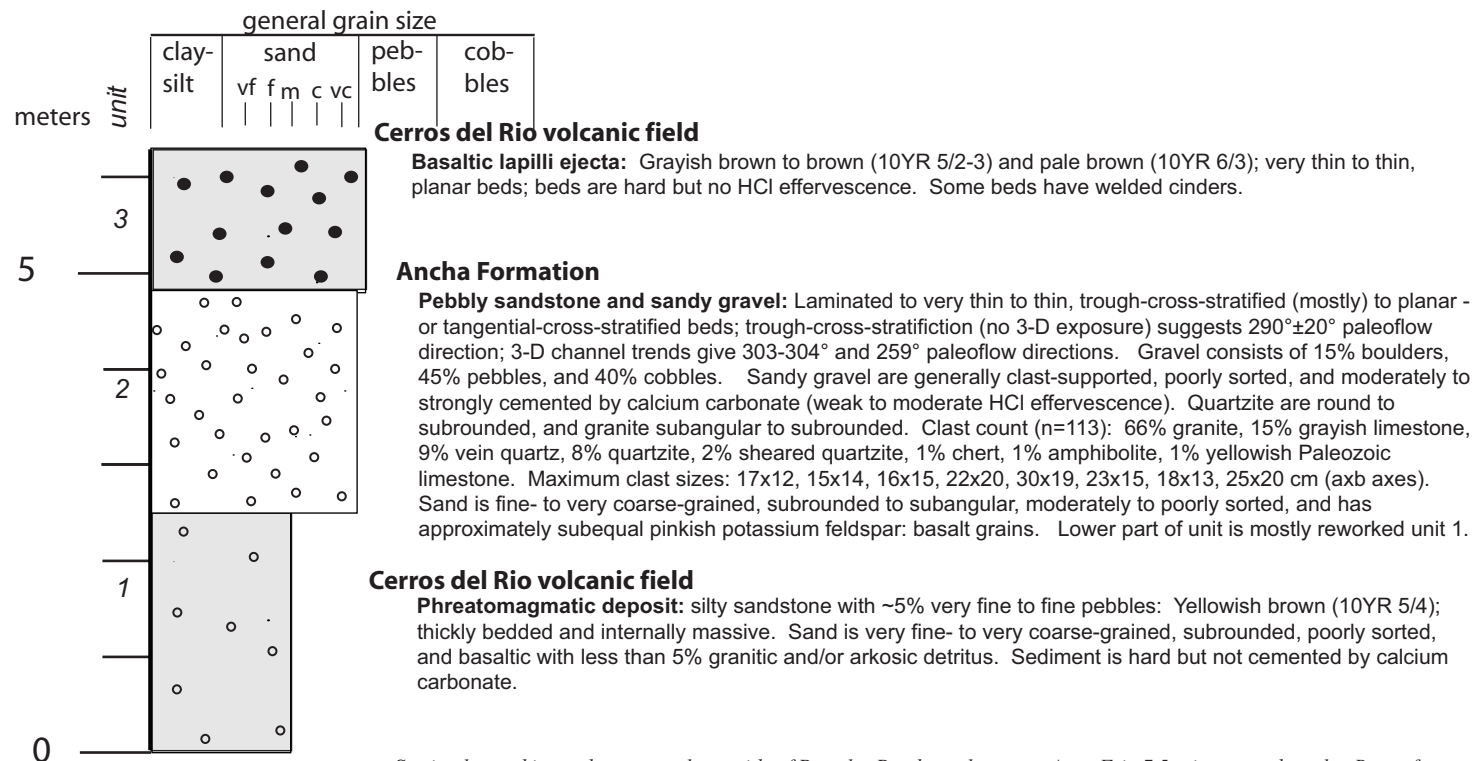
alluvial slope, 1-30% coarse channels



Located on western slope of Canada Ancha, about 580 m north of Portales Pond, in the Agua Fria 7.5-minute quadrangle, New Mexico. Base of section UTM coord: 3952460N, 400760E (NAD 27, zone 13). Measured and described by Daniel Koning, October 9, 2004.

WA-83

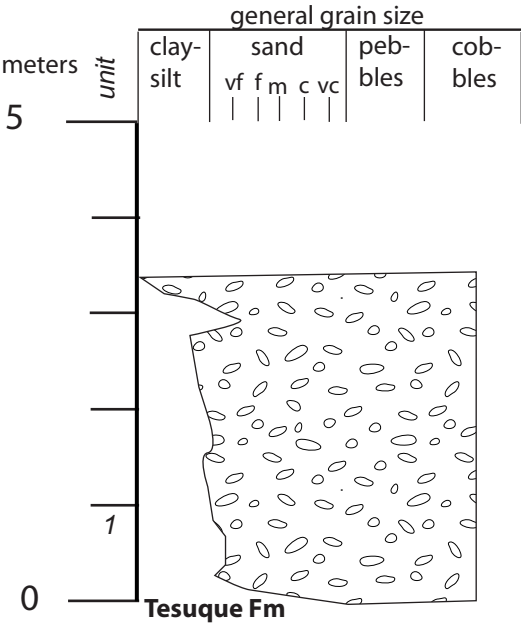
alluvial slope, 1-30% coarse channels



Section located in roadcut on northeast side of Portales Pond, northwestern Agua Fria 7.5-minute quadrangle. Base of section UTM coord: 3951950N, 400660E (NAD 27, zone 13). Measured and described by Daniel Koning, Oct. 9, 2004.

SOUTHWEST

NORTHEAST



Ancha Formation

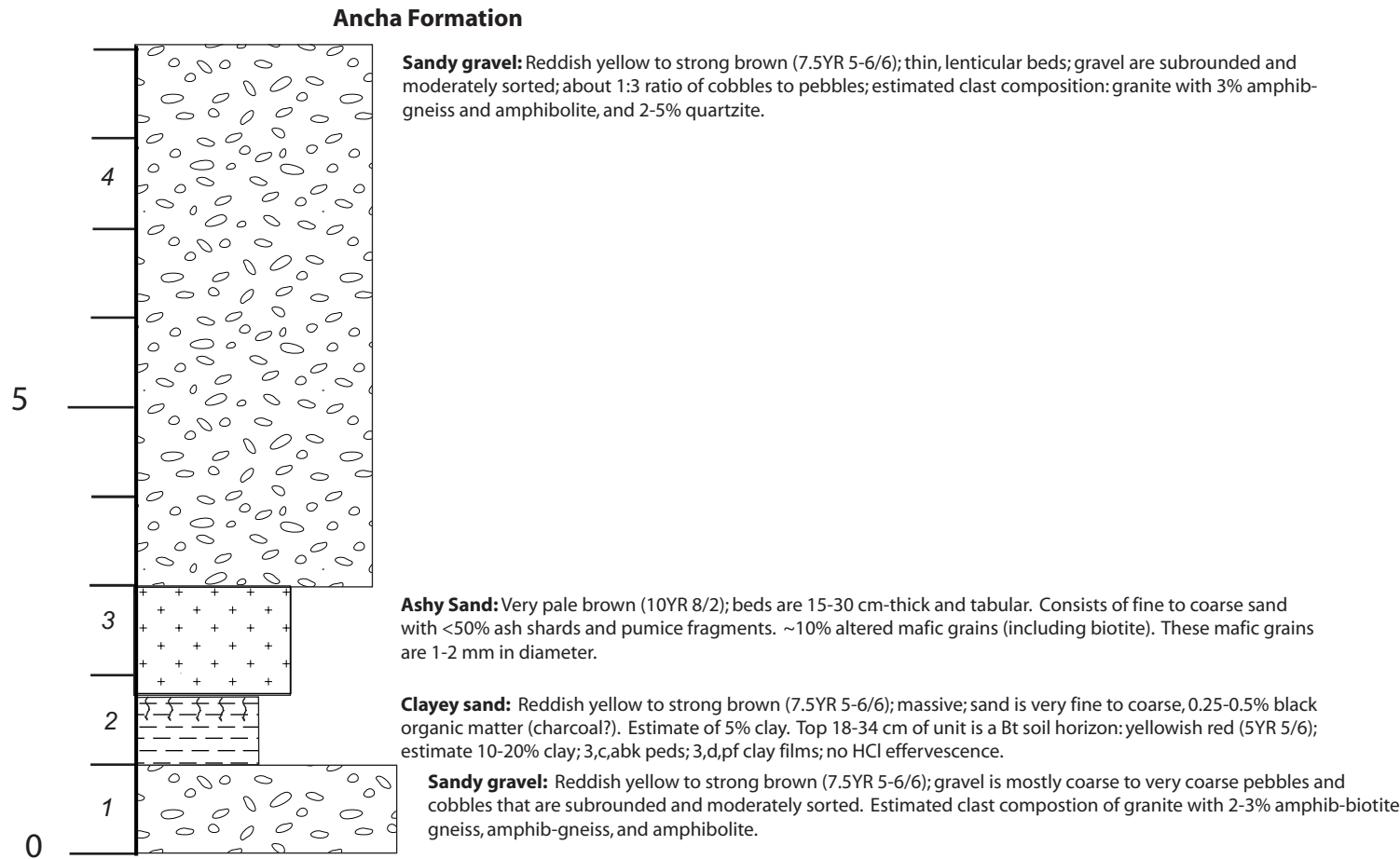
Sandy gravel: Clast-supported sandy gravel, with an estimated 50-60% cobbles and 40-50% pebbles. Beds are very thin to 85 cm-thick, and lenticular to broadly lenticular. Clasts are subrounded (some round) and poorly sorted. Clast count (n=158): 95% granite, 3% quartzite, 1% vein quartz, 1% biotite gneiss, 1% amphibolite (deeply weathered). Max clast sizes: 22x17, 21x16, 17x15, 18x12, 27x17, 20x19 cm (axb axes). Sand is light yellowish brown to brownish yellow (10YR 6/5-6), fU-vcU (mostly cL to vcU), subrounded to subangular, poorly sorted, arkosic and rich in coarse granitic detritus. Non-cemented and weakly consolidated.

Located in modern Santa Fe River, 700 m southwest of bridge of County Rd 62, southwestern Agua Fria 7.5-minute quadrangle, New Mexico. Base of section UTM coord: 3945850N, 406730E (NAD 27, zone 13). Measured and described by Daniel Koning, June 16, 2004.

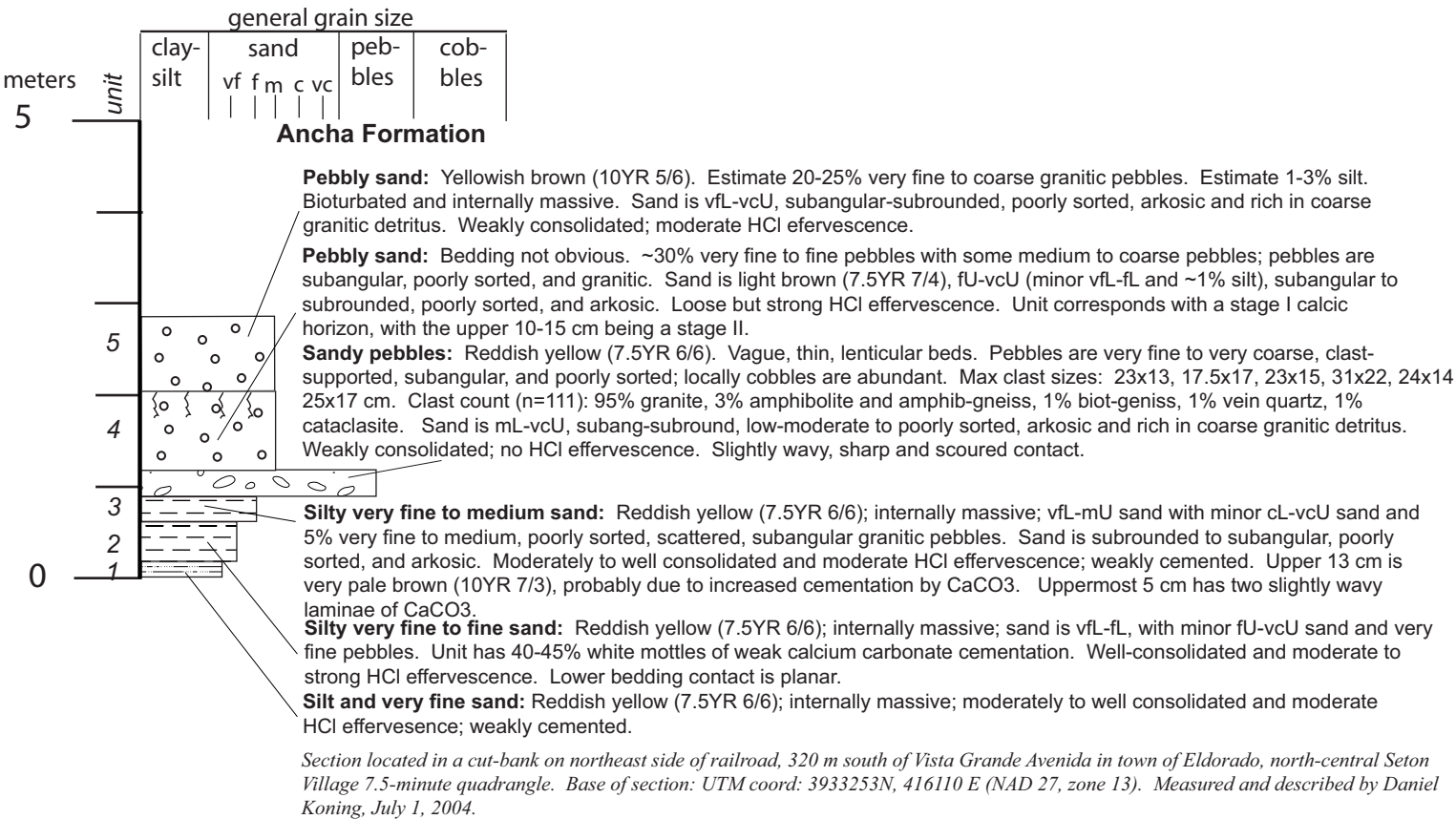
T-36

ancestral Santa Fe River

		general grain size									
meters	unit	clay-silt	sand					pebbles	cob- bles		
			vf	f	m	c	vc				



Located on north slope of Arroyo de los Chamisos in north-central Turquoise Hill, 7.5-minute quadrangle; outcrop is just above the aquaduct and 1.2 km southwest of Capital High School. Base of section: 3942003N, 404416E (NAD 27, zone 13). Measured and described by Daniel Koning, July 26, 1999.

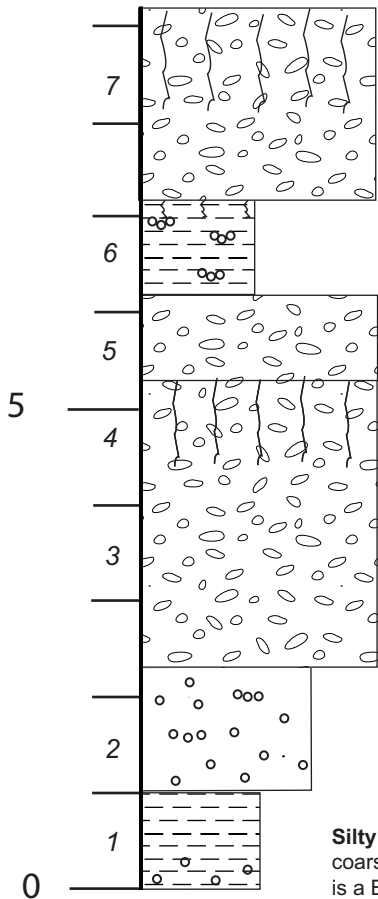


alluvial slope, >35% cs channels



		general grain size								
meters	unit	clay-silt		sand				pebbles	cob-bles	
				vf	f	m	c			

Ancha Formation



Sandy pebbles with about 40% pebbly sand: Sandy pebbles are in very thin, broadly lenticular beds; pebbly sand is planar-laminated. Pebbles are mostly clast-supported, subangular(mostly) to subrounded, moderately to poorly sorted, and granitic. Sand is light yellowish brown to brownish yellow (10YR 6/4-6), fU-vcU, subangular (mostly) to subrounded, poorly (mostly) to moderately sorted, arkosic and rich in coarse granitic detritus. Loose to weakly consolidated. Scoured lower contact with 10-20 cm of relief. Upper 100 cm is a degraded and bioturbated stage II+ calcic horizon; internally massive and very pale brown (10YR 7/3). Using a dip of 1°, the Guaje pumice projects to upper 2 m of exposure.

Silty sand with minor sandy pebble lenses: Generally a silty vFL-fL sand with 10-15% very thin to thin lenses of sandy pebbles; pebbles are very fine to coarse, subangular to subrounded, poorly sorted, and granitic. Internally massive aside from sandy pebble lenses. Sand is has subordinate fU-vcU grains, is subrounded-subangular, poorly sorted, and arkosic. Moderately-well consolidated and moderate HCl effervescence. Lower contact is planar and gradational over 20 cm. Upper 10-15 cm corresponds to a weak stage II calcic horizon.

Sandy pebbles: Very thin to medium, lenticular to broadly lenticular beds. Gravel is clast-supported, subangular to subrounded, and poorly sorted; composition similar to unit 3 but with about 5% amphibolite. Sand is very pale brown (10YR 7/3), mL-vcU, subangular, poorly sorted, arkosic and rich in coarse granitic detritus. Loose to weakly consolidated; no HCl effervescence.

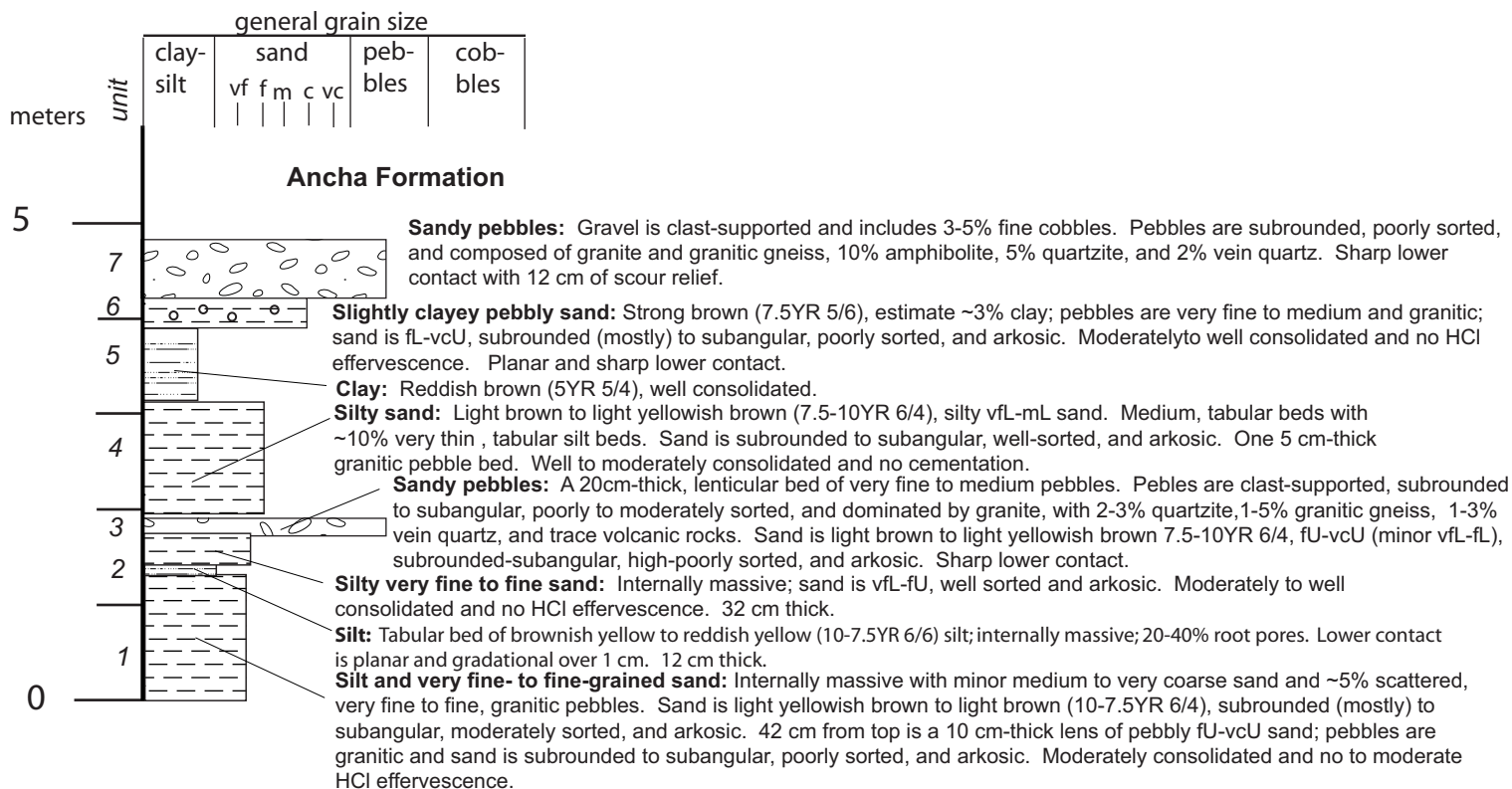
Subequal pebbly sand and sandy pebbles: Pebbly sand has planar-laminated to planar-lenticular, very thin beds; sand is light brown to reddish yellow (7.5YR 4-6/6) and strong brown (7.5YR 5/6); mU-vcU, subangular (mostly) to subrounded, moderately to poorly sorted, and arkosic. Sandy pebbles are clast-supported and in very thin to thin, lenticular beds with minor channel-shaped beds 10-40 cm-thick. Gravel has ~20% cobbles; clasts are subangular to subrounded and poorly sorted; clast count (n=103): 87% granite, 5% biot-gneiss, 4% amphib-biot gneiss and amphib, 2% vein quartz, 1% quartzite, 1% musc-biot gneiss; max clast sizes: 24x17, 12x10, 16x9.5, 15x12, 12x11, 12x7, 12x8 cm (axb axes). Loose to weakly consolidated and no HCl effervescence. Upper 90 cm has a stage III carbonate horizon: white (2.5Y 8/1), more cemented and hard near top, strong HCl effervescence. Lower contact not well-exposed. Unit is approx the same elevation as the Guaje pumice on north wall of Arroyo Hondo (a few 100 m to west).

Gravelly sand: Reddish yellow (7.5YR 6-7/6); massive except for the pebbly lenses described below. 15-20% scattered, very fine to very coarse pebbles that are subangular, poorly sorted, and dominated by granite and granitic gneiss; 10% of gravel are cobbles. 10-20% of unit are very thin to medium lenses of clast-supported, sandy pebbles: very fine to very coarse, subangular to subrounded, and poorly sorted. Sand is vFL-vcU, silty (less silt than underlying unit), and subangular (mostly) to subrounded, poorly sorted, and arkosic. Moderate consolidated with no HCl efferevescence. Lower contact is sharp and scoured (4 cm of relief).

Silty sand: Strong brown to reddish yellow (7.5YR 5-6/6). Unit represents one fining-upward bed. 7-10% very fine to coarse granitic pebbles near base. Sand is vFL-mU, subrounded to subangular, poorly sorted, and arkosic. Upper 11 cm is a Bt soil horizon: 2dpf clay films and 3vf-m, subang blocky, hard peds; reddish yellow (5YR 6/6); lower contact is planar and gradational over 2 cm. Well consolidated, no to weak HCl effervescence. Base of exposure is 1.5 m below that of the QTa/Tt contact on north side of Arroyo Hondo (bearing of N2°W).

Located in first cut-bank along railroad south of Arroyo Hondo, northwest Seton Village 7.5-minute quadrangle. Base of section UTM coord: 3942106N, 409461E (NAD 27, zone 13). Measured and described by Daniel Koning, July 1, 2004.

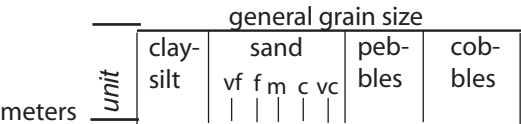
alluvial slope, 15-60% cs channels



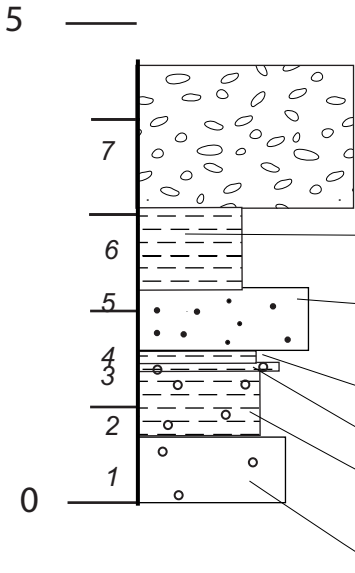
Section measured along west side of Highway 14 about 1.58 km north of San Marcos Spring; northwest Picture Rock quadrangle. Base of section UTM coord: 3925915N, 403036E (NAD 27, zone 13). Measured and described by Daniel Koning, June 17, 2004.

SEO-A125

alluvial slope, 15-40% coarse channels



Ancha Formation



Note: Upper 30 cm on top of unit 5 is a silty vFL-fU sand with minor mL-vcU sand and pebbles. Disturbed because of pedogenesis and surface activity. Max clast sizes of surface gravel: 16x10, 12x9, 13x9.5, 16x6, 20x15, 21x14 cm (axb axes).

Sandy pebbles with about 1/3 pebbly sand: Sandy pebbles are in very thin to thin, lenticular beds; also tangential to planar cross-stratification (laminations and very thin beds) up to 12 cm tall. Pebbles are clast-supported, very fine to fine, subrounded to subangular, and moderately to poorly sorted; max clast size: 5x3.5, 4.5x3, 4.5x4, 6x4.5, 5x4.5 cm (axb axes); clast count (n=104): 88% granite, 5% amphib-gneiss, 2% amphibolite, 2% biot-gneiss, 3% vein quartz. Pebbly sand is planar-laminated or low-angle-cross-stratified (up to 6 cm-thick). Sand is light yellowish brown (10YR 6/4), fU-vcU, moderately sorted within a bed, subangular (minor subrounded), and arkosic. Moderately to weakly consolidated with weak HCl effervescence.

Silty very fine-grained sand: Pink (7.5YR 7/4) with subordinate fL-cl sand. Internally massive. Sand is subrounded to subangular, moderately sorted, and arkosic. Contact not readily visible but probably planar and gradational over 10-20 cm. Well consolidated with moderate-strong HCl effervescence.

Pumiceous sand: Light yellowish brown to very pale brown (10YR 6-7/4) and internally massive. Sand is vFL-fU with minor mL-vcU; subrounded to subangular, moderately sorted, and arkosic. Sand is mixed with subordinate pumice that is sand-size to very fine to medium pebble-size. Moderately consolidated with weak HCl effervescence. Sharp, planar lower contact.

Clayey sand: Strong brown (7.5YR 5/6) clayey (est 3-5% clay) vFL-mU sand; sand is subround to subang, moderately sorted, and arkosic. Moderately consolid. with weak HCl effervescence. Lower contact is planar(?) and gradational.

Muddy sand: Est 3-5% mud. Sand is vFL-vcU, subang-subrnd, poorly sorted, and arkosic. 2-3% very fine to medium pebbles (granitic, subang-subrnd, poorly sorted). Lower contact is slightly wavy and gradational over 2 cm. 5 cm thick.

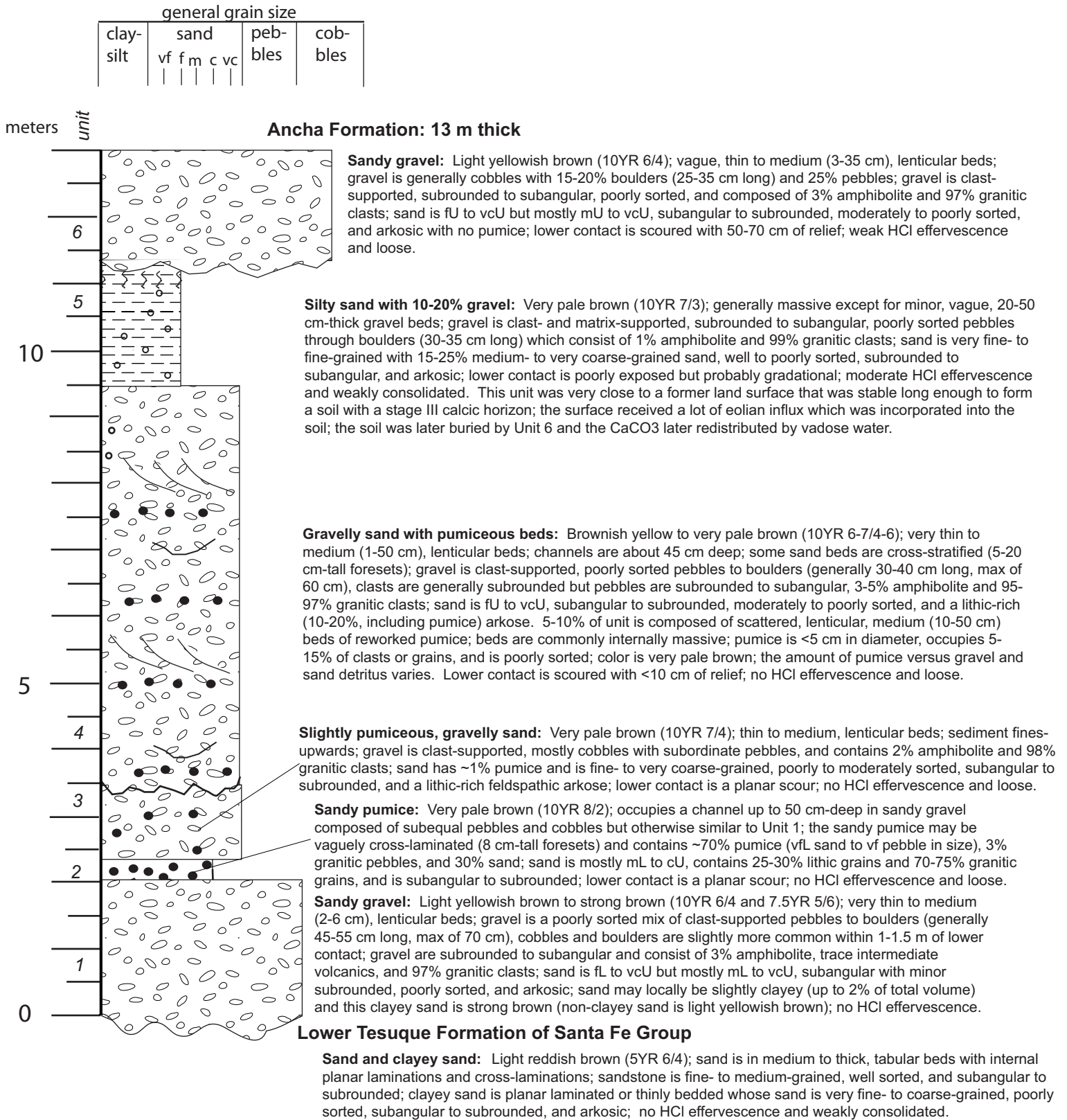
Silty sand with minor pebbles: Light yellowish brown (10YR 6/4) vFL-mU sand with minor cL-vcU sand and 1% very fine to fine pebbles. Internally massive. Estimate 3-5% silt. Sand is subrounded to subangular, moderately to poorly sorted, and arkosic. Moderately consolidated with moderate HCl effervescence. Sharp, planar(?) lower contact.

Sand with minor pebbles: Light yellowish brown (10YR 6/4), very slightly silty (est 1% silt) vFL-vcU sand. Internally massive. Sand is subangular with some subrounded, moderately to poorly sorted, and arkosic; 1-5% very fine to fine, subangular granitic pebbles. Moderately consolidated with weak HCl effervescence. Locally, upper 3-5 cm is strongly cemented by CaCO3 -- often in wavy laminations.

Located in road-cut on east side of Richards Avenue, a short distance (approximately 30 m) north of the southern Interstate 25 overpass, northwest Seton Village 7.5-minute quadrangle. Base of section UTM coord: 3942108N, 409461E (NAD 27, zone 13). Measured and described by Daniel Koning, July 3, 2004.

Arroyo Hondo

alluvial slope, >35% coarse channels

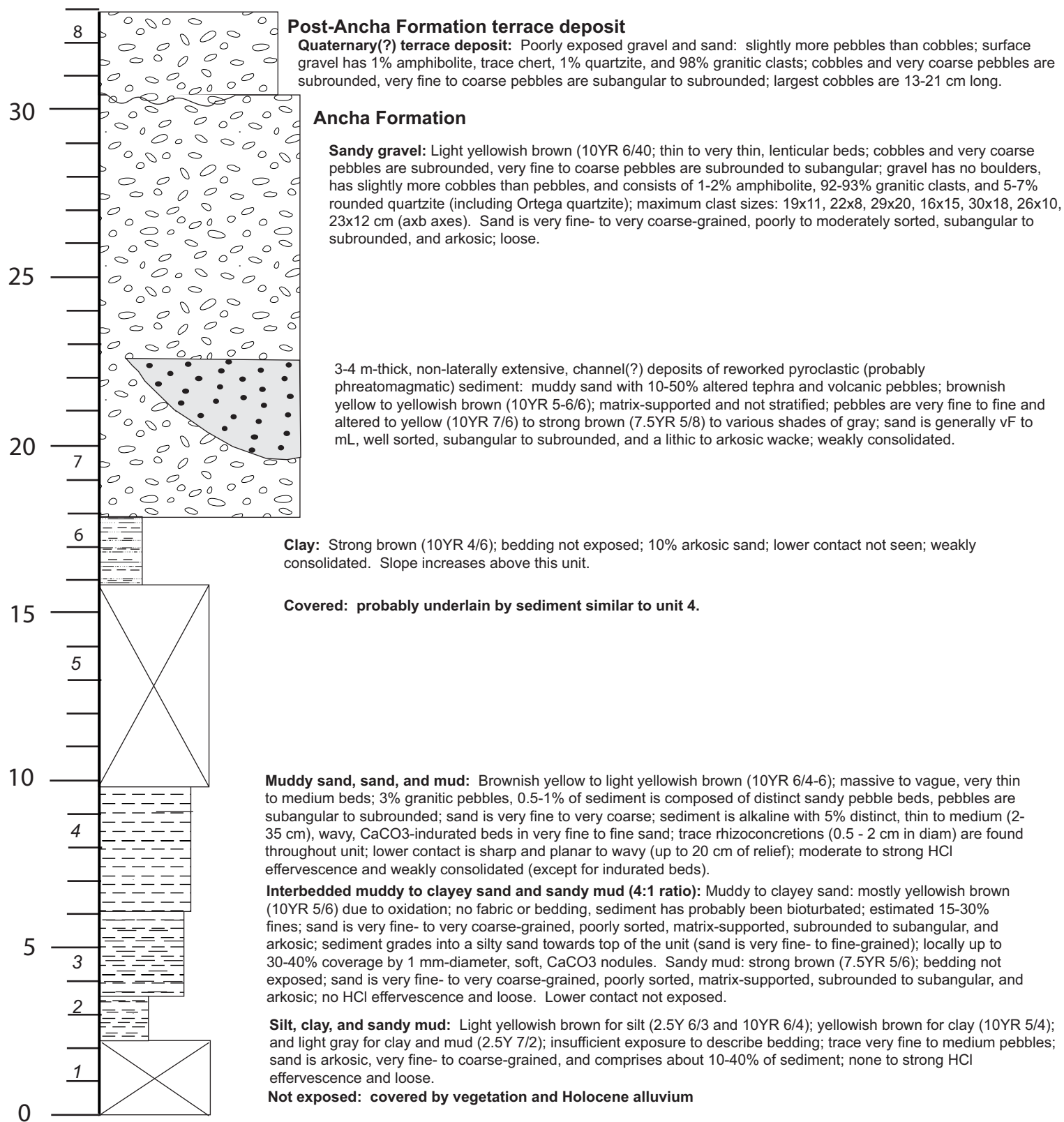


Located on upper north slope of Arroyo Hondo near south terminus of Old Pecos Trail, Santa Fe 7.5-minute quadrangle, Santa Fe County, New Mexico. Measured and described upsection from unit 1, along N71E trend, by D.J. Koning October 16, 2000. Base of measured section at N: 3,942,778 m; E: 414,823 m (UTM zone 13, NAD 27). Section is on private property of Gaylon Duke (46 Old Agua Fria Road W., Santa Fe, N.M., 87505).

Cienega Creek

ancestral Santa Fe River (18-30.5 m) and alluvial slope, 1-30% coarse channels (0-18m)

meters	unit	general grain size					pebbles	cob- bles
		clay- silt	vf	f	m	c	vc	



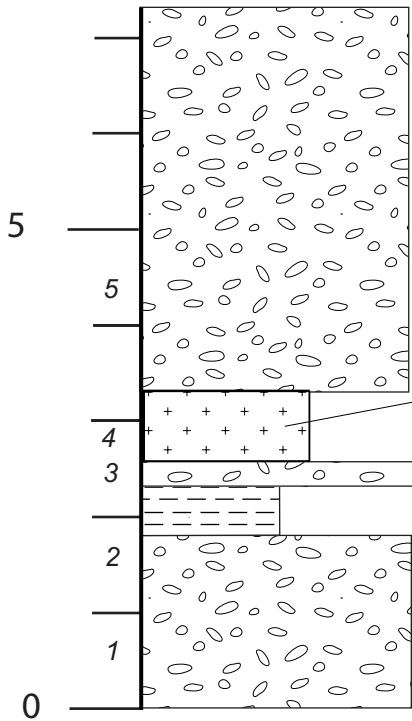
Located on north slope of La Cienega Creek, approximately 550 m west of Arroyo de los Tanques, Turquoise Hill 7.5-minute quadrangle, Santa Fe County, New Mexico. Measured and described upsection from unit 1 by D.J. Koning on October 17, 2000. Section trends N15E for 290 m, then N-S for 335 m. Base of measured section located at bottom of La Cienega Creek at N: 3,937,427 m; E: 400,840 m (UTM zone 13, NAD 27).

T-27
ancestral Santa Fe River



		general grain size								
meters	unit	clay-silt	sand				pebbles	cob- bles		
			vf	f	m	c				

Ancha Formation



Sandy gravel: Bedding not well-exposed. Gravel consists of pebbles and slightly subordinate cobbles that are subrounded and moderately sorted; maximum clast size: 20x11, 16x9, 17x10.5, 10x9, 11x9, 14x9, 14x9 cm (axb axes); clast count (n=137): 95% granite, 1.5% vein quartz, 2.5% biotite- and muscovite-gneiss, 2% amphib-biot gneiss and amphibolite. Sand is light yellowish brown (10YR6/4), slightly silty (est 1% silt), vFL-vcU but mostly mL-vcU), subangular to subrounded, poorly sorted, and arkosic. Weakly consolidated and no HCl effervescence.

Sandy ash to ashy sand: Very pale brown (10YR 8/2). Ash includes glass, pumice shards, quartz crystals, and dark volcanic lithic grains and is fL-cU (lithic grains are mL-vcU and vf pebble-size). Unit includes subequal arkosic, detrital sand that is generally mU-vcL, subrounded to subangular, and moderately sorted. Weakly consolidated and no to little HCl effervescence. Ash seems to project and correlate to the ash dated at 1.63± 0.02 Ma in Koning et al. (2002).

Sandy gravel: As in unit 5.

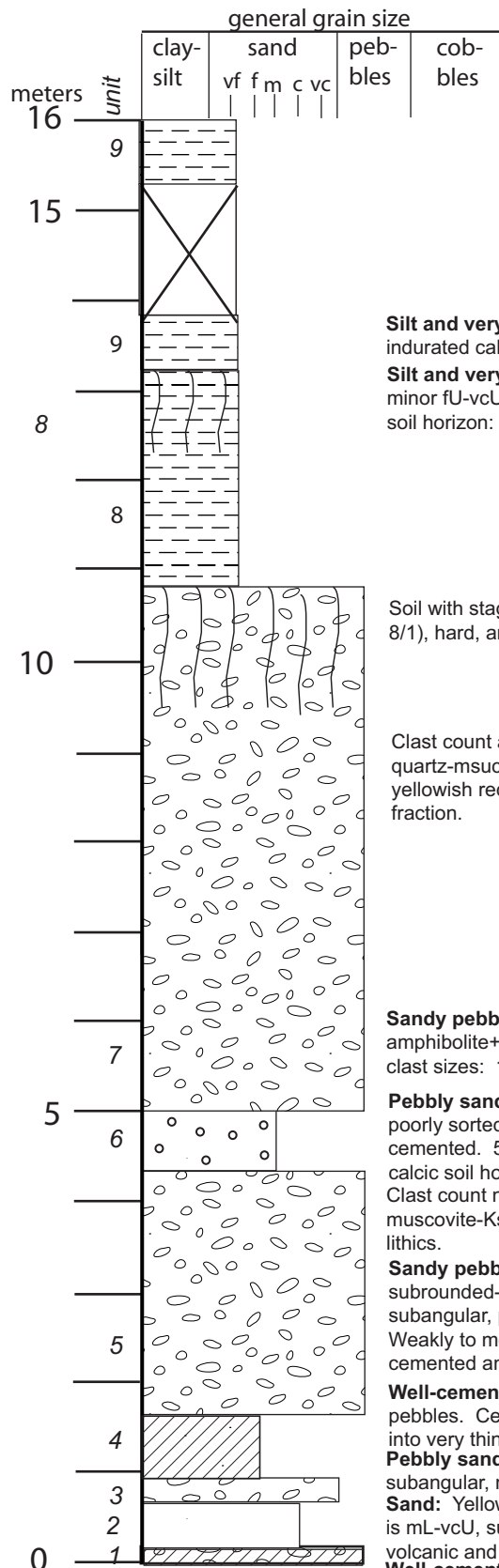
Sandy clay: Strong brown (7.5YR 5/6); sand is generally vFL-vfU. Moderately consolidated; no HCl

Clayey-sandy gravel: Clay and sand color: strong brown (7.5YR 5/6). Gravel is subrounded and moderately sorted; max clast sizes: 16x9, 18x9, 21x11, 15x9, 14x9, 12x8 cm (axb axes); clast count (92% granite, 3% amphibolite, 3% amphib-gneiss, 1% un-identified, 1% quartzite. Sand is mL-vcU (mostly cL-vcU), subrounded to subangular, poorly sorted, and arkosic. Weakly consolidated; no HCl effervescence.

Section measured in cut-bank on east side of Highway 599, where the highway cuts through the south wall of Arroyo de los Chamisos, SW of center of the Turquoise Hill 7.5-minute quadrangle. Base of section UTM coord: 3940902N, 403012E (NAD 27, zone 13). Measured and described by Daniel Koning, July 26, 1999.

SEO-A80

alluvial slope, 25-60% coarse channels



Silt and very fine to fine sand: Sediment similar to unit 8; basal 6 cm has abundant white, cm-size rip-ups of indurated calcium carbonate (probably from erosion of underlying soil).

Silt and very fine to fine sand: Light brown to light yellowish brown (7.5-10YR 6/4); sand is vfl-fL and mixed with minor fU-vcU sand. Massive. Well-consolidated and has strong HCl effervescence. Top 90 cm is a stage III+ calcic soil horizon: very pale brown (10YR 8/2), strong HCl effervescence.

Soil with stage III calcic horizon developed on parent material of pebbly sand; internally massive, white color (7.5YR 8/1), hard, and strong HCl effervescence. Lower horizon conatct gradational over 10 cm.

Clast count a few m below top of unit: 52% granite, 43% latite and other felsic-intermediate volcanic clasts, 4% quartz-msucovite gneiss, 1% quartz-biotite gneiss, 1% muscovite-Kspar gneiss. Gravel are poorly sorted. Sand is yellowish red (5YR 4-5/6, mL-vcU, has 30%-35% sand-size clay bits; more latite-rich lithics than Kspar in the sand fraction.

Sandy pebbles: Similar to sediment in unit 5. Clast count (n=123): 69% granite, 13% latite, 5% amphibolite+amphibolite gneiss, 5% quartzite, 4% granitic gness, 3% vein quartz. 1% muscovite schist. Maximum clast sizes: 13x8, 15x8.5, 19x18, 16x10, 16x9, 14x13, cm (axb axes).

Pebbly sand: Light brown (7.5YR 6/4) pebbly very fine to very coarse sand. Sand is subrounded to subangular, poorly sorted, and has an approximate ratio of 50-60%:50-50% Kspar:volcanic and mafic lithic grains. Loose and not cemented. 5 m to the north, the unit becomes silty and has strong effervescence in HCl; here, it looks like a stage II+ calcic soil horizon.

Clast count near top of unit (n=100): 50% intermediate to felsic volcanic clasts, 45% granite, 2% vein quartz, 2% muscovite-Kspar gneiss, and 1% quartzite. Sand has a ratio of about 50-60% : 40-50% kspar: volcanic and mafic lithics.

Sandy pebbles: Very thin to medium, lenticular to broadly lenticular beds. Pebbles are mostly clast-supported; subrounded-subangular and poorly to moderately sorted. Sand is light brown (7.5YR 6/4), fL-vcU, subrounded to subangular, poorly to moderately sorted, and arkosic; 3-5% clay present as films coating 10-20% of grain surface. Weakly to moderately consolidated and no HCl effervescence -- except for lower 22 cm of unit, which is well-cemented and weathers into very thin to thin plates. Sharp and planar lower contact.

Well-cemented sandstone: Pinkish white to pink (7.5YR 8/2-3); sand is vfl-vcU (mostly vfl-fU), 2-3% scattered pebbles. Cement is hard but not as indurated as unit 1; in lower 30 cm the cement is massive, but above it weathers into very thin to thin plates. Lower conatct is gradational over 4 cm.

Pebbly sand to sandy pebbles; Sand and bedding style similar to underlying unit; pebbles are subrounded to subangular, moderately sorted, and granitic with subordinate volcanic and gneiss clasts. Loose and non-cemented.

Sand: Yellowish red to reddish yellow (5YR 5-6/6); very thin to medium(?), lenticular to broadly lenticular beds; sand is mL-vcU, subrounded to subangular, moderately sorted, and has an estimated 50-60%:40-50% ratio of Kspar : volcanic and mafic lithics. Loose and non-cemented. Sharp lower contact.

Well-cemented pebble conglomerate: Pink (7.5YR 8/3), pebble conglomerate; pebbles are very fine to very coarse, with very sparse fine cobbles, subrounded-subangular, and poorly sorted; clasts are granitic with an estimated 5% quartzite and 25-35% intermediate-felsic volcanic rocks. Very hard due to strong calcium carbonate cementation. Basal contact is an unconformity and relatively planar.

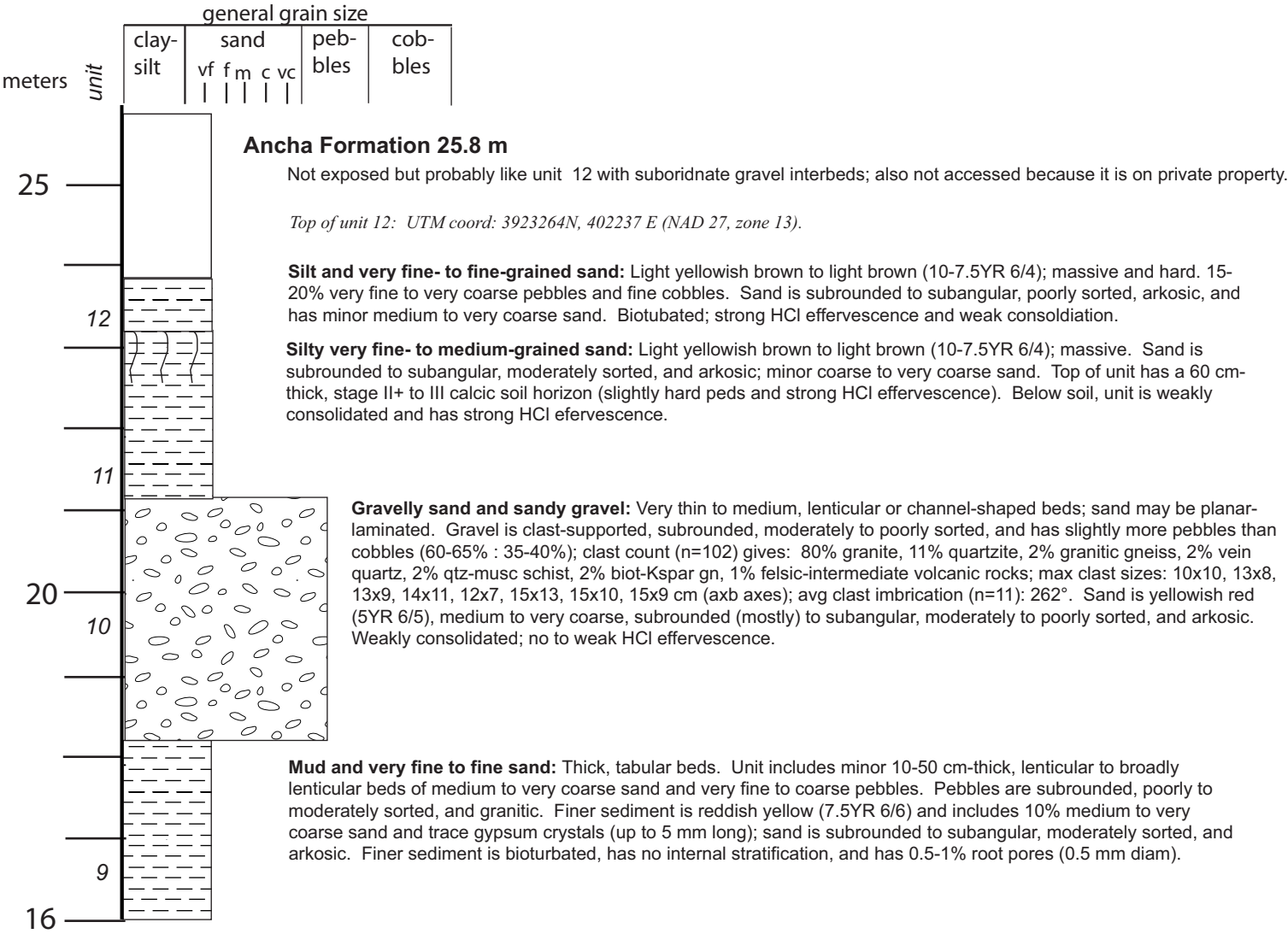
Galisteo Formation

Sandstone: Steeply tilted fL-fU sand; well-sorted, arkosic, and cemented (hard; strong HCl effervescence).

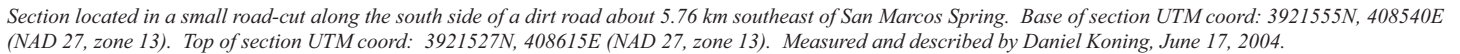
Section measured along west side of Highway 14 about 1.78 km southwest of San Marcos Spring; northwest Picture Rock quadrangle, New Mexico. Base of section UTM coord: 3922778N, 402219 E. Top of section UTM coord: 3923262, 402244E (NAD 27, zone 13). Measured and described by Daniel Koning, June 17, 2004.

SEO-A80 (continued)

alluvial slope, 25-60% coarse channels

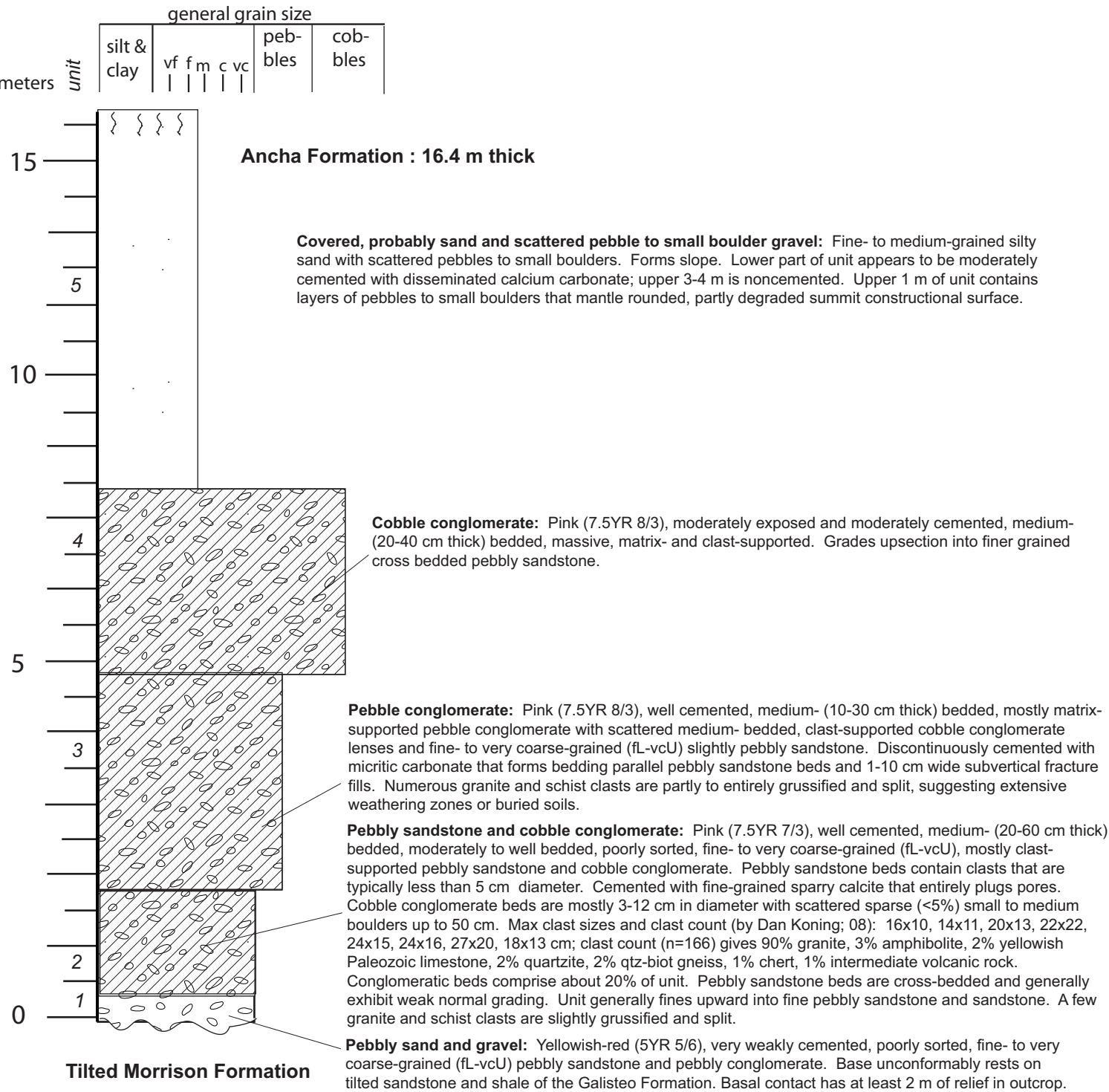


alluvial slope sediment, 25-60% coarse channels



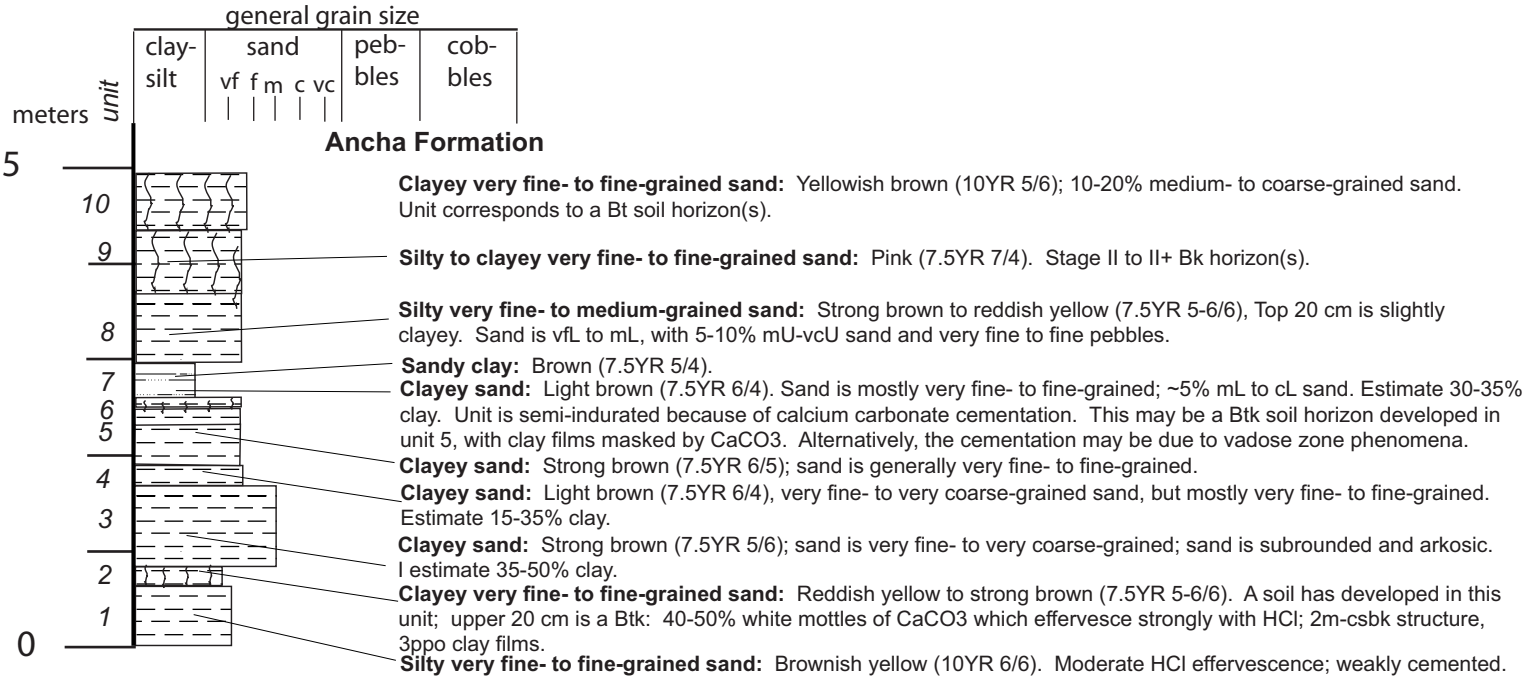
Galisteo 02

alluvial slope, >35% coarse channels



Reference section of the Ancha Formation described on exposures of private dirt road, northwest of intersection of AT&SF shortline to Lamy, NM, and US 285. Described by S.D. Connell and F.J. Pazzaglia on June 9, 2000. Base at N: 3,927,340 m, E: 417,770 m (Zone 13, NAD 83), Galisteo 7.5-minute quadrangle, Santa Fe County, New Mexico.

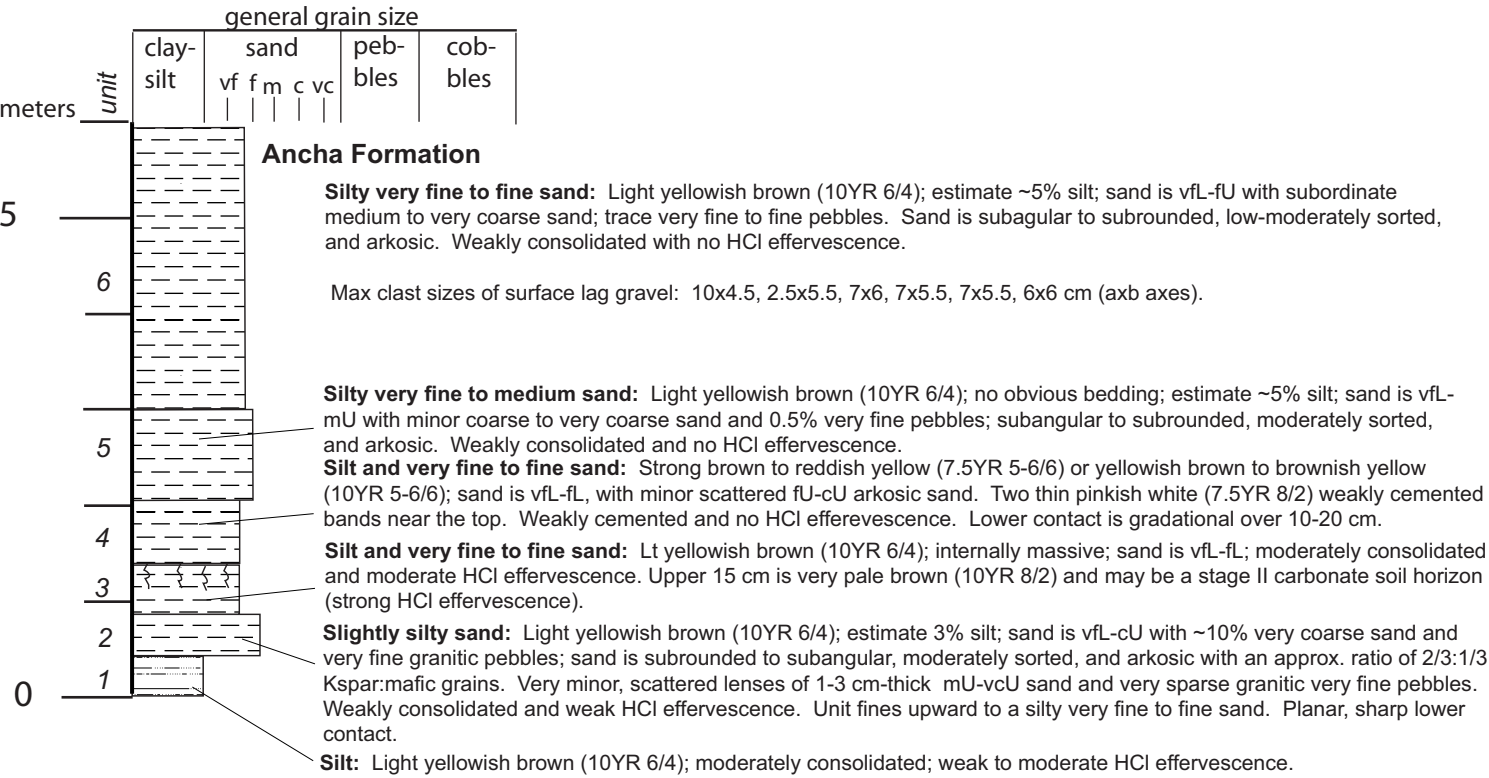
alluvial slope, 1-30% coarse channels



Located in roadcut on east side of Highway 599, about 1 km south of the New Mexico State Penitentiary. Base of section UTM coord: 3934555 N, 404555 E (NAD 27, zone 13). Measured and described by Daniel Koning, August 10, 1999.

SEO-A124

alluvial slope, 1-30% coarse channels



Section located in cut-bank on west side of Interstate 25, immediately southwest of the south-bound on-ramp from Highway 587; southwest Turquoise Hill 7.5-minute quadrangle. Base of section UTM coord: 3934548 N, 398323 E (NAD 27, zone 13). Measured and described by Daniel Koning, July 1, 2004.